

Characterization and trend detection of meteorological drought for a semi-arid area of Parbhani district of Indian state of Maharashtra

S. B. TARATE and PRAVENDRA KUMAR*

*Department of Soil Science and Agriculture Chemistry, School of Agriculture,
Lovely Professional University, Phagwara – 144 411, Punjab, India*

**Department of Soil and Water Conservation Engineering, G. B. Pant University of
Agriculture and Technology, Pantnagar – 263 145, India*

(Received 24 April 2020, Accepted 5 March 2021)

e mail : taratesuryakant01@gmail.com

सार – सूखा, विशेष रूप से शुष्क और अर्ध-शुष्क क्षेत्रों में जल संसाधन प्रबंधन के संबंध में प्रतिकूल प्राकृतिक खतरों में से एक है। वैश्विक उष्णन (ग्लोबल वार्मिंग) और जलवायु परिवर्तन की स्थितियों के कारण, शुष्क तथा अर्द्ध शुष्क क्षेत्रों में सूखे की गंभीरता तथा इसकी प्रवृत्ति का पता लगाना अत्यंत महत्वपूर्ण है। इसलिए, इस अध्ययन में, भारतीय राज्य महाराष्ट्र के परभणी जिले के अर्ध शुष्क क्षेत्र के लिए मानकीकृत वर्षण सूचकांक (SPI) और मानकीकृत आवीक्षण सूखा सूचकांक (RDI_{std}) जैसे विभिन्न मौसम संबंधी सूखे सूचकांकों का उपयोग करके सूखे की प्रचंड गंभीरता का आकलन किया गया। परिणामों से पता चला कि SPI और RDI_{std} समान रूप से व्यवहार करते हैं ताकि सूखे की गंभीरता का पता लगाया जा सके। 37 वर्षों (1983 से 2019) में, SPI में 1 प्रचंड सूखा वर्ष, 6 मध्यम सूखे वर्ष, 22 सामान्य वर्ष, 4 मध्यम आर्द्र वर्ष, 3 अत्यंत आर्द्र वर्ष तथा 1 प्रचंड आर्द्र वर्ष दर्शाए गए जबकि RDI_{std} में 1 प्रचंड सूखा वर्ष, 5 मध्यम सूखे वर्ष, 23 सामान्य वर्ष, 4 मध्यम आर्द्र वर्ष, 3 अत्यंत आर्द्र वर्ष तथा 1 प्रचंड आर्द्र वर्ष दर्शाए गए। SPI और RDI_{std} दोनों सूचकांकों के आधार पर एक प्रचंड सूखा वर्ष देखा गया। साप्ताहिक आँकड़ा विश्लेषण के आधार पर, सूखे, सामान्य और आर्द्र सप्ताह की आवृत्ति क्रमशः 70.58, 15.90 तथा 13.51% पाई गई। अल्पकालिक साप्ताहिक वर्षा विश्लेषण से अध्ययन क्षेत्र में सूखे की लगातार घटनाओं की स्पष्ट तस्वीर का पता चला जो दीर्घावधि पैमाने (मासिक या वार्षिक) की तुलना में अधिक है। सांख्यिकीय रूप से महत्वपूर्ण प्रवृत्ति की पहचान करने के लिए, एक गैर प्राचलिक परीक्षण (Mann-Kendall) और प्राचलिक परीक्षण (Linear regression) का उपयोग किया गया। वार्षिक पैमाने पर, PET और तापमान को छोड़कर वर्षा, SPI, RDI_{std} के मामले में महत्वपूर्ण वृद्धि या कमी की प्रवृत्ति नहीं पाई गई। दोनों परीक्षणों में 0.1 स्तर (α) में PET की सांख्यिकीय रूप से कमी की महत्वपूर्ण प्रवृत्ति को दिखाया गया जो वाष्पीकरण के माध्यम से जल की कमी और सूखे की गंभीरता को कम करेगा। मासिक समय मान पर, $\alpha = 0.01, 0.05$ तथा 0.1 में दोनों परीक्षणों में SPI के मामले में जून महीने के लिए सांख्यिकीय रूप से कम होने का संकेत मिला (अर्थात् बढ़ते हुए सूखे की गंभीरता) जो मॉनसून की आरंभिक अवधि में वर्षा की कमी में वृद्धि का संकेत देता है और खरीफ ऋतु की फसलों की आरंभिक बुवाई में देरी को बढ़ाता है। नवंबर माह के लिए, दोनों परीक्षणों में, $\alpha = 0.1$ में SPI और RDI_{std} की वृद्धि की प्रवृत्ति (अर्थात् सूखे की गंभीरता में कमी) दिखाई दी। PET के मामले में, दोनों परीक्षणों में किसी भी माह में कोई सांख्यिकीय महत्वपूर्ण प्रवृत्ति नहीं दिखाई दी।

ABSTRACT. Drought is one of the adverse natural hazards especially in arid and semi-arid regions regarding water resources management. Under the conditions of global warming and climate change, the investigation of drought severity and its trend in arid and semi-arid regions is of primary importance. Therefore, in this study, drought severity assessment and trend detection were carried out using different meteorological drought indices like Standardized Precipitation Index (SPI) and Standardized Reconnaissance Drought Index (RDI_{std}) for a semi-arid area of Parbhani district of the Indian state of Maharashtra. The results showed that SPI and RDI_{std} behave similarly to detect drought severity except for some slight deviation in detecting moderate and normal drought severity. Out of 37 years (1983 to 2019), SPI showed 1 severe drought year, 6 moderate drought years, 22 normal years, 4 moderate wet years, 3 very wet years and 1 extremely wet year while RDI_{std} showed 1 severe drought year, 5 moderate drought years, 23 normal years, 4 moderate wet years, 3 very wet years and 1 extremely wet year. One severe drought year was observed on the basis of both SPI and RDI_{std} indices. On the basis of weekly data analysis, the frequency of drought, normal and wet week was found to be 70.58, 15.90 and 13.51 per cent, respectively. It is revealed that the short term weekly rainfall analysis provides clear picture of frequent occurrences of drought episodes at the study area as compared to longer (monthly or annual) time scale. For identifying a statistically significant trend, a non-parametric test (Mann-Kendall) and parametric test (Linear regression) were used. At an annual scale, no significant either increasing or decreasing trend was found in the case of precipitation, SPI, RDI_{std} except potential evapotranspiration (PET) and temperature. Both the tests showed a statistically significant decreasing trend of PET at 0.1 level of significance (α) which would reduce water loss through

evapotranspiration and minimize drought severity. At monthly time scale, both the tests indicated statistically decreasing trend for June month in the case of SPI (*i.e.*, increasing drought severity) at $\alpha = 0.01, 0.05$ and 0.1 which indicated increment in shortage of precipitation at early monsoon period and hence, increase in delay of early sowing of Kharif seasonal crops. For November month, both the tests showed an increasing trend of SPI and RDI_{std} (*i.e.*, decreasing drought severity) at $\alpha = 0.1$. In the case of PET, both the tests did not show any statistically significant trend at any month.

Key words – Drought, SPI, RDI_{std}, Mann-Kendall test, Linear regression test.

1. Introduction

Drought is a recurrent, multi-dimensional natural hazard that affects human life adversely. It occurs because of a shortage of precipitation, a high rate of evapotranspiration and overuse of available surface and groundwater resources (Bhelawe *et al.*, 2015; Zarei *et al.*, 2016). Since past few decades, the severity and frequency of occurrence of drought have been increased considerably due to the effects of climate change and global warming (Al-Qinna *et al.*, 2011; Edossa *et al.*, 2010; Liu *et al.*, 2012; Paulo *et al.*, 2016; Shen *et al.*, 2017; Surendran *et al.*, 2017). Drought is responsible to cause enormous impacts on ecology, environment, hydrology and agriculture. Drought is defined as the deficiency of precipitation for an extended duration that severely affects agricultural production; it causes a shortage of water for human or livestock even for consumption and influences the ecological balance (Singh *et al.*, 2014). Different specialist describes the drought in different ways; for instance, according to agriculturist's drought indicate a deficiency of moisture in the top layer of soil which relate drought to moisture shortage in topsoil up to root zone which influences the crops adversely, meteorologists relate it to when deficiency of precipitation occurs for a longer duration, according to hydrologist low runoff or low stream flow denote drought event (Zarei *et al.*, 2016). Over the last few decades, more people are affected due to increased frequency of drought as a natural hazard (Dhar *et al.*, 1979; Ray *et al.*, 1987; Kumar and Kumar, 1989; Dabral, 1996; Marathe *et al.*, 2001; Shrivastava *et al.*, 2008; Lala *et al.*, 2012; Lala *et al.*, 2014). Drought is a complex phenomenon which can be identified by its characteristics like duration, severity and areal extent. Monitoring and forecasting of drought is a research challenge (Rahmat *et al.*, 2015). There are four different types of drought as Meteorological, Agricultural, Hydrological and Socio-economic drought (Zarch *et al.*, 2011). For realizing the drought and its severity, different indices are being used. The severity of drought, its spatial and temporal extent can be estimated using these indices (Zarei *et al.*, 2016). These indices help to analyze historical drought occurrences and to estimate recurrence probability. It helps to analyze this climatic anomaly in terms of their frequency, duration, intensity and spatial extent. Adverse effects of droughts like socioeconomic, agricultural and environmental impacts can be reduced by assessment as well as forecasting of drought

behaviour (Singh and Sharma, 2003; Tiwari *et al.*, 2007; Manikandan and Tamilmani, 2011; Tigkas *et al.*, 2015). It has been proved that different drought indices are the key element for drought monitoring and the management of water resource. In order to assess the drought severity, different meteorological drought indices like the Standardized Precipitation Index (SPI) and Standardized Reconnaissance Drought Index (RDI_{std}) are widely used. The output obtained from these indices helps the academics, researchers, administrators and policy makers for decision making in water conservation and drought mitigation planning purposes. Drought indices can be calculated manually, it can also be calculated using tools or softwares designed for this purpose. The SPI makes the use of precipitation only while RDI_{std} makes the use of precipitation as well as the potential evapotranspiration (PET) simultaneously. Different meteorological parameters like temperature, wind velocity, relative humidity and sunshine duration can also affect the drought magnitude in terms of PET. By keeping all these things in mind, this study was carried out with following objectives: (i) to assess the drought severity using different meteorological drought indices (ii) to develop linear regression equation in between SPI and RDI_{std} meteorological indices and (iii) to detect the trend of drought severity using parametric and non-parametric tests for semi-arid Parbhani district of Maharashtra state of India.

2. Data sets and methodology

2.1. Study area

The Parbhani district of Maharashtra state comes under semi-arid climatic conditions. Parbhani is situated at $19^{\circ} 16' N$ latitude and $76^{\circ} 47' E$ longitude and located at an altitude of 409 m above mean sea level. It comes under moderate to moderately high precipitation zone with an average annual precipitation of 955 mm. The mean maximum and minimum temperature of the study area is $44.6^{\circ} C$ and $21.8^{\circ} C$, respectively. The mean relative humidity varies from 30 to 98 per cent. May is the hottest month while December is the coldest month of the year. April and May are dryer (low relative humidity) months of the year. The location of the study area is shown in Fig. 1. Agriculture is the main occupation of people lives in the study area.

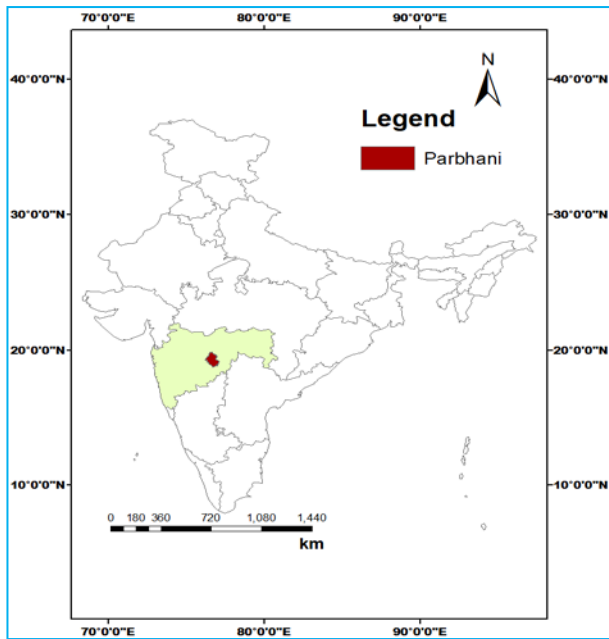


Fig. 1. Location map of the study area

2.2. Precipitation and PET data

Precipitation and other meteorological data (maximum temperature, minimum temperature, wind velocity, relative humidity and bright sunshine hours) from the year 1983 to 2019 were collected from India Meteorological Department (IMD) recognized observatory located in Vasantnao Naik Marathwada Krishi Vidyapeeth, Parbhani, Maharashtra. The monthly variation of rainfall and PET is shown in Fig. 2. Overall, rainfall is maximum in monsoon than summer and winter season. PET of the study area has been calculated using the Hargreaves method. In this study, DrinC (Drought index calculator) software was used for calculating the drought SPI and RDI_{std} indices (Tigkas *et al.*, 2014). The daily rainfall data series were divided in to annual, monthly and weekly rainfall data series for determination of drought severity. In case of weekly data analysis, the last day of every year is counted in the 52nd meteorological week while in case of leap year 29th day of February is counted in the 9th meteorological month.

2.3. Drought indices

2.3.1. SPI

McKee *et al.* (1993) invented SPI at Colorado State University to estimate the precipitation deficit at multiple time scales. The SPI is the transformation of the precipitation time series into a standardized

normal distribution. The computation of SPI consists of fitting cumulative probability distribution function (PDF) on aggregated monthly (k) precipitation series and transforming the cumulative distribution function (CDF) to the CDF of the standard normal distribution with zero mean and unit variance. The gamma PDF $[g(x)]$ is defined as:

$$g(x) = \frac{1}{\beta^\alpha \Gamma(\alpha)} x^{\alpha-1} e^{-\frac{x}{\beta}} \tag{1}$$

where, β is a scale parameter, α is a shape parameter, x is the precipitation amount and $\Gamma(\alpha)$ is the gamma function at α . Scale and shape factors can be estimated using a method of maximum likelihood. The estimated parameters can be used to find the cumulative PDF of observed precipitation events for the given month and particular time scale. The CDF is obtained by integrating Eqn. (1) as:

$$G(x) = \int_0^x g(x)dx = \int_0^x \frac{1}{\beta^\alpha \Gamma(\alpha)} x^{\alpha-1} e^{-\frac{x}{\beta}} dx \tag{2}$$

$$\alpha = \frac{1}{4A} \left(1 + \sqrt{1 + (4A)/3} \right) \tag{3}$$

$$\beta = \frac{\bar{x}}{\alpha} \tag{4}$$

$$A = \ln(\bar{x}) - \frac{\sum \ln(x)}{n} \tag{5}$$

n is the number of precipitation observations and \bar{x} refers to the sample mean of the data. Transformation of the CDF to the CDF of the standard normal distribution with zero mean and unit variance, gives SPI as:

$$SPI = \varphi^{-1}[G(x)] \tag{6}$$

This transformed probability is the SPI. A positive value of SPI indicates that precipitation is above average and a negative value denotes below average precipitation. Severe drought can be represented by a continuous negative value of SPI reaching to -1.0 or less.

2.3.2. RDI

Tsakiris & Vangelis (2005) invented RDI which represents the water balance deficit between input precipitation and an output PET. The initial value (α_k) of

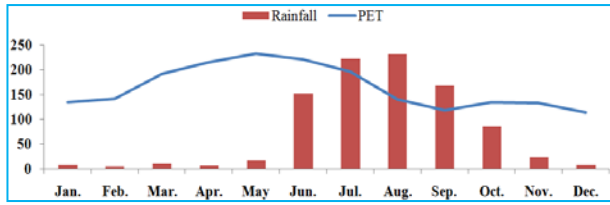


Fig. 2. Mean monthly rainfall (cm) and PET (cm) of the study area

RDI for a certain period, indicated by a certain month (k) during a year, is defined as:

$$\alpha_k = \frac{\sum_{j=1}^{j=k} P_j}{\sum_{j=1}^{j=k} PET_j} \quad (7)$$

where, P_j and PET_j are the precipitation and PET of the j^{th} month of the year. The RDI can be expressed as normalized RDI (RDI_n) and Standardized RDI (RDI_{std}). The RDI_n is computed using the following equation as:

$$RDI_n(k) = \frac{\alpha_k}{\bar{\alpha}_k} - 1 \quad (8)$$

The RDI_{std} is computed using a similar procedure which is used for SPI estimation. The equation is as follows:

$$RDI_{std}(k) = \frac{y_k - \bar{y}_k}{\sigma_k} \quad (9)$$

where, y_k is $\ln \alpha_k^i$, \bar{y}_k is the arithmetic mean, σ_k is the standard deviation.

In this study, the initial values of RDI are fitted to a gamma PDF similar to the one used for SPI. Therefore, the RDI_{std} could be computed following a similar procedure described for SPI and the values compared to the same threshold values as used for SPI drought categorization (Table 1).

2.4. Statistical methods for trend detection

Many statistical techniques (parametric or non-parametric) have been developed to detect trends within time series such as linear regression, Spearman's Rho test, Mann-Kendall test, Sen's slope estimator and Bayesian procedure, etc. Various parametric and non-parametric tests can be used for identifying trends in hydro-meteorological time series. In this study, the Mann-Kendall non-parametric and linear regression parametric tests were used to analyze the trend of drought severity.

TABLE 1

Classifications scale of SPI and RDI_{std} values

SPI/ RDI_{std} values	Category
2.00 or more	Extremely wet
1.50 - 1.99	Very wet
1.00 - 1.49	Moderately wet
0.99 to -0.99	Near normal
- 1.0 to -1.49	Moderate drought
- 1.5 to -1.99	Severe drought
-2 and less	Extreme drought

2.4.1. Mann-Kendall trend detection test

It is a non-parametric distribution free trend detection test. This type of test is the most useful because most of the hydro-meteorological time series data are not normally distributed. The Mann-Kendall test statistic S is defined as:

$$s = \sum_{i=1}^{n-1} \sum_{j=i+1}^n \text{sgn}(X_j - X_i) \quad (10)$$

where, n is a number of data points, X_i and X_j are the data values in time series i and j ($j > i$), respectively and $\text{sgn}(X_j - X_i)$ is the sign function given as:

$$\text{sgn}(X_j - X_i) = \begin{cases} +1 & \text{if } (x_j - x_i) > 0 \\ 0 & \text{if } (x_j - x_i) = 0 \\ -1 & \text{if } (x_j - x_i) < 0 \end{cases} \quad (11)$$

where, the null hypothesis (H_0) is that there is no trend in the data. If H_0 is true, then s is approximately normally distributed with zero mean (μ) and variance (σ^2) is given as:

$$\sigma^2 = \frac{n(n-1)(2n+5)}{18} \quad (12)$$

The z -statistic (critical test statistic values for various level of significance can be obtained from normal probability tables) is given as:

$$z = \begin{cases} (S-1)/\sigma & \text{if } S > 0 \\ 0 & \text{if } S = 0 \\ (S+1)/\sigma & \text{if } S < 0 \end{cases} \quad (13)$$

A positive value of s indicates an increasing trend and *vice versa*. Here, the detection of the trend was carried out at a specific level of significance (α). In this research, $\alpha = 0.01, 0.05$ and 0.1 were used.

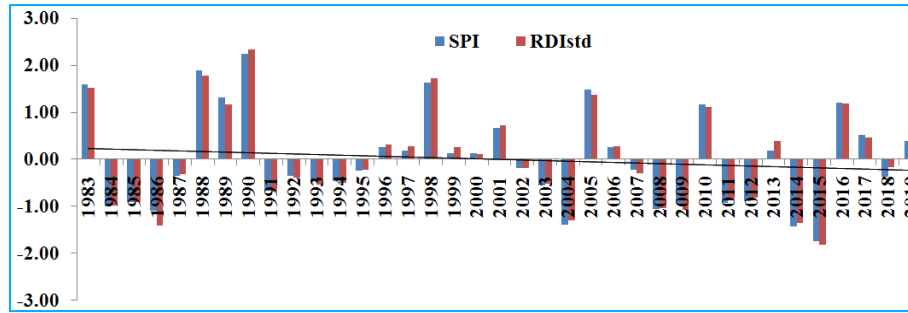


Fig. 3. Historical time series variation of annual SPI and RDI_{std}

2.4.2. Linear regression trend detection test

It is a parametric test that assumes that data are normally distributed. It tests the presence of a linear trend by identifying the relationship between time (x) and the variable of interest (y). The regression gradient (b) can be estimated as:

$$b = \frac{\sum_{i=1}^n (X_i - \bar{x})(y_i - \bar{y})}{\sum_{i=1}^n (X_i - \bar{x})(X_i - \bar{x})} \tag{14}$$

where, X_i is the time, Y_i is the variable of interest, \bar{x} and \bar{y} are the ranks. The intercept could be estimated as:

$$a = \bar{y} - b\bar{x} \tag{15}$$

The test statistic z can be estimated as:

$$z = \frac{b}{\sigma} \tag{16}$$

$$\text{where, } \sigma = \sqrt{\frac{12 \sum_{i=1}^n (y_i - a - bX_i)^2}{n(n-2)(n^2-1)}} \tag{17}$$

The test statistic z follows a Student - t distribution with $n-2$ degrees of freedom under the H_0 (critical test statistic values for different α can be obtained from Student's t statistic tables).

3. Results and discussion

In this study, the correlation between SPI and RDI_{std} was determined at different time scales. The correlation coefficients were observed to be very high at different time scale for a given semi-arid region as presented in Table 2.

TABLE 2

Correlation coefficient of the SPI and RDI_{std} at different time scales

Time scales	1 month	3 month	6 month	12 month
Correlation coefficient	0.99408	0.9963	0.9965	0.9951

The SPI and RDI_{std} indices showed similar behaviour to detect the severity of drought years, except a slight deviation for the year 1986 as shown in Fig. 3. Out of 37 years, SPI showed 1 severe drought year, 6 moderate drought years, 22 normal years, 4 moderate wet years, 3 very wet years and 1 extremely wet year while RDI_{std} showed 1 severe drought year, 5 moderate drought years, 23 normal years, 4 moderate wet years, 3 very wet years and 1 extremely wet year. RDI_{std} and SPI indices showed a nearly equal number of drought event with a slight deviation in their magnitude. The trend line shows the increased drought frequency and magnitude from the year 1983 to 2019 as shown in Fig. 3. Paired t -test indicated that SPI and RDI_{std} values do not differ significantly with a p -value of 0.9861.

The occurrence of the number of drought episodes and their severity for the past 37 years is shown in Fig. 4 for different months. While assessing the drought severity at monthly time scale, for the month of February, June, July, August, October and December, the trend line showed increased drought severity as shown in Fig. 4. The monsoon season including June, July, August, September and October experienced moderate, severe and extreme drought episodes on the basis of monthly time scale. It indicated an abrupt deviation of rainfall from mean rainfall is more in monsoon season which is responsible for the occurrence of number drought episodes in the monsoon within the study area.

3.1. Monthly linear regression in between SPI and RDI_{std} values

Monthly linear regression relationship between SPI and RDI_{std} indices are shown in Fig. 5. The low

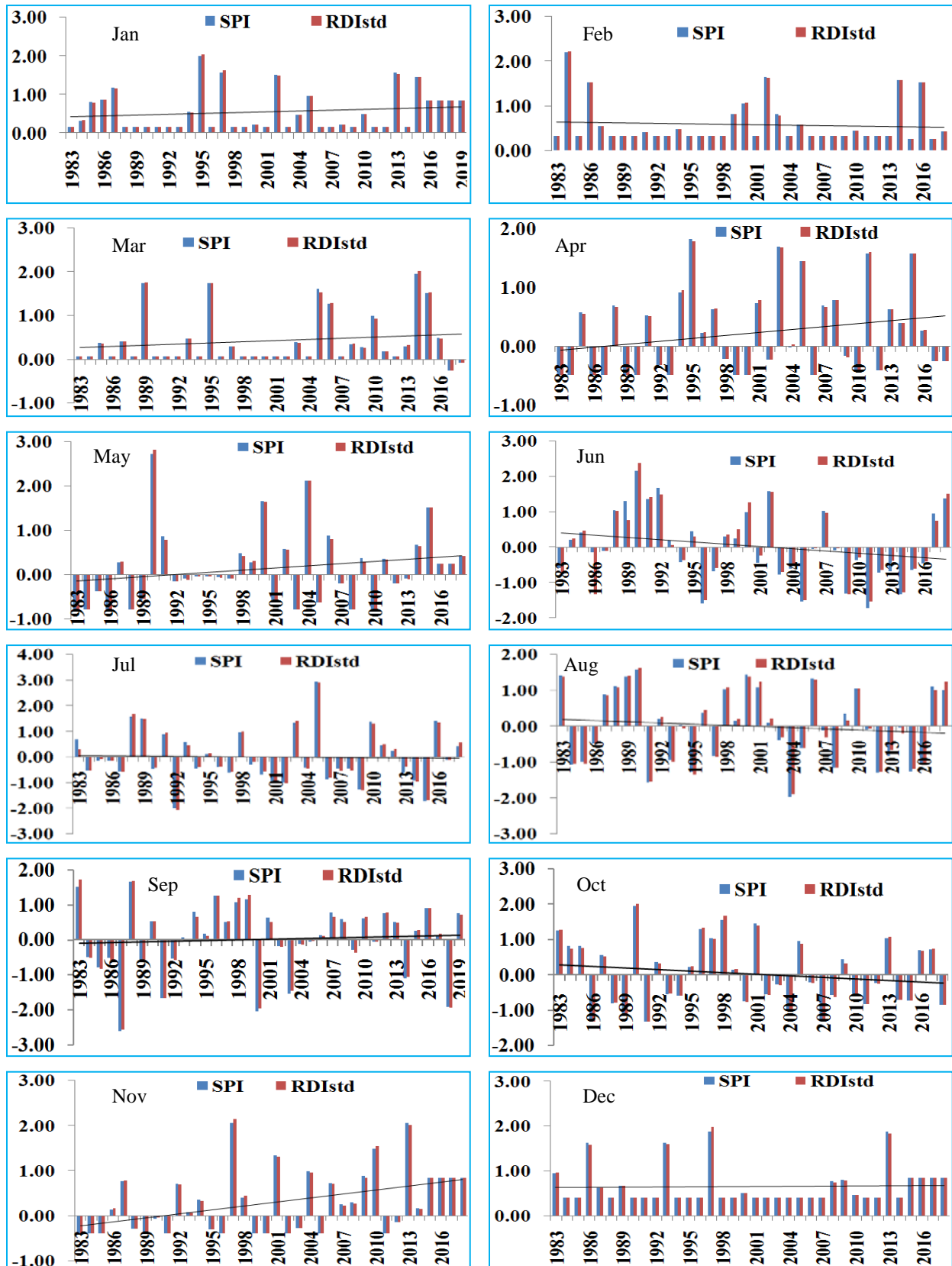


Fig. 4. SPI and RDI_{std} of different months for past historical data

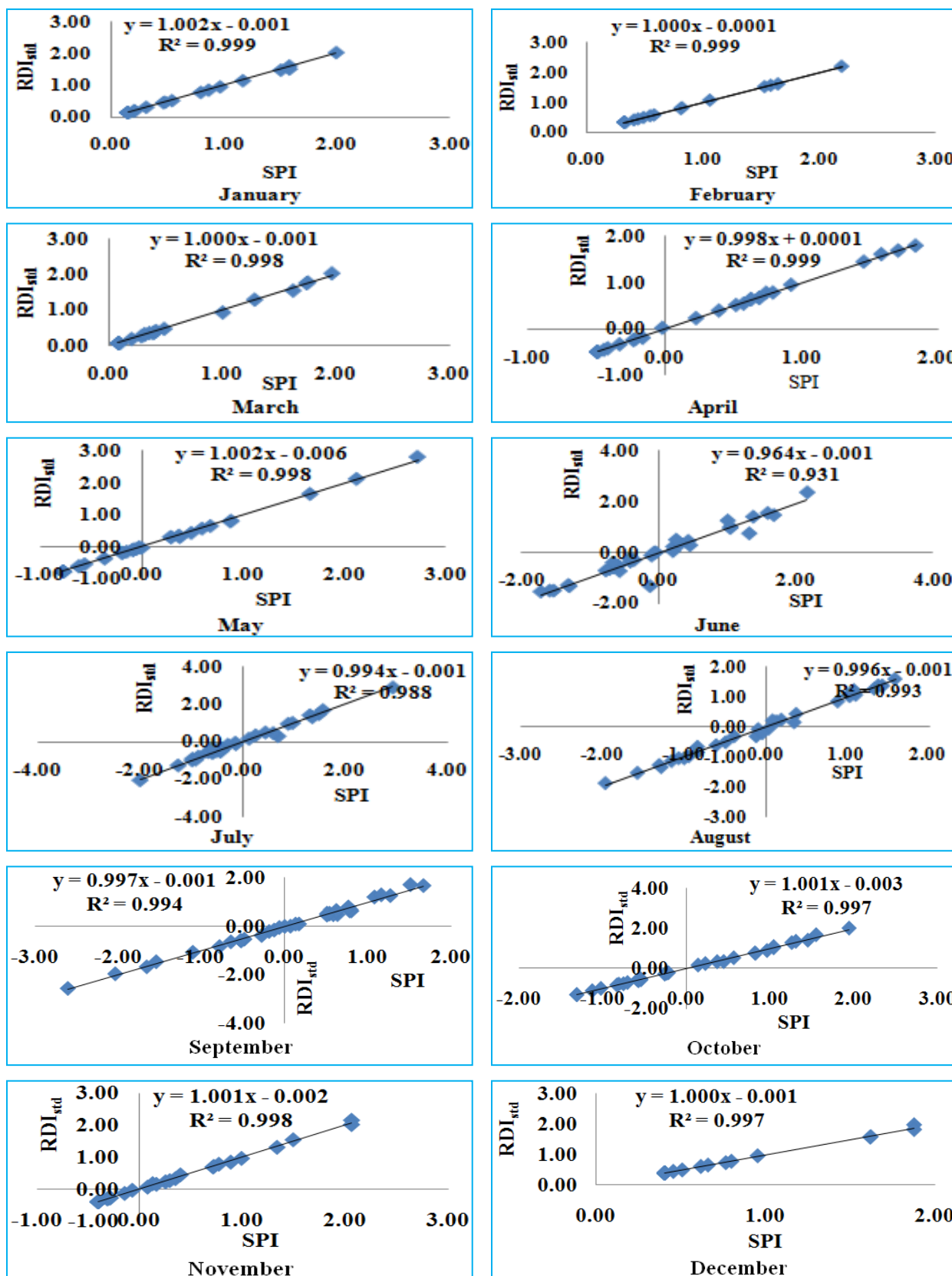


Fig. 5. Monthly linear regression between SPI and RDI_{std} for Parbhani district

TABLE 3
Mann-Kendall non parametric test for trend detection

Time scale	Variable	Test statistic (Z)	Critical values			Total S score	Result
			$\alpha = 0.1$	$\alpha = 0.05$	$\alpha = 0.01$		
(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)
Annual	Precipitation	-0.811	1.645	1.96	2.576	-51	NS
	Temperature	-1.314	1.645	1.96	2.576	-82	NS
	PET	-1.768	1.645	1.96	2.576	-110	S (0.1)
	SPI	-0.811	1.645	1.96	2.576	-51	NS
	RDI _{std}	-0.762	1.645	1.96	2.576	-48	NS
January	Precipitation	-0.227	1.645	1.96	2.576	-15	NS
	Temperature	-1.184	1.645	1.96	2.576	-74	NS
	PET	0.341	1.645	1.96	2.576	22	NS
	SPI	-0.227	1.645	1.96	2.576	-15	NS
	RDI _{std}	-0.243	1.645	1.96	2.576	-16	NS
February	Precipitation	-0.308	1.645	1.96	2.576	-20	NS
	Temperature	0.811	1.645	1.96	2.576	51	NS
	PET	-0.243	1.645	1.96	2.576	-16	NS
	SPI	-0.308	1.645	1.96	2.576	-20	NS
	RDI _{std}	-0.308	1.645	1.96	2.576	-20	NS
March	Precipitation	1.265	1.645	1.96	2.576	79	NS
	Temperature	-1.314	1.645	1.96	2.576	-82	NS
	PET	-0.243	1.645	1.96	2.576	-16	NS
	SPI	1.265	1.645	1.96	2.576	79	NS
	RDI _{std}	1.297	1.645	1.96	2.576	81	NS
April	Precipitation	1.881	1.645	1.96	2.576	117	S(0.1)
	Temperature	-0.519	1.645	1.96	2.576	-33	NS
	PET	-0.778	1.645	1.96	2.576	-49	NS
	SPI	1.881	1.645	1.96	2.576	117	S(0.1)
	RDI _{std}	1.865	1.645	1.96	2.576	116	S(0.1)
May	Precipitation	1.249	1.645	1.96	2.576	78	NS
	Temperature	-1.184	1.645	1.96	2.576	-74	NS
	PET	0	1.645	1.96	2.576	-1	NS
	SPI	1.249	1.645	1.96	2.576	78	NS
	RDI _{std}	1.314	1.645	1.96	2.576	82	NS
June	Precipitation	-2.578	1.645	1.96	2.576	-160	S(0.01)
	Temperature	-0.697	1.645	1.96	2.576	-44	NS
	PET	-1.103	1.645	1.96	2.576	-69	NS
	SPI	-2.578	1.645	1.96	2.576	-160	S(0.01)
	RDI _{std}	-1.946	1.645	1.96	2.576	-121	S(0.1)
July	Precipitation	-1.151	1.645	1.96	2.576	-72	NS
	Temperature	-0.243	1.645	1.96	2.576	-16	NS
	PET	-0.422	1.645	1.96	2.576	-27	NS
	SPI	-1.151	1.645	1.96	2.576	-72	NS
	RDI _{std}	-0.989	1.645	1.96	2.576	-62	NS
August	Precipitation	-0.778	1.645	1.96	2.576	-49	NS
	Temperature	-1.362	1.645	1.96	2.576	-85	NS
	PET	0.243	1.645	1.96	2.576	16	NS
	SPI	-0.778	1.645	1.96	2.576	-49	NS
	RDI _{std}	-0.876	1.645	1.96	2.576	-55	NS

TABLE 3 (Contd.)

(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)
September	Precipitation	0.486	1.645	1.96	2.576	31	NS
	Temperature	-2.027	1.645	1.96	2.576	-126	S(0.05)
	PET	-1.265	1.645	1.96	2.576	-79	NS
	SPI	0.486	1.645	1.96	2.576	31	NS
	RDI _{std}	0.519	1.645	1.96	2.576	33	NS
October	Precipitation	-0.876	1.645	1.96	2.576	-55	NS
	Temperature	0.243	1.645	1.96	2.576	16	NS
	PET	1.119	1.645	1.96	2.576	70	NS
	SPI	-0.876	1.645	1.96	2.576	-55	NS
	RDI _{std}	-0.843	1.645	1.96	2.576	-53	NS
November	Precipitation	1.897	1.645	1.96	2.576	118	S(0.1)
	Temperature	1.476	1.645	1.96	2.576	92	NS
	PET	0.486	1.645	1.96	2.576	31	NS
	SPI	1.897	1.645	1.96	2.576	118	S(0.1)
	RDI _{std}	1.849	1.645	1.96	2.576	111	S(0.1)
December	Precipitation	-0.276	1.645	1.96	2.576	-18	NS
	Temperature	0.081	1.645	1.96	2.576	6	NS
	PET	1.557	1.645	1.96	2.576	97	NS
	SPI	-0.276	1.645	1.96	2.576	-18	NS
	RDI _{std}	-0.276	1.645	1.96	2.576	-18	NS

determination coefficient (R^2) in June was found to be about 0.931 because of abrupt variation of different meteorological parameters in this month. These linear regression equations can be useful to estimate the RDI_{std} drought index with the availability of only rainfall data (SPI) because RDI_{std} drought index actually represents agricultural drought as it depends on PET.

3.2. Distribution of weekly rainfall

Weekly data are more useful for proper planning and management of agricultural activities as compared to monthly, seasonal and annual rainfall data (Bhelawe *et al.*, 2015). The daily rainfall data series were converted in to weekly rainfall data for determination of drought severity. The weekly rainfall was classified as drought, normal and wet weeks. A week which receives rainfall less than or equal to half of the average weekly rainfall is called a dry week while a week which receives rainfall twice the average weekly rainfall is called a wet week and a week which receives rainfall between the limits of weekly rainfall corresponding to dry and wet week is called a normal week (Ramdas and Mallik, 1948; Ray *et al.*, 2016). The average values of weekly rainfall and corresponding values of drought, normal and wet week's rainfall ranges are presented in Table 5. The weekly rainfall distribution shows a better understanding about severity of drought and practices to be followed for short

term planning and management of cropping pattern. The total numbers of dry, normal and wet weeks corresponding to different meteorological weeks considering 37 years rainfall data are presented in Table 5. The analysis of average weekly rainfall indicated that the maximum average weekly rainfall was observed during 32nd week having magnitude of 60.08 mm followed by 34th, 35th, 31st meteorological week while the lowest rainfall was received in 51st meteorological week with average weekly value of 0.04 mm. The weekly analysis indicated that 51st standard meteorological week showed maximum number of drought weeks with frequency of 36 out of 37 followed by 4th, 5th, 7th, 11th, 48th and 49th standard meteorological weeks. The minimum numbers of drought weeks were observed during 24th standard meteorological week which receives 13 drought weeks, 18 normal weeks and 6 wet weeks out of 37 meteorological weeks. The analysis reveals that the frequency of drought, normal and wet week was found to be 70.58, 15.90 and 13.51 per cent, respectively. The analysis indicated that the maximum frequency of drought weeks was observed in winter and summer season as compared to monsoon season. Hence, it indicated the need of supplemental irrigation during this water deficit periods. Hence, short term weekly rainfall analysis provides clear picture of frequent occurrences of drought episodes at the study area as compared to longer (monthly or annual) time scale.

TABLE 4
Linear regression parametric test for trend detection

Time scale	Variable	Test statistic (Z)	Critical values			σ	Result
			$\alpha = 0.1$	$\alpha = 0.05$	$\alpha = 0.01$		
(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)
Annual	Precipitation	-1.367	1.697	2.042	2.75	5.897	NS
	Temperature	-1.84	1.697	2.042	2.75	0.013	S(0.1)
	PET	-1.706	1.697	2.042	2.75	1.359	S(0.1)
	SPI	-1.322	1.697	2.042	2.75	0.019	NS
	RDI _{std}	-1.118	1.697	2.042	2.75	0.019	NS
January	Precipitation	-0.187	1.697	2.042	2.75	0.01	NS
	Temperature	-1.645	1.697	2.042	2.75	0.036	NS
	PET	-0.026	1.697	2.042	2.75	0.097	NS
	SPI	-0.187	1.697	2.042	2.75	0.01	NS
	RDI _{std}	-0.212	1.697	2.042	2.75	0.01	NS
February	Precipitation	-0.599	1.697	2.042	2.75	0.009	NS
	Temperature	0.463	1.697	2.042	2.75	0.019	NS
	PET	0.027	1.697	2.042	2.75	0.123	NS
	SPI	-0.599	1.697	2.042	2.75	0.009	NS
	RDI _{std}	-0.624	1.697	2.042	2.75	0.009	NS
March	Precipitation	1.248	1.697	2.042	2.75	0.011	NS
	Temperature	-1.2	1.697	2.042	2.75	0.023	NS
	PET	-0.883	1.697	2.042	2.75	0.229	NS
	SPI	1.248	1.697	2.042	2.75	0.011	NS
	RDI _{std}	1.243	1.697	2.042	2.75	0.011	NS
April	Precipitation	1.551	1.697	2.042	2.75	0.014	NS
	Temperature	-0.397	1.697	2.042	2.75	0.018	NS
	PET	-0.169	1.697	2.042	2.75	0.169	NS
	SPI	1.551	1.697	2.042	2.75	0.014	NS
	RDI _{std}	1.577	1.697	2.042	2.75	0.014	NS
May	Precipitation	0.693	1.697	2.042	2.75	0.017	NS
	Temperature	-1.286	1.697	2.042	2.75	0.022	NS
	PET	0.296	1.697	2.042	2.75	0.236	NS
	SPI	0.693	1.697	2.042	2.75	0.017	NS
	RDI _{std}	0.637	1.697	2.042	2.75	0.017	NS
June	Precipitation	-2.9	1.697	2.042	2.75	0.017	S(0.01)
	Temperature	-1.453	1.697	2.042	2.75	0.12	NS
	PET	-1.676	1.697	2.042	2.75	1.076	NS
	SPI	-2.9	1.697	2.042	2.75	0.017	S(0.01)
	RDI _{std}	-2.105	1.697	2.042	2.75	0.018	S(0.05)
July	Precipitation	-0.445	1.697	2.042	2.75	0.02	NS
	Temperature	-0.088	1.697	2.042	2.75	0.015	NS
	PET	-1.048	1.697	2.042	2.75	0.227	NS
	SPI	-0.445	1.697	2.042	2.75	0.02	NS
	RDI _{std}	-0.316	1.697	2.042	2.75	0.02	NS
August	Precipitation	-0.837	1.697	2.042	2.75	0.02	NS
	Temperature	-1.482	1.697	2.042	2.75	0.012	NS
	PET	0.276	1.697	2.042	2.75	0.198	NS
	SPI	-0.837	1.697	2.042	2.75	0.02	NS
	RDI _{std}	-0.843	1.697	2.042	2.75	0.02	NS

TABLE 4 (Contd.)

(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)
September	Precipitation	0.551	1.697	2.042	2.75	0.02	NS
	Temperature	-2.426	1.697	2.042	2.75	0.015	S(0.05)
	PET	-0.89	1.697	2.042	2.75	0.193	NS
	SPI	0.551	1.697	2.042	2.75	0.02	NS
	RDI _{std}	0.511	1.697	2.042	2.75	0.02	NS
October	Precipitation	-1.052	1.697	2.042	2.75	0.018	NS
	Temperature	0.75	1.697	2.042	2.75	0.02	NS
	PET	1.146	1.697	2.042	2.75	0.206	NS
	SPI	-1.052	1.697	2.042	2.75	0.018	NS
	RDI _{std}	-1.052	1.697	2.042	2.75	0.018	NS
November	Precipitation	2.029	1.697	2.042	2.75	0.013	S(0.1)
	Temperature	1.319	1.697	2.042	2.75	0.023	NS
	PET	0.695	1.697	2.042	2.75	0.166	NS
	SPI	2.029	1.697	2.042	2.75	0.013	S(0.1)
	RDI _{std}	1.958	1.697	2.042	2.75	0.013	S(0.1)
December	Precipitation	-0.328	1.697	2.042	2.75	0.009	NS
	Temperature	0.218	1.697	2.042	2.75	0.022	NS
	PET	1.535	1.697	2.042	2.75	0.111	NS
	SPI	-0.328	1.697	2.042	2.75	0.009	NS
	RDI _{std}	-0.357	1.697	2.042	2.75	0.009	NS

3.3. Trend detection

3.3.1. Mann-Kendall non-parametric test

The non-significant (NS) or significant (S) trend of different variables at annual and monthly time scale was detected using a Mann-Kendall non-parametric test which is presented in Table 3. At annual scale, no significant trend was found for precipitation, temperature, SPI, RDI_{std} except PET. PET shows a negative trend for $\alpha = 0.1$ at an annual scale. At a monthly time scale, a significant positive trend was observed in April for precipitation, SPI and RDI_{std} at $\alpha = 0.1$. A significant decrease in precipitation and SPI was observed in June at all $\alpha = 0.1, 0.05$ and 0.01 , respectively. A significant decrease in RDI_{std} was observed in June at $\alpha = 0.1$. A significant decrease in temperature was observed in September at $\alpha = 0.1$ and 0.05 . The positive increase in precipitation, SPI and RDI_{std} were observed in November at $\alpha = 0.1$. The significant negative trend in precipitation, SPI and RDI_{std} would cause reduction of water availability in the respective month. A positive trend of temperature and negative trend of precipitation, SPI and RDI_{std} would cause water shortage in the respective month.

3.3.2. Linear regression parametric test

The non-significant (NS) and significant (S) trend of different variables at annual and monthly time scale was

detected using linear regression parametric test which is presented in Table 4. At annual time scale, temperature and PET indicated a decreasing trend at $\alpha = 0.1$. In June, precipitation and SPI indicated a decreasing trend at all $\alpha = 0.1, 0.05$ and 0.01 while RDI_{std} indicated a decreasing trend at $\alpha = 0.1$ and 0.05 . In September, temperature indicated a decreasing trend at $\alpha = 0.1$ and 0.05 . Precipitation, SPI and RDI_{std} indicated increasing in November month at $\alpha = 0.1$.

3.3.3. Comparison of Mann-Kendall non-parametric and linear regression parametric test

At an annual time scale, Mann-Kendall non-parametric test and linear regression parametric test showed a statistically significant decreasing trend of PET at $\alpha = 0.1$ which would reduce water loss through evapotranspiration and minimizes drought severity. At an annual time scale, only a linear regression test showed a statistically significant decreasing trend for the temperature at $\alpha = 0.1$ which would reduce drought severity. At a monthly time scale, for April month, Mann-Kendall non-parametric test showed an increasing trend for SPI and RDI_{std} (*i.e.*, decreasing drought severity) at $\alpha = 0.1$ but linear regression parametric test did not show any significant trend for this month. Both the tests showed a decreasing trend of SPI in June (*i.e.*, increasing drought severity) at $\alpha = 0.01, 0.05$ and 0.1 which indicated a shortage of precipitation in this month. Both the tests

TABLE 5
Analysis of weakly rainfall for drought detection at Parbhani

Meteorological week (1)	Average rainfall (mm) (2)	Rainfall value (mm)			Total number of weeks		
		Drought (less than) (3)	Normal (in between) (4)	Wet (more than) (5)	Drought (6)	Normal (7)	Wet (8)
1	1.48	0.74	0.74-2.97	2.97	32	1	4
2	4.17	2.09	2.09-8.34	8.34	33	0	4
3	1.16	0.58	0.58-2.33	2.33	33	1	3
4	1.06	0.53	0.53-2.13	2.13	35	0	2
5	0.55	0.27	0.27-1.09	1.09	35	0	2
6	2.45	1.23	1.23-4.91	4.91	31	1	5
7	1.16	0.58	0.58-2.31	2.31	34	1	2
8	0.95	0.48	0.48-1.90	1.90	33	0	4
9	1.34	0.67	0.67-2.68	2.68	31	2	4
10	6.09	3.04	3.04-12.18	12.18	30	2	5
11	2.53	1.27	1.27-5.06	5.06	34	0	3
12	0.73	0.36	0.36-1.46	1.46	31	1	5
13	1.44	0.72	0.72-2.88	2.88	32	0	5
14	1.19	0.60	0.60-2.68	2.68	29	4	4
15	2.5	1.25	1.25-5.01	5.01	29	3	5
16	1.05	0.53	0.53-2.11	2.11	29	1	7
17	0.98	0.49	0.49-1.96	1.96	27	2	8
18	2.11	1.06	1.06-4.22	4.22	28	2	7
19	1.11	0.55	0.55-2.22	2.22	27	5	5
20	4.9	2.45	2.45-9.81	9.81	29	3	5
21	6.66	3.33	3.33-13.32	13.32	27	5	5
22	8.53	4.27	4.27-17.07	17.07	26	7	4
23	28.65	14.33	14.33-57.30	57.30	18	13	6
24	41.98	20.99	20.99-83.95	83.95	13	18	6
25	43.03	21.52	21.52-86.06	86.06	14	19	4
26	40.31	20.15	20.15 - 80.60	80.60	19	14	4
27	37.83	18.91	18.91-75.66	75.66	18	15	4
28	21.07	21.07	21.07-84.29	84.29	16	16	5
29	43.45	21.73	21.73-86.91	86.91	18	12	7
30	37.64	37.64	37.64-150.54	150.54	19	15	3
31	50.0	25.01	25.01-100.04	100.04	19	12	6
32	60.08	30.04	30.04-120.16	120.16	25	4	8
33	43.58	21.79	21.79-87.17	87.17	17	12	8
34	59.76	29.88	29.88-119.51	119.51	16	17	4
35	56.07	28.03	28.03-112.13	112.13	15	16	6
36	46.84	23.42	23.42-93.68	93.68	17	12	8
37	31.04	15.52	15.52-62.08	62.08	18	10	9
38	41.68	20.84	20.84-83.37	83.37	15	15	7
39	30.58	15.29	15.29-61.15	61.15	21	10	6
40	33.60	16.80	16.80-67.19	67.19	22	8	7
41	22.17	11.08	11.08-44.33	44.33	24	5	8
42	17.17	8.58	8.58-34.33	34.33	26	6	5
43	11.29	5.65	5.65-22.58	22.58	29	2	6
44	3.91	1.95	1.95-7.82	7.82	29	3	5

TABLE 5 (Contd.)

(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)
45	5.28	2.64	2.64-10.56	10.56	29	3	5
46	4.27	2.13	2.13-8.54	8.54	29	2	6
47	7.68	3.84	3.84-15.37	15.37	29	2	6
48	7.03	3.51	3.51-14.05	14.05	34	1	2
49	4.07	2.04	2.04-8.14	8.14	34	0	3
50	2.05	1.02	1.02-4.09	4.09	33	1	3
51	0.04	0.02	0.02-0.08	0.08	36	0	1
52	2.48	1.24	1.24-4.96	4.96	31	2	4
Total					1358	306	260

showed an increasing trend for SPI and RDI_{std} in November (*i.e.*, decreasing drought severity) at $\alpha = 0.1$. In the case of PET, both the tests did not show any significant trend in any month. Both the tests showed a significant decreasing trend of temperature in September at $\alpha = 0.1$ and 0.05 which would help to reduce the loss of water through evapotranspiration and minimize adverse effects on crop growth due to water shortage.

4. Conclusions

In this study, meteorological drought indices like SPI and RDI_{std} were used to assess the drought severity of the semi-arid Parbhani district of the Indian state of Maharashtra. Here, it was observed that SPI and RDI_{std} behave similarly for detecting the drought severity class except for some slight deviation in detecting moderate and normal drought years. One severe drought year was observed in based on both SPI and RDI_{std} indices. The developed linear regression equations can also be useful to estimate the RDI_{std} drought index with the availability of only rainfall data (SPI) because RDI_{std} drought index actually represents agricultural drought as it depends on PET. On the basis of weekly rainfall analysis, winter and summer seasons were found to be more prone to drought hazards as compared to monsoon season. Hence, it indicated the need of supplemental irrigation during this water deficit periods. Here, it was observed that short term weekly rainfall analysis provides clear picture of frequent occurrences of drought episodes as compared to longer (monthly or annual) time scale. Hence, weekly rainfall data analysis is more useful for short term strategic planning in agriculture. Both the Mann-Kendall non-parametric and linear regression parametric tests showed a statistically significant decreasing trend of SPI in June (*i.e.*, increasing drought severity) at all $\alpha = 0.01, 0.05$ and 0.1 which indicated an increase in shortage of precipitation at the initiation of the monsoon season. The reduction of precipitation at the initiation of monsoon would be responsible for delaying the sowing of Kharif

seasonal crops which may affect the crop yield. Both the tests showed an increasing trend of SPI and RDI_{std} in November (*i.e.*, decreasing drought severity) at $\alpha = 0.1$ which would be beneficial for growing of winter seasonal crops. As the drought episodes are more frequent in arid and semi-arid regions, hence future study can be focused with combined drought index using remote sensing data like Normalized Difference Vegetation Index (NDVI) etc.

Acknowledgements

The authors would like to thank G. B. Pant University of Agriculture and Technology, Pantnagar for providing necessary computer facilities. The contents and views expressed in this research paper/article are the views of the authors and do not necessarily reflect the views of the organizations they belong to.

References

- Al-Qinna, M. I., Hammouri, N. A., Obeidat, M. M. and Ahmad, F. Y., 2011, "Drought analysis in Jordan under current and future climates", *Climatic Change*, **106**, 421-440.
- Bhelawe, S., Chaudhary, J. L., Manikandan, N. and Deshmukh, R., 2015, "Meteorological drought assessment in Raipur district of Chhattisgarh state, India", *Plant Archives*, **15**, 1, 465-469.
- Dabral, P. P., 1996, "Meteorology drought analysis based on rainfall data", *Indian Journal of Soil Conservation*, **24**, 1, 37-40.
- Dhar, O. N., Rakhecha, P. R. and Kolkarni, A. K., 1979, "Rainfall study of severe drought year of India", *International Symposium in Hydrological Aspect of drought*, **1**, 28-36.
- Edossa, D. C., Babel, M. S. and Gupta, A. D., 2010, "Drought Analysis in the Awash River Basin, Ethiopia", *Water Resources Management*, **24**, 1441-1460.
- Kumar, D. and Kumar, S., 1989, "Drought analysis based on rainfall data", *Indian Journal of Soil Conservation*, **17**, 1, 55-59.
- Lala, I. P. R., Bora, P. K., Singh, A. K., Singh, R., Singh, N. J. and Feroze, S. M., 2014, "Impact and Assessment of Meteorological Drought on Rice Based Farming System in East Garo Hills District of Meghalaya, India", *Journal of Agri. Search*, **1**, 4, 227-232.

- Lala, I. P. R., Singh, V. R., Singh, A. K., Singh, R. and Feroze, S. M., 2012, "Meteorological Drought Assessment in Barapani, Meghalaya", *Journal of Indian Water Resources Society*, **32**, 1-2.
- Liu, L., Hong, Y., Bednarczyk, C. N., Yong, B., Shafer, M. A., Riley, R. and Hocker, J. E., 2012, "Hydro-Climatological Drought Analyses and Projections Using Meteorological and Hydrological Drought Indices: A Case Study in Blue River Basin, Oklahoma", *Water Resources Management*, **26**, 2761-2779.
- Manikandan, M. and Tamilmani, D., 2011, "Assessment of occurrence and frequency of drought using rainfall data in Coimbatore, India", *Asian journal of environmental science*, **6**, 2, 136-146.
- Marathe, R. A., Mohanty, S. and Singh, S., 2001, "Meteorological drought analysis based on rainfall data of Nagpur", *Journal of Soil and Water Conservation*, **45**, 1-5.
- McKee, T. B., Doesken, N. J. and Kleist, J., 1993, "The relationship of drought frequency and duration to time scale", Proceedings of the Eighth Conference on Applied Climatology, *American Meteorological Society*, Anaheim (CA), 179-184.
- Paulo, A., Martins, D. and Pereira, L. S., 2016, "Influence of Precipitation Changes on the SPI and Related Drought Severity: An Analysis Using Long-Term Data Series", *Water Resources Management*, **30**, 5737-5757.
- Rahmat, S. N., Jayasuriya, N. and Bhuiyan, M., 2015, "Assessing drought using meteorological drought indices in Victoria, Australia", *Hydrology Research*, **46**, 3, 463-476.
- Ramdas, L. A. and Malik, A. K., 1948, "Agricultural situation in India", Technical Bulletin, ICAR, New Delhi.
- Ray, C. R., Senapati, P. C. and Lal, R., 1987, "Investigation of drought from rainfall data at Gopalpur, Orissa", *Indian Journal of Soil Conservation*, **15**, 1, 15-19.
- Ray, M., Patro, H. and Mishra, N., 2016, "Analysis of rainfall for drought occurrences in Mayurbhanj, Odisha", *Annals of Plant and Soil Research*, **18**, 3, 280-283.
- Shen, Q., Liang, L., Luo, X., Li, Y. and Zhang, L., 2017, "Analysis of the spatial-temporal variation characteristics of vegetative drought and its relationship with meteorological factors in China from 1982 to 2010", *Environmental Monitoring Assessment*, **189**, 471.
- Shrivastava, K. S., Raveendra K. R. and Pandey, A., 2008, "Assessment of meteorological droughts in North Lakhimpur district of Assam", *Journal of Indian Water Resources Society*, **28**, 2, 26-31.
- Singh, B. and Sharma, M. K., 2003, "A study of pattern of rainfall of crop research station-Masodha (U. P.)", *MAUSAM*, **54**, 2, 552-556.
- Singh, B., Arya, C. K., Singh, J. and Mourya, K. K., 2014, "Analysis of rainfall data for storage and irrigation planning in humid south-eastern plain of Rajasthan in India", *Journal of Applied and Natural Science*, **6**, 1, 214-219.
- Surendran, U., Kumar, V., Ramasubramoniam, S. and Raja, P., 2017, "Development of Drought Indices for Semi-Arid Region Using Drought Indices Calculator (DrinC) - A Case Study from Madurai District, a Semi-Arid Region in India", *Water Resource Management*, **31**, 3593-3605.
- Tigkas, D., Vangelis, H. and Tsakiris, G., 2015, "DrinC: a software for drought analysis based on drought indices", *Earth Science Informatics*, **8**, 3, 697-709.
- Tiwari, K. N., Paul, D. K. and Gontia N. K., 2007, "Characterization of meteorological drought", *Hydrology*, **30**, 1-2, 15-27.
- Tsakiris, G. and Vangelis, H., 2005, "Establishing a drought index incorporating evapotranspiration", *European Water*, **9/10**, 3-11.
- Zarch, M. A. A., Malekinezhad, H., Mobin, M. H., Dastorani, M. T. and Kousari, M. R., 2011, "Drought Monitoring by Reconnaissance Drought Index (RDI) in Iran", *Water Resource Management*, **25**, 3485-3504.
- Zarei, A. R., Moghimi, M. M. and Mahmoudi, M. R., 2016, "Parametric and Non-Parametric Trend of Drought in Arid and Semi-Arid Regions Using RDI Index", *Water Resource Management*, **30**, 5479-5500.
-