Relationship between Vertical Currents and Intensity of Precipitation

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(Received 27th Angust, 1946.)

A¹E of condensation in a titing column of saturated air can be easily found out when it is assumed that the column is thermally and physically isolated. For purpose of evaluating the rate of precipitation further assumption is made, viz., all the products of condensation are immediately detached from the system and fall out as precipitation. On these assumptions formulas for the amounts of precipitation have been given by several writers. Fulks¹ recently derived a workable formula from which the rate of precipitation can be evaluated when the physical state of the upper layers and the velocity of the ascending layer are known. The subject has been briefly discussed here following Holmboe².

Take a sample saturated air of $(1+w_s)$ tons having one ton of dry air and w. ton of water vapour. While it ascends under pseudo-adiabatic process, precipitation from this mass will be at the rate of $-\frac{d w_s}{dt}$.

We write $-\frac{dw_s}{dt} = -\frac{dw_t}{dz} \cdot \frac{dz}{dt} = \frac{-dw_s}{dx} \cdot v_z$

For simplifying calculation we take $v_z = 1$ m/sec to begin with. We may then say that when the mass of $(1+w_s)$ tons of the saturated air is lifted one metre the amount of precipitation will be- $\frac{dw_s}{dz}$ ton. Per ton of saturated air, the precipitation is therefore $-\frac{1}{1+w}$. $\frac{dw_s}{dz}$ ton. If we consider in the above sample of air a column with cross-sectional area of one square metre and height of one dynamicdecimetre, the mass of the column will be $\frac{1}{\alpha}$ ton. The precipitation from this column will therefore be

$$
P_1 = -\frac{1}{\alpha g} \frac{d w_s}{(1 + w_s)} \cdot \frac{d w_s}{dz} \cdot or = -\frac{1}{\alpha (1 + w_s)} \cdot \frac{d w_s}{d \phi}
$$

Here, α is the sp. vol. of the saturated air,

ws the mising ratio,

 ϕ the dynamic height,

and g the acceleration of gravity.

During the pseudo-adiabatic pr cess the heat released by condensation is used up in heating the entire mass of the air. Therefore $Ldw_s = (1 + w_s) (C_p dT - \alpha dp)$ or

dividing by
$$
d\phi
$$
, $- L \frac{dw_s}{d\phi} = (1 + w_s) (C_p \frac{dT}{d\phi} - \alpha \frac{dp}{d\phi}) \cdots (2)$

Eliminating $\frac{dw_8}{d\phi}$ between (1) and (2) and substituting

$$
d\phi = -\alpha
$$
 dp, we get, $P_1 = \frac{1 + C_p \frac{dT}{d\phi}}{\alpha L}$

Here, C_p is the sp. heat of moist air at constant pressure, and L the heat of condensation or of sublimation according as the process takes place in the rain stage of snow stage.

If we assume that the condensed water will fall on ground over an area of one square metre the rain collected per second will be P_1 . 10³ millimetres, or at the same rate the rainfall is 3.6×10^{6} P, mm/hr.

For ease of calculation, a column of 100 dynamic metres in height, ascending with a velocity of v_{y} m/sec. may be taken. In that case the rate of precipitation becomes

$$
P_{100} = 3.6 \times 10^{9} \frac{1 + C_{\text{P}} \frac{dT}{d\phi}}{\alpha L} \cdot v_{\text{z}} \qquad \cdots \qquad (4)
$$

Or, putting $3.6 \times 10^{9} \frac{1 + C_{\text{P}} \frac{dT}{d\phi}}{\alpha L} = r$
We may write $P_{100} = r, v_{\text{z}}$

Precipitation from the entire column would be $P = \Sigma r_v v_z$... (5) when the total height of the saturated layer is broken up into layers of 100 dy.m. thickness. The value of r may be obtained from RAOB data and that of P readily from rainfall intensity record; the value of vz can thus be determined.

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t \mathbf{p} \bar{u} ŧ $\mathbf u$ \mathbf{p} 942 25 $R₅$ 600 4 100 900 21 90 500 100 -4 95 850 19 400 -14 ××. 750 13 100 500 -28 \sim 700 11 100 200 -46 ċ,

Take the rainfall at 2130 hrs. I.S.T. on 22nd June, 1945. From radiosonde ascent at 20.30 hrs. we find the conditions of the upper layers as follows :-

Ascent curve indicated that the surface air had a lapse rate just greater than the dry-adiabatic. It was, therefore, indifferent and would rise with small perturbations. Rising from the ground a parcel of air would get saturated at about 900 mb. level and therea'ter continue to rise due to buoyancy until it came to equilibrium with its environment at about 200 mb. level.

Assuming $v_z=1$ m/sec, the rate of precipitation has been calculated from the entire saturated layer as follows :-

Sample calculation :- Take the layer 800-700 mb. Mean value in the layer $p=75.0$ cb., T=286 °.5 K $w_s = 13.3 \times 10^{-3}$

Thickness of the layer is given by

H₇₀₀ -- H₈₀₀ = 28.7 (273+t_m °C) in $\frac{p_{700}}{p_{800}}$

$$
\frac{d\mathbf{T}}{d\phi} = \frac{-5}{11880} = -0.42 \times 10^{-3} \text{ deg./dy. dm.}
$$

\nC_p = C_{pd} (1+0.9w) = 1004 (1+0.9 × .013) = 1016 kJton.
\nC_p
$$
\frac{d\mathbf{T}}{d\phi} = -1016 \times 0.42 \times 10^{-3} = -0.426
$$

\n
$$
\alpha = \frac{R_d \mathbf{T}}{p} \cdot \frac{287 + 286.5 (1 \times .61 \times .013)}{75.0} = \frac{1107 \text{ m}^3/\text{ton}}{1107 \text{ m}^3/\text{ton}}
$$

\nr = 3.6 × 10⁹ ×
$$
\frac{1 + C_p \frac{d\mathbf{T}}{d\phi}}{\alpha L} = 3.6 \times 10^9 \times \frac{1 - 0.426}{1107 \times 2.5 \times 10^8}
$$

\n= 0.746 mm/hr.

= 28.7 (273+13.5) in
$$
\frac{8}{7}
$$
 = 1188 dyn. metres

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The rain commenced at 2130 hrs., i.e., about one hour after the ascent was taken. The structure of the rainfall as shown by the intensity record is given below :-

The average rate of rainfall, as found from the record, is $\frac{1 \cdot 2 \times 60}{52} = 1 \cdot 39$ in./hr. or = 35.2 mm./hr. Rate of precipitation, as calculated, is 26.7 mm./hr. The average rate of ascent of the saturated column is therefore v_a av. $\frac{35.2}{26.7}$ = 1.3 m./sec.

During the fall the maximum intensity attained was 6 in./hr. or 152 mm./hr. for a very short time. It will not be incorrect to assume that only v_z changed during the period and that other factors contributing to the rate of precipitation remained more or

less constant. Thus,
$$
v_{\text{z}} = \frac{152}{26.7} = 5.7 \text{ m./sec.}
$$

A few rainfall records have been studied together with the upper-air conditions prevailing shortly before the fall, and proceeding in the way mentioned above the speed values of the vertical currents of saturated air have been derived. Unfortunately the number of occasions selected for study so far has not been large. However, the results of study are given in Tables 1 and 2 below :

TABLE 1.

Date	Rainfall			Duration of Intensity of Rainfall in minutes								
	Time I S.T.	Dura- tion in	Amt. in h. m. inches	0.5''/ hr.	$1^{\nu}/$ hr.	20/ hr.	5''/ hr.	4''/ hr.	$5^{\prime\prime}/$ hr.	ϵ "/ hr.	Aver- age	Max.
$22 - 6 - 1915$	21.30	$0 - 52$	1.2	34	24	15 $(5 + 5)$ $+5)$	8 $(3+5)$	3	$\overline{\bf 2}$	2	1.4	$6 - 0$
$6 - 7 - 19 + 5$	19.30	$0 - 55$	$0 - 9$	26	22	11 $(3+2)$ $+6)$	3	$\bf{2}$	$\ddot{}$		0.88	4.0
17-7-1945	22.00	$0 - 45$	$1-3$	38	19	9 $(2+3)$	$\bf{2}$	\cdots	\bullet		$1 - 73$	3·0
$21 - 7 - 1915$			4.87		92	$\frac{+2}{+2}$ 31	12	6				
	22.11	$3 - 48$		134		$(5 + 5)$ $+2$ $+3+2$	$(2+3)$ $2+7$	$(2+4)$	\cdots	\cdots	$1 - 28$	4.0
						$+2+3$ $+4$ $\frac{1}{8}$						
$14 - 8 - 1945$	$19 - 45$	$1 - 17$	1.3	57	23	$(3+3)$ $+2)$	$\boldsymbol{3}$	$\overline{2}$	\cdots	\bullet	$1-0$	4.0

Structure of Rainfall.

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Speed of vertical currents in convective tayer and their duration, as derived from analysis of rainfall data in table 1. V_z vertical velocity in metre per sec. and d the duration in minutes.

Table 1 shows among other things how long it rained at different intensities. For intensities greater than 2 in./hr. the number of showers together with their duration in minutes are shown by figures in bracket. The last two columns give respectively the average and the maximum values of intensity during the fall.

In respect of each rainfall the speed, which the ascending currents of saturated air might attain, has been derived and shown in Table 2 together with their duration in minutes. The figures in bracket indicate the number of occasions this value of the speed was found during the fall. It will be seen that the highest value of v_z is about 5.7 metres per sec., or about 12.7 miles per hour, which has been obtained during the rainfall on 22nd June, 1945. The average speed is found to be about 1.3 m./sec.

As the speed of ascending currents is closely related to intensity of rainfall the structure of convectional rain is of particular interest. Thunder showers occuring at Poona during 1939-45 were therefore grouped in Table 3 according to month, and from the records of Intensity Raingauge⁸ information was collected about the intensities and their duration during each fall. During heavy rain the showers are quite distinct. Duration of each shower has therefore been indicated by figures in bracket. It will be seen that thunder-storms during September and October generally give more intense rain at Poona and the maximum intensity of rain recorded there is about 6 in./hr. During the 35 thunder-storms 211 showers occurred which had intensity of 2 in /hr. or greater. These high intensity showers have been grouped in Table 4 according to the maximum intensity attained by them. The duration of showers of a certain intensity has been averaged and its mean value also given.

TABLE 3.

Duration, in minutes, of intensity of thunder showers at Poona. Figures in bracket indicate duration of separate showers.

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TABLE 4.

Thundershowers having intensity of 2 in. | hr. or greater.

It will be seen that out of 211 showers only 5 reached an intensity of 5.5 in./hr. while 56 of them reached a maximum intensity of 2 in /hr. only. Generally a shower with an intensity of 4 in./hr., or greater, lasts for 2 to 3 mins. and that with an intensity between 2 to 4 in./hr. lasts for about 3 minutes only.

$R E F E R E N C E S$:

- 1. Mon. Weather Rev. Washington 63, 10, 291 (1935).
- 2. Holmboe, J., Forsythe, G.E. and Gustin, W. Dysamic Meteorology, 141 (1945).
- 3. Jour. Sc. Instts., 22, 5, 92 (1945).