

Relationship between Vertical Currents and Intensity of Precipitation

J. M. SIL

(Received 27th August, 1946.)

RA^TE of condensation in a rising column of saturated air can be easily found out when it is assumed that the column is thermally and physically isolated. For purpose of evaluating the rate of precipitation further assumption is made, *viz.*, all the products of condensation are immediately detached from the system and fall out as precipitation. On these assumptions formulas for the amounts of precipitation have been given by several writers. Fulks¹ recently derived a workable formula from which the rate of precipitation can be evaluated when the physical state of the upper layers and the velocity of the ascending layer are known. The subject has been briefly discussed here following Holmboe².

Take a sample saturated air of $(1+w_s)$ tons having one ton of dry air and w_s ton of water vapour. While it ascends under pseudo-adiabatic process, precipitation from this mass will be at the rate of $-\frac{dw_s}{dt}$.

$$\text{We write } -\frac{dw_s}{dt} = -\frac{dw_s}{dz} \cdot \frac{dz}{dt} = -\frac{dw_s}{dz} \cdot v_z$$

For simplifying calculation we take $v_z = 1$ m/sec to begin with. We may then say that when the mass of $(1+w_s)$ tons of the saturated air is lifted one metre the amount of precipitation will be $-\frac{dw_s}{dz}$ ton. Per ton of saturated air, the precipi-

tation is therefore $-\frac{1}{1+w_s} \cdot \frac{dw_s}{dz}$ ton. If we consider in the above sample of air a column with cross-sectional area of one square metre and height of one dynamic—decimetre, the mass of the column will be $\frac{1}{\alpha g}$ ton. The precipitation from this column will therefore be

$$P_1 = -\frac{1}{\alpha g (1+w_s)} \cdot \frac{dw_s}{dz} \text{ or } = -\frac{1}{\alpha (1+w_s)} \cdot \frac{dw_s}{d\phi}$$

Here, α is the sp. vol. of the saturated air,

w_s the mixing ratio,

ϕ the dynamic height,

and g the acceleration of gravity.

During the pseudo-adiabatic process the heat released by condensation is used up in heating the entire mass of the air. Therefore $Ldw_s = (1+w_s)(C_p dT - \alpha dp)$ or

$$\text{dividing by } d\phi, \quad -L \frac{dw_s}{d\phi} = (1+w_s) \left(C_p \frac{dT}{d\phi} - \alpha \frac{dp}{d\phi} \right) \dots (2)$$

Eliminating $\frac{dw_s}{d\phi}$ between (1) and (2) and substituting

$$d\phi = -\alpha dp, \text{ we get, } P_1 = \frac{1+C_p \frac{dT}{d\phi}}{\alpha L}$$

Here, C_p is the sp. heat of moist air at constant pressure, and L the heat of condensation or of sublimation according as the process takes place in the rain stage of snow stage.

If we assume that the condensed water will fall on ground over an area of one square metre the rain collected per second will be $P_1 \cdot 10^9$ millimetres, or at the same rate the rainfall is $3.6 \times 10^9 P_1$ mm/hr.

For ease of calculation, a column of 100 dynamic metres in height, ascending with a velocity of v_z m/sec. may be taken. In that case the rate of precipitation becomes

$$P_{100} = 3.6 \times 10^9 \frac{1+C_p \frac{dT}{d\phi}}{\alpha L} \cdot v_z \dots (4)$$

$$\text{Or, putting } 3.6 \times 10^9 \frac{1+C_p \frac{dT}{d\phi}}{\alpha L} = r$$

We may write $P_{100} = r \cdot v_z$

Precipitation from the entire column would be $P = \sum r \cdot v_z \dots (5)$ when the total height of the saturated layer is broken up into layers of 100 dy.m. thickness. The value of r may be obtained from RAOB data and that of P readily from rainfall intensity record; the value of v_z can thus be determined.

Take the rainfall at 2130 hrs. I.S.T. on 22nd June, 1945. From radiosonde ascent at 20.30 hrs. we find the conditions of the upper layers as follows:—

p	t	u	p	t	u
942	25	85	600	4	100
900	21	90	500	-4	100
850	19	95	400	-14	..
750	13	100	300	-28	..
700	11	100	200	-46	..

Ascent curve indicated that the surface air had a lapse rate just greater than the dry-adiabatic. It was, therefore, indifferent and would rise with small perturbations. Rising from the ground a parcel of air would get saturated at about 900 mb. level and thereafter continue to rise due to buoyancy until it came to equilibrium with its environment at about 200 mb. level.

Assuming $v_z=1$ m/sec. the rate of precipitation has been calculated from the entire saturated layer as follows:—

Layer. mbs.	H dyn. m.	r mm./hr.	$\frac{Hr}{100}$
800—700	1188	0.75	8.9
700—600	1287	0.50	6.4
600—500	1428	0.44	6.2
500—400	1700	0.31	5.2
Total ..			26.7 mm/hr.

Sample calculation:—Take the layer 800—700 mb.

Mean value in the layer $p=75.0$ cb., $T=286.5$ K

$$w_s = 13.3 \times 10^{-3}$$

Thickness of the layer is given by

$$H_{700} - H_{800} = 28.7 (273 + t_m \text{ } ^\circ\text{C}) \ln \frac{p_{700}}{p_{800}}$$

$$= 28.7 (273 + 13.5) \ln \frac{8}{7} = 1188 \text{ dyn. metres.}$$

$$\frac{dT}{d\phi} = \frac{-5}{11880} = -0.42 \times 10^{-3} \text{ deg./dy. dm.}$$

$$C_p = C_{pd} (1 + 0.9w) = 1004 (1 + 0.9 \times 0.13) = 1016 \text{ kJ/ton.}$$

$$C_p \frac{dT}{d\phi} = -1016 \times 0.42 \times 10^{-3} = -0.426$$

$$\alpha = \frac{R_d T}{p} \frac{287 + 286.5 (1 \times 0.61 \times 0.13)}{75.0} = 1107 \text{ m}^3/\text{ton}$$

$$r = 3.6 \times 10^9 \times \frac{1 + C_p \frac{dT}{d\phi}}{\alpha L} = 3.6 \times 10^9 \times \frac{1 - 0.426}{1107 \times 2.5 \times 10^6} = 0.746 \text{ mm./hr.}$$

The rain commenced at 2130 hrs., *i.e.*, about one hour after the ascent was taken. The structure of the rainfall as shown by the intensity record is given below :—

Date	Rainfall		Periods of intensity in minutes							
	Total amt. in inches	Duration	0.5"/hr	1"/hr.	2"/hr.	3"/hr.	4"/hr.	5"/hr.	6"/hr.	
22nd June, 1945.	1.2	52 mins.	9	7	5	3				1st shower.
			25	17	10	5	3	2	2	2nd shower.

The average rate of rainfall, as found from the record, is $\frac{1.2 \times 60}{52} = 1.39$ in./hr. or = 35.2 mm./hr. Rate of precipitation, as calculated, is 26.7 mm./hr. The average rate of ascent of the saturated column is therefore v_z av. = $\frac{35.2}{26.7} = 1.3$ m./sec.

During the fall the maximum intensity attained was 6 in./hr. or 152 mm./hr. for a very short time. It will not be incorrect to assume that only v_z changed during the period and that other factors contributing to the rate of precipitation remained more or less constant. Thus, v_z max. = $\frac{152}{26.7} = 5.7$ m./sec.

A few rainfall records have been studied together with the upper-air conditions prevailing shortly before the fall, and proceeding in the way mentioned above the speed values of the vertical currents of saturated air have been derived. Unfortunately the number of occasions selected for study so far has not been large. However, the results of study are given in Tables 1 and 2 below :

TABLE 1.
Structure of Rainfall.

Date	Rainfall			Duration of Intensity of Rainfall in minutes							Average	Max.
	Time I S.T.	Duration in h. m.	Amt. in inches	0.5"/hr.	1"/hr.	2"/hr.	3"/hr.	4"/hr.	5"/hr.	6"/hr.		
22-6-1945	21-30	0—52	1.2	34	24	15 (5+5 +5)	8 (3+5)	3	2	2	1.4	6.0
6-7-1945	19-30	0—55	0.9	26	22	11 (3+2 +6)	3	2	0.88	4.0
17-7-1945	22-00	0—45	1.3	38	19	9 (2+3 +2 +2)	2	1.73	3.0
21-7-1945	22-11	3—48	4.87	134	92	31 (5+5 +2 +3+2 +2+3 +4 +5)	12 (2+3 2+7)	6 (2+4)	1.28	4.0
14-8-1945	19-45	1—17	1.3	57	23	8 (3+3 +2)	3	2	1.0	4.0

TABLE 4.

Thunderstorms having intensity of 2 in./hr. or greater.

Intensity attained		5.5"/ hr.	5"/ hr.	4.5"/ hr.	4"/ hr.	3.5"/ hr.	3"/ hr.	2.5"/ hr.	2"/ hr.
Number of showers	..	5	10	14	21	28	36	41	56
Per cent	..	2.4	4.7	6.6	10.0	13.3	17.0	19.5	26.6
Average duration in mins. during a shower	..	2.6	2.9	3.07	3.0	3.04	3.1	3.8	3.94

It will be seen that out of 211 showers only 5 reached an intensity of 5.5 in./hr. while 56 of them reached a maximum intensity of 2 in./hr. only. Generally a shower with an intensity of 4 in./hr., or greater, lasts for 2 to 3 mins. and that with an intensity between 2 to 4 in./hr. lasts for about 3 minutes only.

REFERENCES:

1. *Mon. Weather Rev. Washington* 63, 10, 291 (1935).
2. Holmboe, J., Forsythe, G.E. and Gustin, W. *Dynamic Meteorology*, 141 (1945).
3. *Jour. Sc. Instts.*, 22, 5, 92 (1945).