



## Trend assessment of rainfall, temperature and relative humidity using non-parametric tests in the national capital region, Delhi

JITENDRA RAJPUT, N. L. KUSHWAHA\*, D. R. SENA, D. K. SINGH and INDRA MANI

*Division of Agricultural Engineering, ICAR-IARI, Pusa, New Delhi – 110 012*

*(Received 27 December 2021, Accepted 22 November 2022)*

\*e mail : [nand.kushwaha@icar.gov.in](mailto:nand.kushwaha@icar.gov.in)

सार — जल संसाधन उपयोग की योजना और सिंचाई संबंधी निर्णय लेने के लिए स्थानीय, क्षेत्रीय और वैश्विक पैमानों पर वर्षा और तापमान की स्थानिक-अस्थायी विविधताओं को समझना महत्वपूर्ण है। वर्तमान जाँच में भारत के राष्ट्रीय राजधानी क्षेत्र (एनसीआर), दिल्ली, में 31 वर्षों (1990-2020) में आईएआरआई मौसम विज्ञान केंद्र, पूसा, नई दिल्ली से प्राप्त वर्षा, तापमान और सापेक्ष आर्द्रता की प्रवृत्तियों और विविधताओं की पहचान करने का प्रयास किया गया है। समय के साथ परिवर्तन की प्रवृत्ति दिशा और परिमाण का आकलन करने के लिए वार्षिक, मासिक और ऋतुनिष्ठ विश्लेषण हेतु थैल सेन ढलान परीक्षण के बाद मान-केंडल (MK) परीक्षण का उपयोग करके सांख्यिकीय प्रवृत्ति विश्लेषण किया गया। चर की समय श्रृंखला में विभक्ति बिंदु का पता लगाने के लिए पेटिट परीक्षण किया गया है। वार्षिक अधिकतम तापमान ( $T_{max}$ ), न्यूनतम तापमान ( $T_{min}$ ) और वर्षा के डेटा में कोई प्रवृत्तियाँ नहीं पाई गईं जबकि, वार्षिक सापेक्ष आर्द्रता में सकारात्मक वृद्धि की प्रवृत्ति दिखाई दी। जनवरी और दिसंबर के सर्दी के महीनों में अधिकतम तापमान ( $T_{max}$ ) में सांख्यिकीय रूप से महत्वपूर्ण गिरावट की प्रवृत्ति का पता चला है। इसका तात्पर्य जनवरी और दिसंबर के सर्दी के महीनों में तापमान में गिरावट से है। साथ ही, अधिकतम तापमान ( $T_{max}$ ) के ऋतुनिष्ठ विश्लेषण से शीत ऋतु में कमी की प्रवृत्ति का पता चला है। इसी तरह, न्यूनतम तापमान ( $T_{min}$ ) से जनवरी और दिसंबर के सर्दी के महीनों में कमी की प्रवृत्ति का पता चला है। हालांकि, अप्रैल में न्यूनतम तापमान ( $T_{min}$ ) में सांख्यिकीय रूप से वृद्धि की प्रवृत्ति देखी गई, जो बढ़ते तापमान के कारण जायद मौसम की फसलों की खेती के लिए प्रतिकूल वातावरण पैदा कर सकती है। इसके अलावा, न्यूनतम तापमान ( $T_{min}$ ) के ऋतुनिष्ठ विश्लेषण से मॉनसून, शरद ऋतु और शीतऋतु के दौरान कमी की प्रवृत्ति का पता चला है। सापेक्ष आर्द्रता ने फरवरी, जून, जुलाई, अगस्त, सितंबर और मॉनसून ऋतु को छोड़कर वार्षिक, मासिक और ऋतुनिष्ठ आधार पर सकारात्मक वृद्धि की प्रवृत्ति दर्शाई। प्रवृत्ति परिवर्तन विश्लेषण में वार्षिक अधिकतम तापमान और न्यूनतम तापमान के डेटा श्रृंखला में कोई परिवर्तन बिंदु नहीं दिखाई दिया। यह देखा गया कि 2005 के अगस्त में न्यूनतम तापमान के डेटा में परिवर्तन की प्रवृत्ति है। साथ ही, न्यूनतम तापमान ( $T_{min}$ ) डेटा श्रृंखला के लिए मॉनसून और शीतऋतु में परिवर्तन बिंदु का पता चला है। वार्षिक औसत वर्षा डेटा में परिवर्तन बिंदु को 2012 में चिह्नित किया गया। सापेक्ष आर्द्रता में भी अप्रैल, नवंबर में वार्षिक और मॉनसून ऋतु के पैमाने पर डेटा श्रृंखला में परिवर्तन बिंदु दिखे। इस शोध पत्र में जल विज्ञान डिजाइन और फसल सिंचाई निर्णयों पर इन परिवर्तनों के शिक्षाप्रद प्रभाव पर चर्चा की गई है।

**ABSTRACT.** Understanding rainfall and temperature spatio-temporal variations at the local, regional, and global scales is vital for water resource utilization planning and making irrigation decisions. The present investigation attempts to identify trends and variations in the rainfall, temperature and relative humidity over 31 years (1990-2020) in the National Capital Region (NCR), Delhi, India, obtained from IARI meteorological station, Pusa, New Delhi. The statistical trend analyses were carryout using Mann-Kendall (MK) test followed by Theil Sen slope test for annual, monthly and seasonal analysis to assess the trend direction and magnitude of the change over time. Pettitt's test employed to detect the inflection point in the variable's time series. The annual  $T_{max}$ ,  $T_{min}$  and rainfall showed an absence of trends in the data. While, annual relative humidity showed a positive increasing trend.  $T_{max}$  indicated a statistically significant decreasing trend in the winter months of January and December. This implies a dip in the temperature in the winter months of January and December. Also,  $T_{max}$  seasonal analysis revealed a decreasing trend during winter season. Similarly,  $T_{min}$  revealed a decreasing trend for the winter months of January and December. Although, a statistically increasing trend for  $T_{min}$  was seen in April, which may cause a harsh environment for cultivating the Zaid season crops due to increased warming. In addition,  $T_{min}$  seasonal analysis revealed a decreasing trend during monsoon, autumn and winter seasons. Relative humidity showed positive increasing trends on annual, monthly, and seasonal basis except during February, June, July, August, September and Monsoon season. The trend change analysis showed no change point in the annual  $T_{max}$  and  $T_{min}$  data series. It was observed that  $T_{min}$  in August showed a change point in the data trend in 2005. Also, the change point was detected in monsoon and winter seasons for  $T_{min}$  data series. The change point in the

annual average rainfall data was marked in 2012. Relative humidity also showed change points in data series in April, November, Annual and Monsoon season scales. A didactic implication of these changes on hydrologic design and crop irrigation decisions was discussed in this paper.

**Key words** – Trend analysis, Man-Kendall test, Pettitt test, Theil-sen's slope, Climate analysis, India.

## 1. Introduction

Recent years have indicated a severe challenge to global food production due to alterations in climatic behavior besides others. Hindrances posed by climate change at various scales to agricultural sustainability attracted scientific intellect to pursue intensive efforts to understand their pattern and behavior to offset its effects (Kushwaha *et al.*, 2016). Climate change has caused several impacts on land, water, and the environment at the global level. It threatens human existence due to food insecurity, water scarcity, farming failures, extreme events, infectious diseases, biodiversity loss and displacement due to the availability of deprived resources (Kushwaha *et al.*, 2022). Rainfall and temperature are commonly utilized to explain climate change impacts water resources. Different studies have indicated that temperatures and rainfall have consistently registered monotonous changes to their trends in the past decades (Pingale *et al.*, 2014). New *et al.* (2001) retrieved rainfall data from a different nation to detect rainfall trends and they reported an increasing trend during the 20<sup>th</sup> century. A 150 years study led by authors Jones *et al.* (1999) on surface temperatures showed that the average temperature increased by 0.37 °C and 0.32 °C during 1925-1944 and 1978-1997 at the global scale. Several studies have been performed in India to assess the spatio-temporal variability of rainfall at various scales. According to Lal (2001), the random variability in rainfall in India during the past 100 years has no systematic shift either on a yearly or seasonal scale. However, rainfall fluctuation was noticed in India on temporal and spatial scales (Pattanaik, 2005). The rainfall trend analysis of 45 districts in Madhya Pradesh, India, was conducted, and 1978 was determined the most likely year of change in yearly rainfall (Duhan and Pandey, 2013). Researchers (Rajani *et al.*, 2020) applied discrete wavelet transform (DWT) based Mann-Kendall test in the central Gujarat region for long-term trend analysis of rainfall. They reported a significant decrease in the rainfall amount across the stations. District level climate variables trend assessment helps develop proper agriculture plans to safeguard the crops and achieve better crop production (Khavse and Chaudhary, 2022). Extreme long-term records tend to absorb the trends due to a possible cyclic trend acquaintance. At the same time, a short-term, especially a decadal change, is likely to exhibit a seasonal tendency in variations as reported by Mahapatra and Mohanty (2006), found that the seasonal rainfall variation

was more significant from 1980 to 1999 than from 1901-1990. Another authors (Phukan and Saha, 2022) analyzed rainfall trends in the rainfall over Tripura. They reported increasing and decreasing trends at the different stations within the state, which shows the spatial variability in the rainfall. Though, there was no significant trend was seen in the mean seasonal rainfall but substantial increase in the heavy rainfall events-based hazards over the central India would occur (Goswami *et al.*, 2006). The rainfall data sources (resolutions of products) also affect the trend assessment (Li *et al.*, 2022).

Most temperature trend analyses in India or elsewhere focus on analyzing seasonal and yearly temperature trends considering a single station or a cluster of stations. The study by Srivastava *et al.* (1992) on decadal climate trends over India offered the first clear signs that the diurnal discrepancy of surface temperatures over India is quite distinctive from that over most world regions. Pant and Kumar (1997) researchers examined yearly and season-wise air temperature data from 1881 to 1997 and discovered a significant temperature increase of 0.57°C per century. Author Jaswal (2010) evaluated India's monthly maximum and minimum temperature and found variability in the warmer months across the northern and southern India. Sharma and Chaudhary, 2014 reported a cooling trend in the winter temperature in the Bastar plateau agro-climatic zone of Chhattisgarh. Rupa Kumar *et al.* (1994) found a rising trend in the T<sub>max</sub> data and no trend for T<sub>min</sub> at 121 stations in India from 1901 to 1987, causing an increase in average and diurnal temperature spectrum. Relative humidity is a crucial climatic variable that governs the aerodynamics portion of the atmospheric evaporating needs (Wang and 40 Dickinson, 2012). According to Talae *et al.* (2012), the yearly relative humidity in Iran's southern and northern coastal regions risen by 1.03%/decade. Kaur *et al.* (2022) concluded that relative humidity was trending upward, while wind speed, pan evaporation and sunlight hours were trending downward, suggesting that overcast and damp weather conditions were becoming more common in the Punjab state. The non-parametric Mann-Kendall test (MK test) is frequently used for trend assessment in meteorological variables (Gellens 2000; Das *et al.*, 2008; Das *et al.*, 2020; Shadmani *et al.*, 2012; Pingale *et al.*, 2014; Korade and Dhorde 2016). The possible explanation for using this non-parametric statistical test is that it can see monotonic trends (both increasing and decreasing) in the variable's data. Many studies employed Theil-Sen

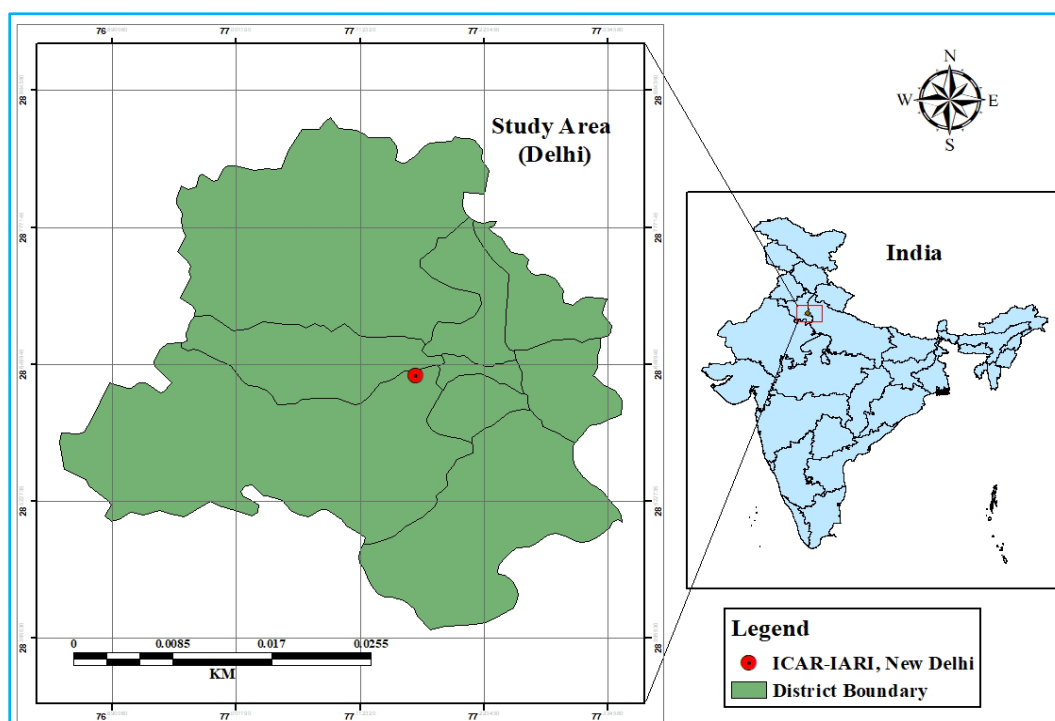


Fig. 1. Study area and IARI meteorological station map

approach for calculating the actual magnitude of slope (Hirsch *et al.*, 1982; Gallego *et al.*, 2006; Chen & Grasby 2009; and Shadmani *et al.*, 2012). Pettitt's (1979) test identifies shifting points in data series (Zhang *et al.*, 2008; Gebremicael *et al.*, 2013).

In a motivation derived from the above studies, the present study utilized 31 years of rainfall, temperature and relative humidity (average) data to investigate trends in a sub-tropical climate region of Delhi, India. The MK and Theli-Sen slope tests were employed to identify trend identification and rate estimation in the variables data. At the same time, the Pettitt technique was followed to see the inflection point describing the shifting in the time-series data.

## 2. Materials and methods

### 2.1. Study area description

Study area IARI Pusa is located in the semi-arid region of the National Capital Region (NCR), Delhi, India. Delhi has a semi-arid, subtropical climate with hot, dry summers and cold winters and is located in the Agro-ecoregion-IV. The annual average temperature is 25.5° C. May, and June is the warmest months, with an average maximum temperature of 39° C in the range of 43.9 to 45.0° C for the last 30 years. The winter season lasts from December through February. January's average

temperature of 14° C is regarded as the coldest. On the other hand, the lowest temperature falls below 1° C. 75% of the annual rainfall (710 mm) occurs during the monsoon season (June to September). The location map of the selected study area is shown in Fig. 1.

### 2.2. Dataset

Long-term rainfall, temperature and relative humidity data (observed daily for 31 years) were investigated to detect trends during 1991 to 2020 in National Capital Region, Delhi, India. The data were acquired from the meteorological station, Division of Agricultural Physics, IARI, Pusa, Delhi, India. Tables 1, 2, 3 and 4 summarizes central tendency and dispersion measures for  $T_{max}$ ,  $T_{min}$ , rainfall and relative humidity (average). The annual rainfall was maximum with the value of 1524.3 mm in 2013 and yearly minimum rainfall amounting to 452.7 mm was obtained in 2002. The analysis of the present investigation is the monthly and annual minimum temperature ( $T_{min}$ ) and the maximum temperature ( $T_{max}$ ) during 1990-2020. Annually maximum temperature observed was 42.60 °C in 2012, whereas the minimum temperature was found in 2005. Relative humidity (RHavg.) analysis showed that monthly average minimum and maximum values were found as 46.35 and 77.35 % in May and August, respectively.

TABLE 1

Monthly values of measures of central tendency and dispersion for Tmax

Parameter	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec
Mean	19.51	23.27	29.15	36.26	39.65	38.78	34.89	33.44	33.51	32.66	27.76	22.20
Median	19.60	23.10	29.20	36.00	39.75	38.90	34.30	33.50	33.60	32.70	27.80	22.60
Mode	18.60	23.10	29.90	35.40	40.10	40.30	34.30	33.60	33.40	32.60	26.90	22.6
Std. Deviation	1.52	1.51	1.79	1.61	1.34	2.06	1.55	1.01	1.15	1.14	0.85	1.51
Variance	2.30	2.29	3.19	2.59	1.79	4.26	2.39	1.02	1.32	1.29	0.72	2.29
Skewness	0.04	0.94	0.47	0.40	-0.35	-0.40	1.24	0.47	-0.27	-1.20	-0.51	-1.75
Kurtosis	-0.10	2.34	-0.37	0.30	0.76	-0.02	1.73	-0.08	0.36	2.91	0.90	3.31
Range	6.60	7.40	6.50	7.50	6.40	8.80	6.90	4.20	5.00	5.80	3.80	6.20
Minimum	16.30	20.80	26.80	32.80	36.10	33.80	32.50	31.60	30.90	28.90	25.30	17.70
Maximum	22.90	28.20	33.30	40.30	42.50	42.60	39.40	35.80	35.90	34.70	29.10	23.90
25th percentile	18.45	22.25	27.60	35.40	38.80	37.35	33.90	32.60	32.89	32.24	27.30	21.75
50th percentile	19.60	23.10	29.20	36.00	39.75	38.90	34.30	33.50	33.60	32.70	27.80	22.60
75th percentile	20.55	24.35	30.00	37.30	40.60	40.30	35.50	33.95	34.10	33.40	28.23	23.20

TABLE 2

Monthly values of measures of central tendency and dispersion for Tmin

Parameter	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec
Mean	6.48	9.14	13.58	19.28	24.19	26.77	26.70	25.95	23.93	17.41	11.49	6.84
Median	6.70	9.20	13.40	19.10	24.00	26.80	26.60	25.90	24.00	17.50	11.60	6.80
Mode	5.50	8.40	13.70	19.10	23.90	27.10	25.70	25.70	24.00	18.70	10.60	6.10
Std. Deviation	1.11	1.11	0.96	1.18	1.05	1.04	1.07	0.62	0.78	1.29	1.38	1.24
Variance	1.23	1.23	0.92	1.40	1.11	1.08	1.14	0.38	0.60	1.67	1.90	1.55
Skewness	-0.01	0.10	0.64	0.13	0.29	0.05	1.23	-0.57	0.38	-0.26	1.26	-0.13
Kurtosis	-0.56	-0.11	0.64	-0.30	0.59	-0.21	3.96	0.88	1.27	-0.88	3.46	-0.50
Range	4.67	4.80	4.20	4.70	5.10	4.40	6.00	2.80	3.70	4.90	7.20	5.20
Minimum	4.30	6.90	11.60	17.10	21.80	24.60	24.40	24.20	22.30	14.90	9.00	4.10
Maximum	8.97	11.70	15.80	21.80	26.90	29.00	30.40	27.00	26.00	19.80	16.20	9.30
25th percentile	5.50	8.40	13.00	18.65	23.50	26.05	26.14	25.65	23.55	16.40	10.60	6.10
50th percentile	6.70	9.20	13.40	19.10	24.00	26.80	26.60	25.90	24.00	17.50	11.60	6.80
75th percentile	7.25	9.80	13.85	20.10	24.90	27.34	27.15	26.45	24.34	18.46	12.10	7.90

<sup>a</sup> More than one mode exists, only the first is reported

### 2.3. Methodology

Trend analysis, a technique for determining the spatio-temporal variations in several climate parameters, is the simplest procedure to detect the variation of

variables in the data time series. This is indeed a critical problem for a country such as India, as our economy relies primarily on agriculture, heavily reliant on monsoon rainfall. Hence, any shift in that time of the year can devastate the nation's farming prospects, therefore, the

**TABLE 3**

**Monthly values of measures of central tendency and dispersion for rainfall**

Parameter	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec
Mean	19.73	24.21	24.03	9.15	33.39	79.74	217.43	236.52	121.99	16.76	4.87	8.78
Median	11.40	13.00	8.00	4.00	27.20	72.80	181.20	219.10	105.90	0.00	0.00	0.20
Mode	0.00	0.00	0.00	0.00	0.00	4.60	7.20	66.30	5.40	0.00	0.00	0.00
Std. Deviation	22.37	28.91	46.70	12.72	35.30	60.95	149.12	108.73	85.35	28.14	9.07	18.27
Variance	500.36	835.51	2181.24	161.80	1246.19	3714.34	22238.00	11822.83	7284.81	791.96	82.25	333.82
Skewness	1.08	1.42	3.13	2.07	1.35	1.29	0.83	0.50	0.39	1.86	2.26	2.49
Kurtosis	0.05	1.49	9.69	4.24	1.56	2.05	0.13	0.18	-0.90	3.10	4.96	5.63
Range	72.90	109.40	201.80	51.80	136.60	266.40	571.40	455.60	308.80	109.00	37.00	68.00
Minimum	0.00	0.00	0.00	0.00	0.00	4.60	7.20	66.30	5.40	0.00	0.00	0.00
Maximum	72.90	109.40	201.80	51.80	136.60	271.00	578.60	521.90	314.20	109.00	37.00	68.00
25th percentile	1.90	1.70	1.10	0.90	5.40	34.05	122.65	162.95	50.75	0.00	0.00	0.00
50th percentile	11.40	13.00	8.00	4.00	27.20	72.80	181.20	219.10	105.90	0.00	0.00	0.20
75th percentile	33.20	31.20	21.35	11.80	46.35	104.60	294.45	307.35	191.00	27.50	5.40	6.30

<sup>a</sup> More than one mode exists, only the first is reported

**TABLE 4**

**Monthly values of measures of central tendency and dispersion for relative humidity**

Parameters	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec
Mean	73.25	68.15	60.39	47.87	46.35	55.35	73.40	77.32	72.58	62.65	63.72	69.45
Median	72.60	67.80	60.24	47.37	47.31	54.48	75.00	79.19	72.90	61.50	63.50	68.77
Mode	62.60	70.39	50.19	38.65	42.82	36.32	52.53	65.36	70.38	53.84	52.73	54.90
Std. Deviation	5.94	7.13	7.42	7.11	6.93	9.33	6.65	5.53	7.21	5.52	5.70	5.65
Variance	35.23	50.87	55.10	50.48	47.99	87.03	44.16	30.55	51.96	30.52	32.54	31.93
Skewness	-0.08	0.54	0.51	0.10	0.00	-0.18	-1.51	-0.50	-0.35	0.46	0.11	-0.13
Kurtosis	-1.17	0.54	-0.72	-1.39	-0.64	-0.30	2.64	-0.16	-0.44	-0.57	-0.32	0.47
Range	19.81	30.41	26.27	22.27	26.86	37.90	29.50	23.29	28.53	21.39	23.78	26.29
Minimum	62.60	54.40	50.19	37.75	33.24	36.32	52.53	65.36	57.25	53.84	52.73	54.90
Maximum	82.40	84.80	76.47	60.02	60.10	74.22	82.03	88.65	85.78	75.23	76.52	81.20
25th percentile	68.48	63.82	55.13	41.49	41.70	50.67	70.43	73.33	68.81	58.63	59.58	66.69
50th percentile	72.60	67.80	60.24	47.37	47.31	54.48	75.00	79.19	72.90	61.50	63.50	68.77
75th percentile	78.23	70.25	64.84	54.61	50.21	61.86	77.56	80.55	77.38	66.77	68.33	73.32

economy. So, trend analysis is capable of resolving global climate change challenges. The main objective of this trend investigation is to analyse changes in meteorological parameters (rainfall temperature and relative humidity). Thirty-one years of meteorological observations were analyzed to examine rainfall and temperature data

patterns. The MK test was employed to detect the trend, and the magnitude of slopes was analyzed employing the Theli-Sen technique.

Additionally, the Pettitt test was employed to determine the time series' change point (Pettitt 1979).

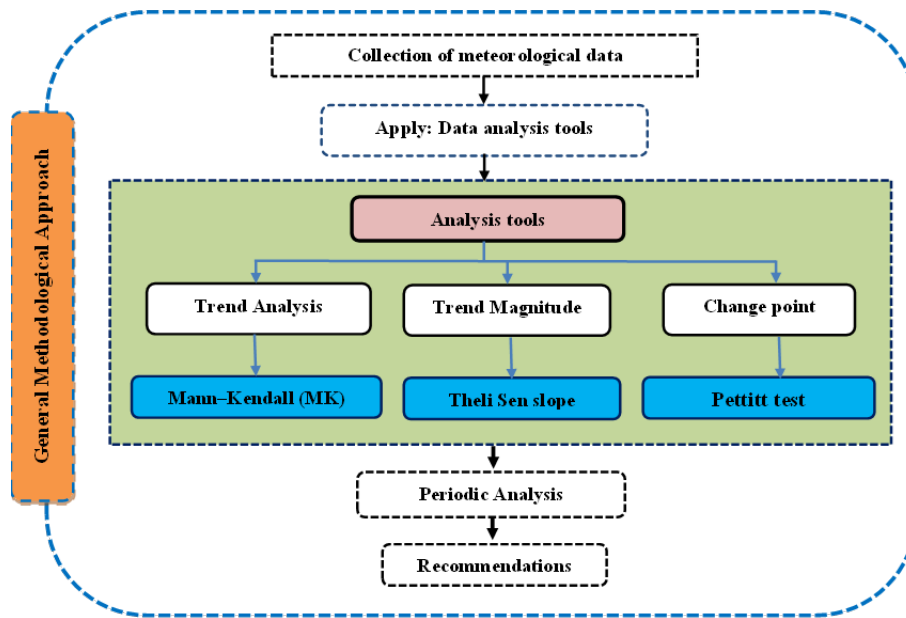


Fig. 2. Flow diagram of the methodology

The flow methodology diagram followed in this investigation has displayed in Fig. 2.

2.3.1. Mann-Kendall test

The MK is a popular technique to identify trends in a time series dataset. This method finds a monotonic trend in the investigation parameter during the intended period. A monotonic upward (downward) trend indicates that the variable continuously rises (falls) over the period. Also, the null hypothesis H0 and alternative hypothesis Ha denote the absence and presence, respectively. The MK test is assessed as per the Shadmani *et al.* (2012):

$$\text{sgn}(x_i - x_j) = \begin{cases} +1 & \text{if } x_j > x_i \\ 0 & \text{if } x_j = x_i \\ -1 & \text{if } x_j < x_i \end{cases}$$

$$S = \sum_{i=1}^{n-1} \sum_{j=i+1}^n \text{sgn}(x_i - x_j)$$

where  $x_i$  and  $x_j$  show the data at times  $i$  and  $j$ , respectively and  $n$  is the data period of the data. If  $S$  is +, means the variable continuously rises in time;  $a - S$  value means a declining trend. The following equation is utilized in cases where  $n$  is more significant than 10.

$$\text{Var}(S) = \frac{n(n-a)(2n+5) - \sum_{i=1}^p t_i(t_i-1)(2t_i+5)}{18}$$

where  $p$  means the number of tied groups,  $t_i$  is the number of data values in the  $p$ th group. When the variance of time is provided in the preceding equation, the standard  $Z$  can be expressed by the following equation:

$$Z = \begin{cases} \frac{S-1}{\sqrt{\text{Var}(S)}} & \text{if } S > 0 \\ 0 & \text{if } S = 0 \\ \frac{S+1}{\sqrt{\text{Var}(S)}} & \text{if } S < 0 \end{cases}$$

2.3.2 Theil-Sen approach

This approach is utilized to compute the slope magnitude after the trend tests. The Theil-Sen techniques shown by the equation (Shadmani *et al.*, 2012) is as follows:

$$\beta = \text{median} \left( \frac{X_j - X_i}{j - i} \right) \quad \text{for all } i < j$$

where  $X_i$  and  $X_j$  means the sequential data values of the time series in the years  $i$  and  $j$ , respectively, the estimated  $\beta$  is estimated slope trend magnitude in the data series.

2.3.3. Pettitt's test

Pettitt's method is widely utilized to identify shifting points in variables time series values. For a time series

TABLE 5

Trend analysis of temperature, rainfall and relative humidity

Tests	Tmax			Tmin			Rainfall			RHavg		
	M-K Test Value (S)	Theli Sen slope	Trend	M-K Test Value (S)	Theli Sen slope	Trend	M-K Test Value (S)	Theli Sen slope	Trend	M-K Test Value (S)	Theli Sen slope	Trend
Jan	-112	-0.05	-	-100	-0.0381	-	42	0.18	N	133	0.3035	+
Feb	21	0.0061	N	-4	0	N	-71	-0.3077	N	70	0.124	N
Mar	62	0.0286	N	36	0.01	N	102	0.48	+	105	0.3109	+
Apr	41	0.0271	N	153	0.05	+	102	0.1833	+	142	0.3476	+
May	20	0.0059	N	-75	-0.0292	N	45	0.4571	N	144	0.3407	+
Jun	4	0	N	-53	-0.0222	N	-3	-0.02	N	81	0.2264	N
Jul	14	0	N	-92	-0.0261	N	107	5.1529	+	85	0.1718	N
Aug	41	91	N	-79	-0.2	N	31	0.85	N	29	0.0436	N
Sep	50	0.0211	N	-57	-0.0125	N	-11	-0.5238	N	40	0.1024	N
Oct	49	0.0167	N	3	0	N	1	0	N	157	0.2648	+
Nov	-67	-0.02	N	-31	-0.015	N	-25	0	N	101	0.2188	+
Dec	-117	-0.04	-	-109	-0.0556	-	6	0	N	113	0.2032	+
Annually average	-20	-0.0026	N	-61	-0.0094	N	97	6.47	N	201	0.2546	+
Spring	57	0.0286	N	1	0	N	49	1.9983	N	105	0.237	+
Summer	23	0.0095	N	7	0.0014	N	37	1.2083	N	161	0.2905	+
Monsoon	53	0.0159	N	-108	-0.0235	-	71	5.6	N	91	0.1568	N
Autumn	-8	-0.0015	N	-100	-0.032	-	-14	-0.159	N	137	0.2507	+
Winter	-148	-0.0577	-	-146	-0.0375	-	-11	-0.1167	N	170	0.2727	+

(Note : N= No statistically significant trend, + = Increasing trend, - = Decreasing trend,  $\alpha = 5\%$ , Spring-February to March, Summer: April to June, Monsoon: July to September, Autumn: October to November, and Winter: December to January)

like  $X_1, X_2, \dots, X_n$  with a length  $n$ , let  $t$  be the time of the change point. Then, the samples  $(X_1, X_2, \dots, X_t)$  and  $(X_{t+1}, X_{t+2}, \dots, X_n)$  can be obtained by changing the time series at time  $t$ . The test statistic  $U_t$  was given by (Chen *et al.*, 2009) and calculated as follows:

$$U_t = \sum_{i=1}^t \sum_{j=i+1}^n \text{sgn}(x_i - x_j)$$

$$\text{sgn}(x) = \begin{cases} +1 & \text{if } x > 0 \\ 0 & \text{if } x = 0 \\ -1 & \text{if } x < 0 \end{cases}$$

At time  $t$ , the greatest  $jU_j$  can be regarded as the most important change point. The following equation may be used to estimate the significant change probability  $P(t)$  at the change point (Chen *et al.*, 2009):

$$P(t) = 1 - \exp\left(\frac{-6U^2}{n^3 + n^2}\right)$$

When the approximated probability exceeds  $(1-\alpha)$ , the change point is taken at a statistically significant level of  $\alpha$ .

### 3. Results and discussion

#### 3.1. Temperature trend analysis

The annual, monthly temperature data (maximum and minimum) showed no significant trend, implying no considerable variation in the annual minimum and maximum temperature during the investigation period. After the MK test had identified the trends, Theli Sen slope was estimated to quantify the slope magnitude. The MK Test value and Theli Sen slope for maximum temperature are shown in Table 5. A statistically decreasing trend was observed for the maximum temperature ( $T_{max}$ ) during January and December (Fig. 3). The MK test value for January was -112 and Theli Sen slope magnitude of -0.0500 during the period, while for December month, the MK test value obtained was -117 and Theli Sen slope magnitude of -0.04. Results indicated a decreasing  $T_{max}$  during the winter, affecting



**Figs. 3(a&b).** The significant trend observed in monthly  $T_{max}$  time series data in (a) January and (b) December

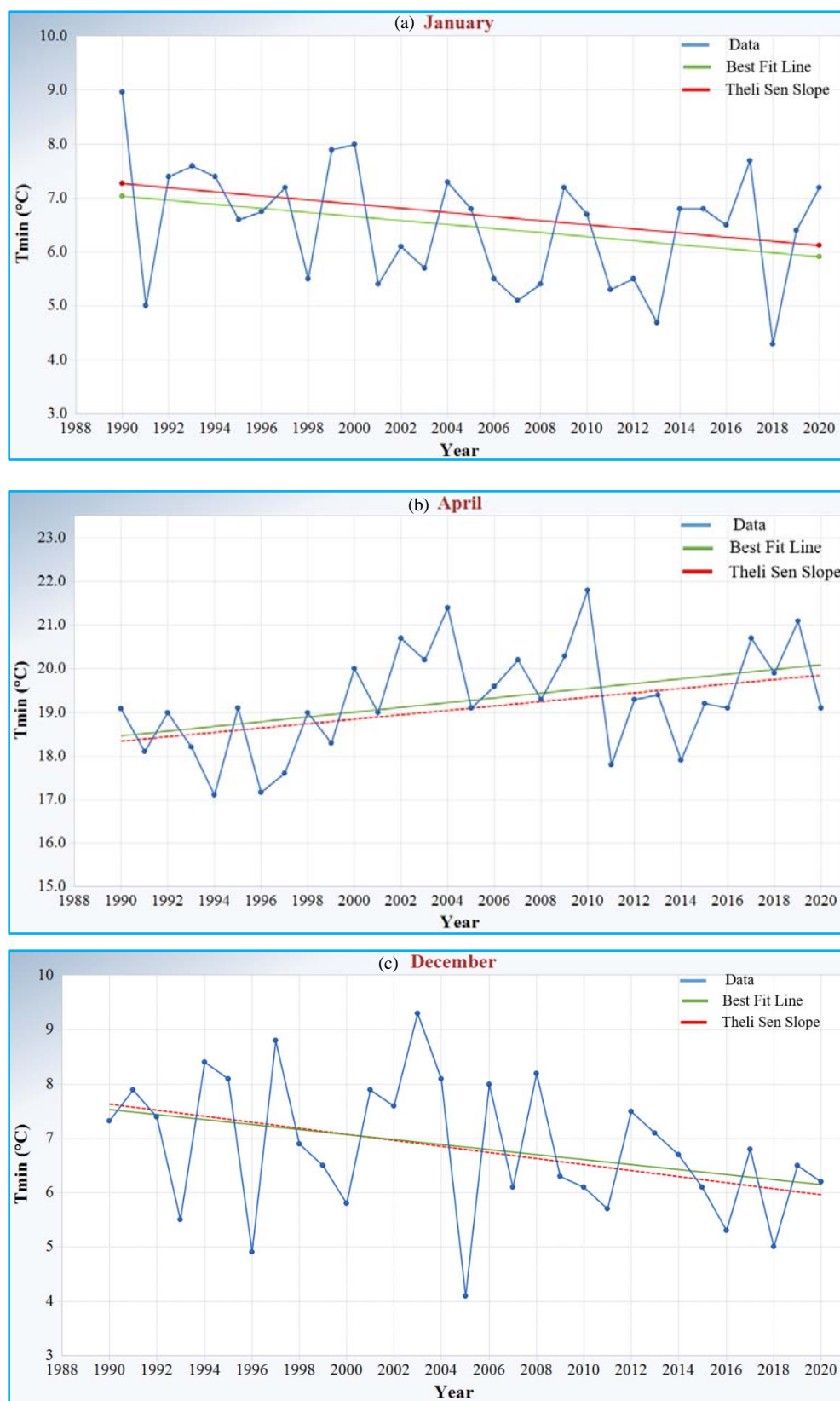
agriculture crop yield. Trend analysis of  $T_{min}$  revealed a significant declining trend for January with an MK test value of -100 and Theli Sen slope of -0.0381. A statistically positive increasing trend was seen in April. The MK test value of 153 and Theli Sen slope of 0.050 for April was observed. A statistically decreasing trend was identified for December. The MK test value of -109 and Theli Sen slope of -0.0556 was found for December. For the remaining months, no trend was seen for  $T_{max}$  and  $T_{min}$ . Seasonal analysis of  $T_{max}$  time series data showed that there has been a statistically decreasing trend for winter season. This downward temperature trend may affect vegetables which are sensitive to lower temperatures. Similarly, for  $T_{min}$ , seasonal analysis showed statistically decreasing trend during monsoon,

autumn and winter seasons. Though decreasing temperatures could lower the crop water demands but may affect the crop development by hampering the growing degree days requirements to achieve biological crop maturity. Another negative effect associated with lower temperatures is extension of crop durations. Fig. 4 shows the significant trends observed in the monthly  $T_{max}$  data series.

### 3.2. Trend change point analysis of temperature data

The change point analysis of temperature (maximum and minimum) data was performed using the Pettitt test. The Pettitt test showed no change point for annual  $T_{max}$





**Figs. 4(a-c).** The significant trend observed in monthly  $T_{max}$  time series data in (a) January, (b) April and (c) December

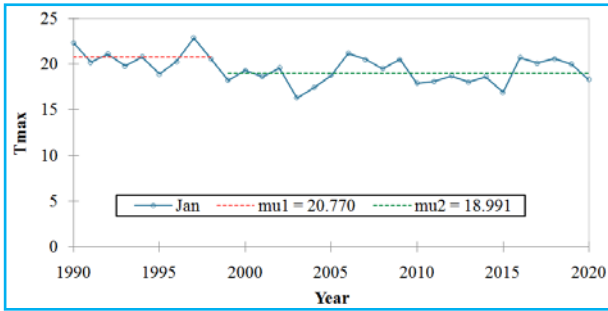


Fig. 5. Trend change point identification in  $T_{max}$  monthly time-series data

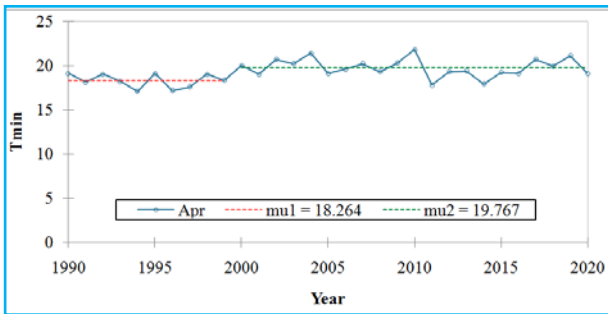


Fig. 6. Trend change point identification in  $T_{min}$  monthly time-series data

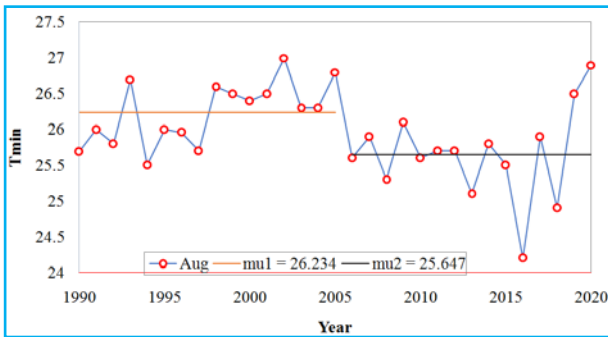


Fig. 7. Trend change point identification in  $T_{min}$  monthly time-series data

and  $T_{min}$ . However, monthly  $T_{max}$  and  $T_{min}$  showed a shifting point in the time series trend. For the  $T_{max}$ , a trend change point was observed for January 1998 (Fig. 5). The average  $T_{max}$  value before the change point (1990-1998) was 20.770 ( $^{\circ}\text{C}$ ), while the average  $T_{max}$  value post the change point (1999-2020) was found as 18.991 ( $^{\circ}\text{C}$ ). This finding showed a 1.78 ( $^{\circ}\text{C}$ ) decline in the time series trend. This result indicated that the  $T_{max}$  during January is on decreasing trend. This finding agreed with the MK test results, which showed a decreasing trend for  $T_{max}$  in January. Pettitt was then applied to monthly  $T_{min}$  values and it was observed that  $T_{min}$  showed a change point in 1999 (Fig. 6) for April. The average

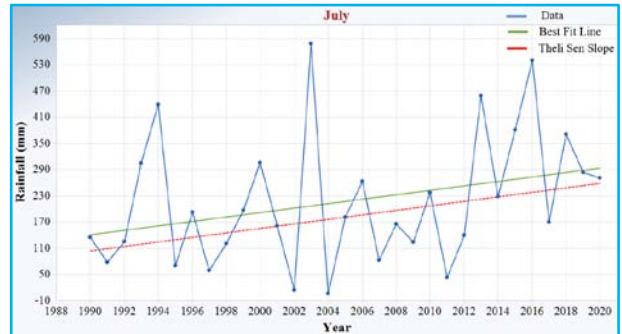
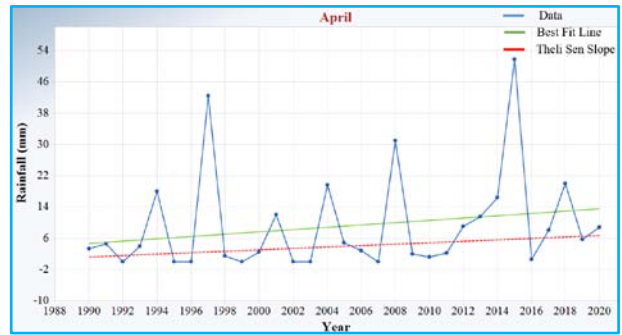
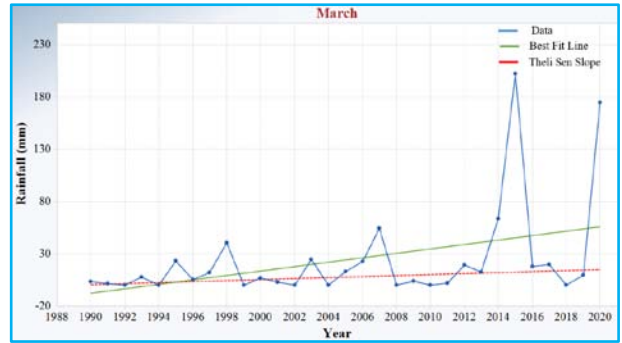


Fig. 8. The significant trend observed in monthly rainfall time series data in (a) March, (b) April and (c) July

$T_{min}$  value before the trend change point (1990-1999) and post-change point (2000-2020) were found as 26.23 ( $^{\circ}\text{C}$ ) and 25.64 ( $^{\circ}\text{C}$ ), respectively. This demonstrates a decline in the  $T_{min}$  of 0.59 ( $^{\circ}\text{C}$ ) in the time series trend. Also,  $T_{min}$  showed a changed point in 2005 for August. The mean value of data series before and after changed points were 26.234 and 25.647  $^{\circ}\text{C}$ , respectively as depicted in Fig. 7.  $T_{max}$  seasonal analysis showed no change point in data series. However,  $T_{min}$  seasonal analysis revealed change point for monsoon and winter seasons. The mean value of  $T_{min}$  data series prior to change point for monsoon season was 25.828  $^{\circ}\text{C}$  which decreased to 25.207  $^{\circ}\text{C}$ , which clearly indicates a declining in the temperature. Similarly, the mean values of  $T_{min}$  data series during the winter season were 7.108 and 6.258  $^{\circ}\text{C}$  prior and post the change point. This further revealed a declining trend of the  $T_{min}$ .

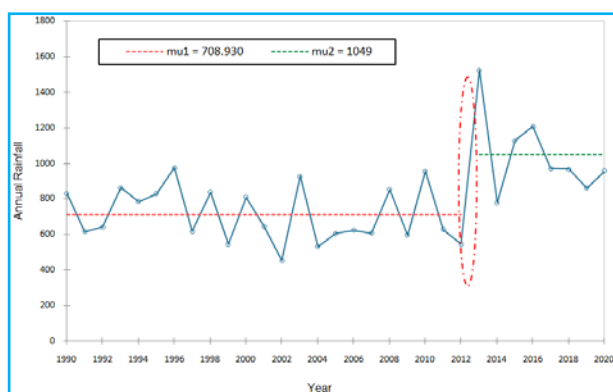


Fig. 9. Change or break point of annual rainfall data series

### 3.3. Rainfall trend analysis

Annual, monthly and seasonal rainfall trends were investigated individually. Annual rainfall showed no statistically significant trend, implying no substantial variation in the chosen station's annual rainfall value. Once the MK test had identified the trends, Theli Sen slope was estimated to quantify the slope magnitude. Monthly and annual MK Test values and Theli Sen slope for rainfall are displayed in Table 5. For the monthly rainfall, a statistically rising trend was found for March and April [Figs. 8 (a&b)]. The MK test value for March was 102 with the Theli Sen slope magnitude of 0.4800. At the same time, April has obtained the MK test value of 102 with a Theli Sen slope magnitude of 0.1833. Only July showed a statistically increasing trend [Fig. 7(c)] during the monsoon period with a MK test value of 107 and Theli Sen slope of 5.1529. Thus, monthly trend results imply increasing rainfall amounts in March, April and July. The seasonal analysis of rainfall showed no trends in the rainfall time series data.

### 3.4. Rainfall Trend Change Identification Analysis

The change point analysis of data series is vital in designs of hydrologic structures. Average annual rainfall time-series data (1990-2020) were subjected to the Pettitt test for change point identification. The shifting point in the annual average rainfall data was observed in 2012 (Fig. 9). The mean value of average annual rainfall before the change point (1990-2012) was 708.93 mm and a mean value of 1049 mm for the annual rainfall data series was found post the change point (2013-2020). The difference in the magnitude of these two data series means values is 340.07 mm. This indicates that there has been rise in the rainfall amount during the last decade temperature for January. Pettitt test was also applied to monthly data and revealed no trend change point in the monthly data series. No change point was reported by the seasonal analysis of the rainfall data.

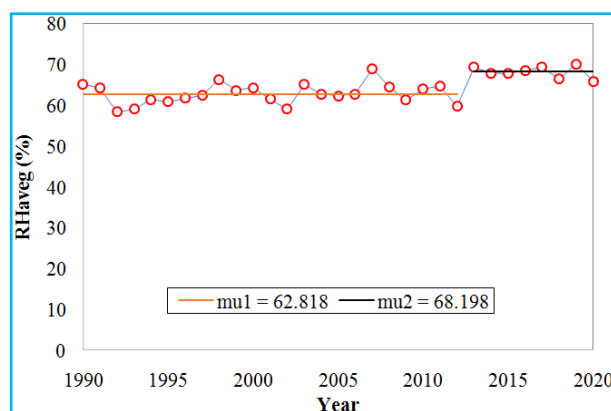


Fig. 10. Change or break point of annual relative humidity data series

### 3.5. Relative Humidity Trend Analysis

The annual average relative humidity (RHavg.) showed a statistically positive increasing trend with M-K test value (S) of 201 and slope magnitude of 0.2546. The monthly trend analysis revealed that relative humidity is on increasing trend during January, March, April, May, October, November and December. However, no trend was seen in the remaining of the months. The seasonal trend analysis showed positive increasing trend during spring, summer, autumn and winter seasons. However, the monsoon season did not reveal a significant trend. The MK Test value and Theli Sen slope for relative humidity are shown in Table 5. The M-K test value was found highest for annual average data series. While, the highest Theli sen slope was observed for summer season. The corresponding lower values for a statistically significant trend was reported for March (105) and December (0.2032). The analysis revealed that relative humidity has been increasing round the year except monsoon season. Though higher relative humidity is required for plant growth but exceeding a critical value may cause the plant growth development by influencing the atmospheric evaporating demand. It is fact that relative humidity is temperature dependent and cannot be regulated artificially to greater extent under open field conditions. However, relative humidity can be regulated using humidifier and dehumidifier under protected cultivation technologies.

### 3.6. Relative Humidity Trend Change Identification Analysis

The change point analysis of annual average relative humidity revealed a shift point in data time series. The change point was found in year 2012 and the average value of relative humidity prior to change point was 62.818% compared to mean value of 68.198% after the change point (Fig. 10). This clearly shows that there has been an increased in the relative humidity post the change

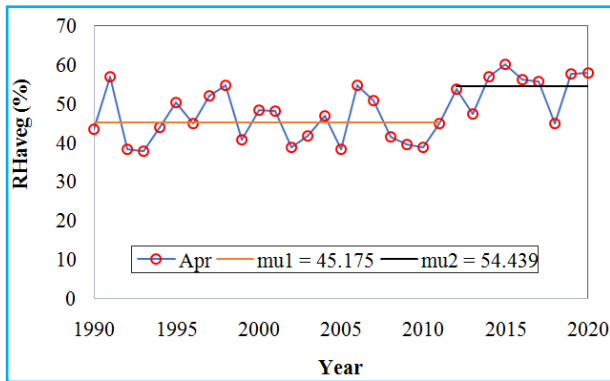


Fig. 11. Change or break point of April relative humidity data series

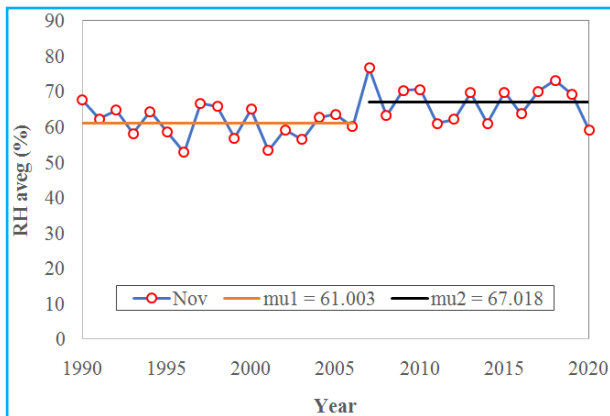


Fig. 12. Change or break point of November relative humidity data series

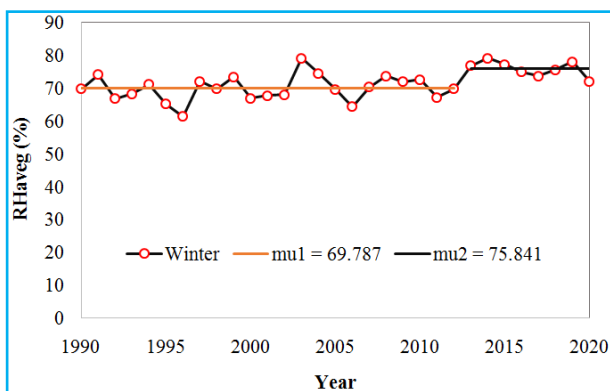


Fig. 13. Change or break point of Winter season relative humidity data series

point by 5.380%. The reason could be in the increased in the relative humidity due to increase in the winter rainfalls events which are associated with the lowering in the temperature. The monthly change point analysis showed change point in the relative humidity trend in April and

November and change occurred in 2011 and 2006, respectively (Figs. 11 and 12). The average relative humidity before the change point was 45.175 % while post change point, average relative found to be 54.439 % for April. For November, the value of relative humidity prior to the change point was 61.003 % which was increased to 67.018 % post the change point. The seasonal analysis showed that there has been break in the trend of relative humidity and the change point was observed in 2012 during winter season (Fig. 13). The mean value of the data time series after the change point was 75.841 % compared to mean value (69.787 %) of data time series prior to trend change point. The remaining months and seasons showed no change or shift point in the time series relative humidity data.

#### 4. Conclusion

Based on historical rainfall temperature records (maximum and minimum) and relative humidity the trend change was examined using three tests (Man Kendall, Theli Sen slope and Pettitt) for Annual, monthly and seasonal data. The outcomes of the tests could not find any specific and significant trends magnitude in the time series data of annual rainfall,  $T_{max}$  and  $T_{min}$  data. But a positive increasing trend was observed in relative humidity data series. Also, disaggregating the data to monthly and seasonal scale,  $T_{max}$  indicated a statistically significant downward trend in January and December, unleashing a mixed effect on crop phenology which has both inhibitory and regulatory impact on growth. But this reduced temperature trend could help water saving by reducing the evapotranspiration rate. Also, a declining trend of  $T_{max}$  during winter season further justified the declining temperature. The  $T_{min}$  showed a statistically significant declining trend in January and December. This dip in the  $T_{min}$  could be detrimental to the *Rabi* season crops and cause chilling injuries. In addition,  $T_{min}$  revealed change declining trend during monsoon, autumn and winter seasons. Notably, the change point for the monthly  $T_{max}$  data series was detected in January 1998. It was observed that  $T_{min}$  in April showed an inflection point in the data series trend in 1999. Further, monsoon and winter showed change points for  $T_{min}$  data. Monthly rainfall indicated a statistically significant increasing trend in March, April and July. This indicated an increase in monsoon and non-monsoon season rainfall magnitude. A typically significant inflection point describing a pattern change in the annual rainfall was detected in 2012. The mean value of average annual rainfall before the change point (1990-2012) was 709 mm and the 1049 mm mean of mean annual rainfall data series was found post-change point (2013-2020). Relative humidity showed positive increasing trends for annual, month and seasonal data series. The implications of such changes to series that

exhibited no trend define multiple methods to identify the trends and avoid biased outcomes deciphered from approaches using a single technique.

#### Acknowledgement

The authors are thankful to the Head, Division of Agricultural Physics, ICAR-Indian Agricultural Research Institute, Pusa Campus, New Delhi, India, for providing the meteorological data.

*Disclaimer* : The contents and views expressed in this study are the views of the authors and do not necessarily reflect the views of the organizations they belong to.

#### References

- Chen, Z. and Grasby, S., E., 2009, "Impact of decadal and century-scale oscillations on hydroclimate trend analyses", *Journal of Hydrology*, **365**, 1, 122-133.
- Das, H. P., Dhotre, A. K. and Rase, D. M., 2008, "Trend variability and trend over Pune", *MAUSAM*, **59**, 3, 291-296.
- Das, J., Mandal, T., Saha, P. and Bhattacharya, S. K., 2020, "Variability and trends of rainfall using non-parametric approaches : A case study of semi-arid area", *MAUSAM*, **71**, 1, 33-44.
- Duhan, D. and Pandey, A., 2013, "Statistical analysis of long term spatial and temporal trends of precipitation during 1901-2002 at Madhya Pradesh, India", *Atmospheric Research*, **122**, 136-149.
- Gallego, M., García, J., Vaquero, J. and Mateos, V., 2006, "Changes in frequency and intensity of daily precipitation over the Iberian Peninsula", *Journal of Geophysics. Res. Atmos.*, **111**, D24105.
- Gebremicael, T., Mohamed, Y., Betrie, G. van der Zaag, P. and Teferi, E., 2013, "Trend analysis of runoff and sediment fluxes in the Upper Blue Nile basin: a combined analysis of statistical tests, physically-based models and landuse maps", *J. Hydrol.*, **482**, 57-68.
- Gellens, D., 2000, "Trend and correlation analysis of k-day extreme precipitation over Belgium", *Theor. Appl. Climatol.*, **66**, 1-2, 117-129.
- Goswami, B. N., Venugopal, V., Sengupta, D., Madhusoodanan, M. S. and Xavier, P. K., 2006, "Increasing trend of extreme rain events over India in a warming environment", *Science*, **314**, 1442-1445.
- Hirsch, R. M., Slack, J. R. and Smith, R. A., 1982, "Techniques of trend analysis for monthly water quality data", *Water Resour. Res.*, **18**, 1, 107-121.
- Jaswal, A. K., 2010, "Recent winter warming over India-spatial and temporal characteristics of monthly maximum and minimum temperature trends for January to March", *MAUSAM*, **61**, 2, 163-174.
- Jones, P. D., New, M., Parker, D. E., Martin, S. and Rigor, I. G., 1999, "Surface air temperature and its changes over the past 150 years", *Rev. Geophys.*, **37**, 2, 173-199.
- Kaur, N., Singh, M., S. Kaur, 2022, "Long term monthly and inter-seasonal weather variability analysis for the lower Shivalik foothills of Punjab", *MAUSAM*, **73**, 1, 173-180.
- Khavse, R. and Chaudhary, J. L., 2022, "Trend assessment in climate variable by Mann Kendall test of Bastar district of Chhattisgarh", *MAUSAM*, **73**, 1, 79-82.
- Kushwaha, N. L., Bhardwaj, A. and Verma, V. K., 2016, "Hydrologic response of Takarla-Ballowal watershed in Shivalik foot-hills based on morphometric analysis using remote sensing and GIS", *J. Indian Water Resour. Soc.*, **36**, 1, 17-25.
- Kushwaha, N. L., Rajput, J., Shirsath, P. B., Sena, D. R. and Mani, I., 2022, "Seasonal climate forecasts (SCFs) based risk management strategies : A case study of rainfed rice cultivation in India", *J. Agrometeorology*, **24**, 1, 10-17.
- Korade, M. S. and Dhorde, A. G., 2016, "Trends in surface temperature over Mumbai and Ratnagiri cities of coastal Maharashtra, India", *MAUSAM*, **67**, 2, 455-462.
- Lal, M., 2001, "Climatic change implications for India's water resources", *J. Ind. Water Resources Soc.*, **21**, 101-119.
- Li, Y., Qin, X., Liu, Y., Jin, Z., Liu, J., Wang, L., and Chen, J., 2022, "Evaluation of long-term and high-resolution gridded precipitation and temperature products in the Qilian mountains, Qinghai-Tibet plateau, *Frontiers in Environmental Science*, **10**, 906821.
- Mohapatra, M. and Mohanty, U. C., 2006, "Spatio-temporal variability of summer monsoon rainfall over Orissa in relation to low-pressure systems", *Journal of Earth System Science*, **115**, 2, 203-218.
- New, M., Todd, M., Hulme, M. and Jones, P., 2001, "Precipitation measurements and trends in the twentieth century", *Int. J. Climatol.*, **21**, 15, 1889-1922.
- Pant, G. B. and Kumar, K. R., 1997, "Climates of South Asia", John Wiley, Chichester, UK.
- Pattanaik, D. R., 2005, "Variability of oceanic and atmospheric conditions during active and inactive periods of storm over the Indian Region", *International Journal of Climatology*, **25**, 1523-1530.
- Pettitt, A. N., 1979, "A non-parametric approach to the change-point detection", *Journal of Applied Statistical*, **28**, 2, 126-135.
- Pingale, S. M., Khare, D., Jat, M. K. and Adamowski, J., 2014, "Spatial and temporal trends of mean and extreme rainfall and temperature for the 33 urban centres of the arid and semi-arid state of Rajasthan, India", *Atmospheric Research*, **138**, 73-90.
- Phukan, R. and Saha, D., 2022, "Analysis of rainfall trends over Tripura", *MAUSAM*, **73**, 1, 27-36.
- Rajani, N. V., Tiwari, M. K., Chinchorkar, S. S., Pampaniya, N. K., 2020, "Long-term trend analysis of rainfall using hybrid Discrete Wavelet Transform (DWT) based Man-Kendall tests in central Gujarat region, India", *MAUSAM*, **71**, 2, 209-224.
- Rupa Kumar, K., Krishan, K. and Pant, G. B., 1994, "Diurnal asymmetry of surface temperature trends over India", *Geophys. Res. Lett.*, **21**, 677-680.
- Shadmani, M., Marofi, S. and Roknian, M., 2012, "Trend analysis in reference evapotranspiration using Mann-Kendall and Spearman's Rho tests in arid regions of Iran", *Water Res. Manage.*, **26**, 1, 211-224.
- Sharma, G. K. and Chaudhary, J. L., 2014, "Time trends in temperature of Bastar plateau agro-climatic zone of Chhattisgarh", *MAUSAM*, **65**, 1, 29-36.

Srivastava, H. N., Dewan, B. N., Dikshit, S. K., Rao, P. G. S., Singh, S. S. and Rao, K. R., 1992, "Decadal trends in climate over India", *MAUSAM*, **43**, 1, 7-20.

Talaei, P. H., Sabziparvar A. A. and Tabari, H., 2012, "Observed changes in relative humidity and dew point temperature in coastal regions of Iran", *Theor. Appl. Climatol.*, **110**, 385-393. doi:10.1007/s00704-012-0630-1.

Wang, K., Dickinson, R. E., 2012, "A review of global terrestrial evapotranspiration: observation, modeling, climatology and climatic variability", *Rev. Geophys.*, **50**.

Zhang, S., Lu, X. X., Higgitt, D. L., Chen, C. T. A., Han, J. and Sun, H., 2008, "Recent changes of water discharge and sediment load in the Zhujiang (Pearl River) Basin, China", *Global Planet. Change*, **60**, 3, 365-380.

