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Measurement of the two components of air earth conduction current in the tropics

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ABSTRACT. An attempt has been made for the first time at Poona (India) to measure both components of air earth current corresponding to the positive and negative ionic conductivities. To avoid electrode effect, at a distance of about 15 m above the ground, two duraluminium test plates (1 m square approximately) are being used one above the other separated by 20 cm. To measure the two components of air earth current, the top and bottom plates are connected to the two pairs of quadrants respectively of a quadrant electrometer and the needle is connected to a steady battery of 22.5 volts. To avoid displacement currents due to potential gradient changes two radioactive collectors are used one for each plate so that the correction due to the displacement current is rectified at the plates themselves. Full details of measurement of the two components of air earth current using the above apparatus are given and their importance in the study of the various problems in atmospheric electricity has also been indicated.

Measurements at Poona during the years 1967, 1968 of the potential gradient and conduction current during widespread fog, mist showed considerable increase in potential gradient and decrease in conduction current often changing from positive to negative values. Thus, the atmospheric electrical elements undergoing specific variations about 1 to 2 hours before the onset of fog and about $\frac{1}{2}$ to $1\frac{1}{2}$ hours before dissipation will be of special importance for meteorological forecasters, in agreement with the results of Dolezalek, Israel and Kasemir.

1. Introduction

The origin of air earth current and the factor which control its variation constitute one of the most important problems of atmospheric electricity. Continuous direct measurements of the air earth current are meagre especially from the tropics. The most successful continuous direct records of the air earth current were that of Simpson (1910) at Simla. The disadvantage of Simpson's method was that the record gave the charge brought down by the air earth current *plus* any induced charge produced by changes in the earth's field, which may have occurred during the period of measurement. It was necessary therefore to allow for field changes by applying corrections obtained from a simultaneous potential gradient record. An improvement of Simpson's method was made by Scrase (1933) using a compensating

condenser which would automatically compensate the field changes. In practice it is found difficult to compensate the field changes using Scrase's method as the slightest variation in the adjustment of the condenser alters the compensation for field changes. It can perhaps be argued as Chalmers (1957) had expressed that Scrase's method is not truly a direct measurement of air earth current but only apparently so, and that what is really measured is the unipolar conductivity of the air between the wire net and the plate, by measuring the current for a given applied potential gradient to the region. The effect of connecting the wire net to the collector, rather than of having it at some other potential, is to make the conditions between the net and plate as close as possible to the natural conditions in the atmosphere. The current measured is not quite the natural current to an area of the earth,

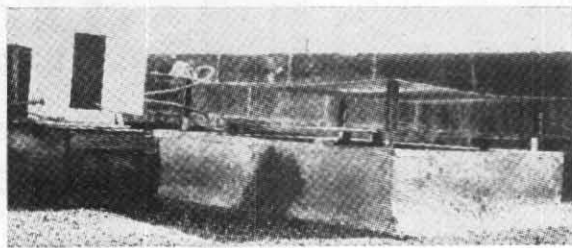


Fig. 1

Position of two duraluminium plates about a metre square one above the other supported by plexiglass insulators for measurement of the two components of air earth conduction current. The top plate receives positive ions coming from above and bottom plate receives negative ions coming from below (earth)

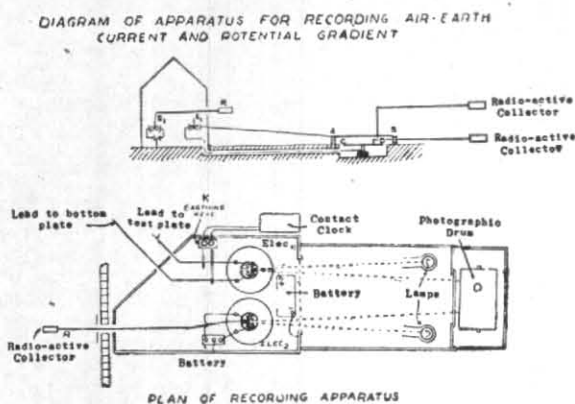


Fig. 2

Scrase's apparatus (1933) with modification for the measurement of two components of air earth current by using two duraluminium plates (AB, CD, one metre square) one above the other using plexiglass insulators. Compensation for the displacement current due to field changes has been arranged by fixing two radio-active collectors one for each plate

when the net is not at the exact potential of the air in its neighbourhood. The improvement made in the present study for the continuous measurement of air earth current is by using two duraluminium plates (Fig. 1), about a metre square, one above the other, separated by a very small distance of 20 cm, using plexiglass insulators. The whole apparatus is at the potential of this level. The exact details of the apparatus and the method of calculation of air earth current are given below.

2. Description of the apparatus

Fig. 2 shows the plan of the recording apparatus for recording air earth current and potential gradient. AB, CD are two duraluminium plates about a metre square and 0.2 cm thick supported by plexiglass insulators in a shallow concrete pit. The bottom plate is arranged to be above the ground level. The air gap between the plate and ground is about 3 cm across while the depth of the

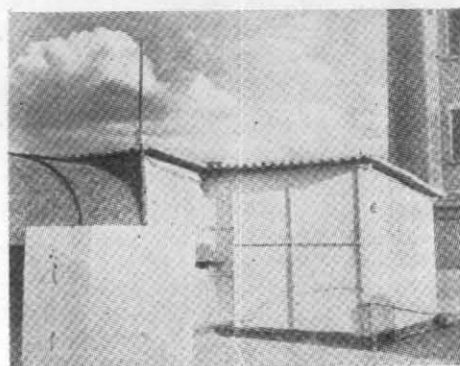


Fig. 3

A view of newly constructed Faraday Cage Hut wherein sensitive electrometers, sensitive galvanometers, photographic recorders for recording potential gradient, air earth current, rain charge, point discharge current from an artificial metal point, point discharge current from tree are housed

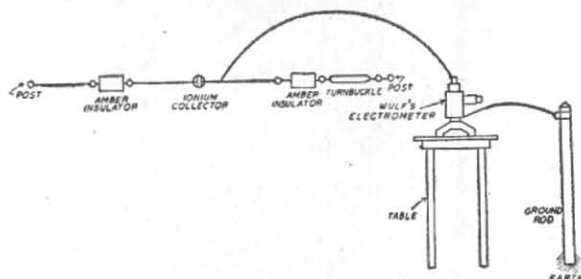


Fig. 4. Diagram showing arrangement for absolute potential gradient measurement

pit is 15 cm. The bottom plate CD is connected by coaxial cable to one pair of quadrants of Dolezalek electrometer E_1 . The top plate AB is connected by coaxial cable to the second pair of quadrants of this electrometer E_1 . The needle of another quadrant electrometer E_2 is connected to a radioactive collector R fixed to an insulated rod projecting through a hole in the recording hut, a Faraday cage (Fig. 3), and the needle of the first quadrant electrometer E_1 is connected to a steady battery of 22.5 volts.

By means of electromagnetic keys K worked by a contact clock both pairs of quadrants of the electrometer E_1 are earthed for one minute at regular intervals of ten minutes. To avoid displacement currents due to potential gradient changes, two radioactive collectors are used one for each plate so that the correction due to the displacement current is rectified at the plates themselves. The whole set up has been fixed on the third terrace of the observatory building (Fig. 2) and as such the measurements of air earth current and potential gradient are made far away from the earth's surface as suggested by Israel (1963).

As shown in the plan of recording apparatus, two lamps are used, one for the measurement of potential gradient and another for the measurement of air earth current. The spots of light coming from the two lamps are reflected by the mirrors of the two electrometers E_1 and E_2 and allowed to fall on the same photographic paper so that we get the continuous record of potential gradient and air earth current on the same photographic paper. Fig. 4 shows the arrangement for absolute potential gradient measurement.

3. Calibration and reduction of records

Air earth current

The magnitude of the air earth current is obtained by the formula—

$$i = \frac{C \cdot V}{A t}$$

where, i is the current per unit area, C is the capacity of the test plate system, A is the area of the plate and V is the change of potential recorded in t seconds.

If k is the scale value of the electrometer in volts per division and d is the deflection in cm recorded in t seconds, the calculation of air earth current in practical units is given by—

$$i = \frac{C k}{A t} \frac{d}{9 \times 10^{11}} \text{ (amp per sq.cm)} \quad (1)$$

The capacity between the two plates is measured accurately by means of a standard bridge. The value of C obtained is 544 cm. The scale value of the electrometer E_1 is determined by applying a known voltage (1.45 volts) after both pairs of quadrants have been earthed momentarily and then insulated. The sensitivity is such that one cm in the record (at 100 cm from the electrometer) is equivalent to 1.45 volts on the test plate. The area A of the plates being 11881 cm² and the time interval t between successive earth marks being 9 min (540 sec), the calibration factor given by formula (1) is obtained as follow—

$$i = \frac{544 \times 1.45 \times d}{109 \times 100 \times 450 \times 9 \times 10^{11}}$$

$$= 136.5 \times 10^{-18} \text{ amp/sq cm}$$

where, d is the deflection in cm reached in 540 sec.

The scale value k is measured once a month and the factor is altered if necessary by applying proper voltage to the needle of the electrometer E_1 ; generally 22.5 volts is applied to the needle of the electrometer E_1 .

Analysis of records

In consequence of the vertical electric field and of the conductivity of the air, an electric current flows downward in fine weather through the lower atmosphere into the earth, the positive and negative ions, moving downwards and upwards respectively. The magnitude of the conduction current in the free air is given by—

$$i = F (\lambda_+ + \lambda_-)$$

where, i = conduction current intensity,

F = intensity of electric field

λ_+ = positive conductivity,

λ_- = negative conductivity.

If there is a vertical gradient of space charge density, turbulent transfer will cause a vertical convection current and indeed careful consideration is necessary in attempting to derive total vertical current into the ground from measurements of the field and the polar conductivity of the air. By the method described already it is possible, however, to measure the current directly.

The conductivity of the air increases fairly rapidly with increasing height above the ground, for two reasons; ionic mobility is greater at lower air densities and the number of ions present is also greater owing to the enhanced effect of the down-coming cosmic radiation. A vertical current independent of height, is thus maintained through a considerable vertical thickness despite the decrease in the vertical field at higher levels.

The circulation of electricity in the atmosphere above any place may be expressed by the simple equation,

$$i = \frac{V}{R}$$

where, i is the air earth current, V is the potential difference between the ionosphere and the ground, R is the total effective resistance of the air column above the place. The potential gradient F at the surface is the product of the current and the specific resistance of the air r (i.e., the reciprocal of the positive conductivity at the surface). Thus we may write—

$$F = ir = V (r/R)$$

Whipple (1932) has shown how the current flowing from the air into the earth is better represented by the product of the potential gradient and the positive conductivity.¹

During the Third International Conference on Atmospheric and Space Electricity held at Montreal in 1963, Prof. H. Israel had emphasized the importance of measuring simultaneously the atmospheric electric parameters, viz., potential

TABLE 1
Hourly means of potential gradient (volts/m) at Poona during Nov 1965-Dec 1966

Time (IST)	1965		1966											
	Nov	Dec	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec
0-1	182	212	152	134	112	84	51	36	31	33	74	108	146	169
1-2	160	190	135	105	104	75	48	35	32	36	74	103	123	136
2-3	128	152	144	96	104	75	43	26	29	36	67	98	107	115
3-4	94	146	97	82	96	69	44	29	23	31	67	90	92	107
4-5	104	139	98	82	93	79	43	35	26	30	61	89	100	100
5-6	119	143	108	93	101	98	51	40	30	45	67	101	123	123
6-7	154	209	143	134	138	127	63	37	41	38	96	128	161	176
7-8	221	225	175	209	196	175	75	38	40	45	120	180	230	222
8-9	274	291	283	272	217	188	70	50	45	44	138	231	238	308
9-10	333	400	303	270	212	146	61	46	35	40	116	221	222	314
10-11	265	337	251	236	155	135	51	44	37	38	115	188	184	262
11-12	188	238	191	174	121	97	36	36	31	36	109	147	176	192
12-13	147	209	153	120	96	83	31	32	37	32	89	116	146	138
13-14	121	198	104	89	71	70	31	29	37	37	70	109	138	107
14-15	127	172	92	69	65	48	30	32	30	33	88	100	138	123
15-16	134	172	81	68	61	22	26	31	25	28	77	101	130	115
16-17	122	182	75	71	55	31	36	33	24	21	76	104	146	123
17-18	181	241	112	96	46	60	47	20	28	15	68	111	199	184
18-19	226	275	151	143	67	99	67	29	32	23	83	123	215	260
19-20	249	304	185	159	107	117	76	22	35	26	82	137	207	285
20-21	242	267	195	179	93	111	73	44	31	29	76	150	184	253
21-22	227	289	190	166	93	96	68	45	33	37	83	176	176	238
22-23	206	314	182	157	104	77	60	35	30	32	77	152	146	207
23-24	169	285	179	138	108	90	51	36	28	32	77	131	146	199

TABLE 2
Seasonal and annual means of potential gradient (F), conductivity (λ) and air earth current (i) during January to December 1966

Time (GMT)	Potential gradient (volts/m)				Conductivity (λ) ($10^{-14} \Omega^{-1}m^{-1}$)				Air earth current (i) ($\times 10^{-12} \text{ amp}/m^2/\text{sec}$)			
	Summer	Equinox	Winter	Year	Summer	Equinox	Winter	Year	Summer	Equinox	Winter	Year
0030	41	92	112	82	—	1.04	1.89	1.73	—	0.956	2.116	1.420
0130	45	122	153	107	—	0.78	1.20	1.23	—	0.956	1.843	1.311
0230	49	168	209	142	1.81	0.62	0.88	0.87	0.887	1.047	1.843	1.230
0330	52	193	275	174	2.10	0.83	0.89	0.97	1.092	1.604	2.457	1.689
0430	45	174	277	165	2.58	0.84	1.01	1.04	1.160	1.467	2.739	1.723
0530	43	148	233	141	2.54	1.08	1.23	1.27	1.092	1.604	2.867	1.792
0630	35	120	183	113	2.75	1.51	1.49	1.62	0.955	1.809	2.730	1.826
0730	33	96	139	89	2.48	1.71	2.16	1.99	0.82	1.64	3.00	1.77
0830	33	80	109	74	2.48	2.01	2.75	2.37	0.82	1.60	3.00	1.76
0930	31	75	105	71	2.22	2.37	2.53	2.43	0.69	1.77	2.66	1.72
1030	27	65	99	64	2.78	1.94	1.90	2.40	0.75	1.26	2.87	1.54
1130	29	66	104	66	2.82	1.86	2.49	2.22	0.82	1.23	2.59	1.47
1230	27	71	148	82	2.27	1.78	1.66	1.70	0.61	1.26	2.46	1.40
1330	38	93	192	108	1.98	1.25	1.32	1.29	0.75	1.16	2.53	1.40
1430	40	111	209	120	1.88	1.05	1.17	1.15	0.75	1.16	2.46	1.38
1530	44	107	203	118	1.40	1.05	1.11	1.08	0.61	1.13	2.25	1.28
1630	46	112	193	117	1.48	1.01	1.17	1.11	0.68	1.13	2.25	1.30
1730	39	103	173	105	1.40	1.13	1.26	1.20	0.55	1.16	2.18	1.26
1830	37	101	165	101	2.77	1.35	1.65	1.60	1.02	1.37	2.73	1.62
1930	38	95	150	90	2.15	1.40	1.87	1.74	0.82	1.33	2.80	1.57
2030	38	89	125	84	1.79	1.46	2.24	1.81	0.68	1.30	2.80	1.52
2130	33	86	115	78	1.24	1.48	2.14	1.75	0.41	1.27	2.46	1.37
2230	32	81	95	69	3.41	1.40	2.59	1.95	1.09	1.14	2.46	1.35
2330	33	81	95	70	4.14	1.35	2.59	2.27	1.37	1.09	2.46	1.59

TABLE 3

Maximum and minimum times of atmospheric parameters (F , λ , i) for the various seasons

	Max time (GMT)						Min Time (GMT)					
	Potential gradient		Conductivity		Air earth current		Potential gradient		Conductivity		Air earth current	
Summer	0330	1630	1130	2330	0430	1830	1030	2230	0230	1530	0930	2130
Equinox	0330	1630	0930	2130	0630	1830	1030	2330	0230	1630	0030	1530
Winter	0430	1430	1030	2230	0730	2030	1030	2230	0230	1530	0230	1530
Year	0330	1430	0930	2330	0630	1830	1030	2230	0230	1530	0230	1530

gradient, conductivity of air and the air earth current, not at the ground but at an altitude of at least two metres to avoid the electrode effect. During the discussion between Prof. Israel, Dr. Chalmers and Dr. Kasemir, on the measurement of air earth current, Dr. Kasemir said that the air earth current at a height above the earth can be measured by using two plates one above the other separated by a very small distance. The two output signals can be fed into a balance amplifier, the output of which yields the current. Dr. Kasemir had mentioned that the technical problem concerning the measurement of air earth current are quite large, but theoretically it should be possible to solve them.

A considerable number of simultaneous records of potential gradient and air earth current were obtained during 1966 and 1967. During 1966, records of air earth current and potential gradient for undisturbed conditions, hourly means centering at the exact hours were obtained. Mean values of the conductivity are obtained by dividing the monthly means of the hourly values of the current by the corresponding means of the potential gradient. Table 1 gives the mean monthly hourly values of potential gradient of 1966. Table 2 (a, b and c) gives the mean hourly values of potential gradient, conductivity and air earth current for the various seasons and year as a whole. Table 3 gives the times of maximum and minimum values of potential gradient, conductivity and air earth current. Maximum values are given in the first three columns and minimum values in the next three columns. The seasons are each composed of four months, *viz.*, *winter* including November to February, *equinox* including March, April, September and October and *summer* including May to August. Curves showing the seasonal and annual variations of the three elements are given in Fig. 6.

4. Comparison of the seasonal and annual variations of potential gradient, conductivity and air earth current of 1966

The potential gradient and air earth current curves follow the usual form with a maximum about mid-winter and minimum about mid-summer. The maximum of conductivity occurs in summer, at about 2330 GMT. The variation of conductivity is extremely large. It is seen that the daily variation of conductivity is very closely related to that of the potential gradient but the changes are in the opposite direction. The conductivity curves are in fact almost exact mirror images of the potential gradient curves. The characteristic changes of these two elements are closely associated with atmospheric pollution showing a double oscillation. The connection between pollution and potential gradient has been studied by Chree and Watson (1924) and by Whipple (1929) who attributes to Simpson an explanation of the double oscillation of pollution as resulting from a combination of two effects—the variation in the rate of production of pollution and the variation in the stability of the atmosphere. In the winter, the air earth current more or less follows the potential gradient variation, closely, there being a very prominent maximum in air earth current at about 0730 GMT, while the potential gradient maximum is at about 0430 GMT and fairly strong minimum at 0230 GMT for air earth current at 2230 GMT for potential gradient. In summer the maximum occurs in the air earth current at 0430 GMT and minimum at 2130 GMT, while the maximum for potential gradient occurs at 0330 GMT and minimum occurs at 1030 GMT. In the equinoctial and annual curves of the air earth current, the changes are small and somewhat irregular. This is presumably because the opposing effects of summer and winter almost equal to each other. Typical fair weather records of air earth current and potential gradient for the

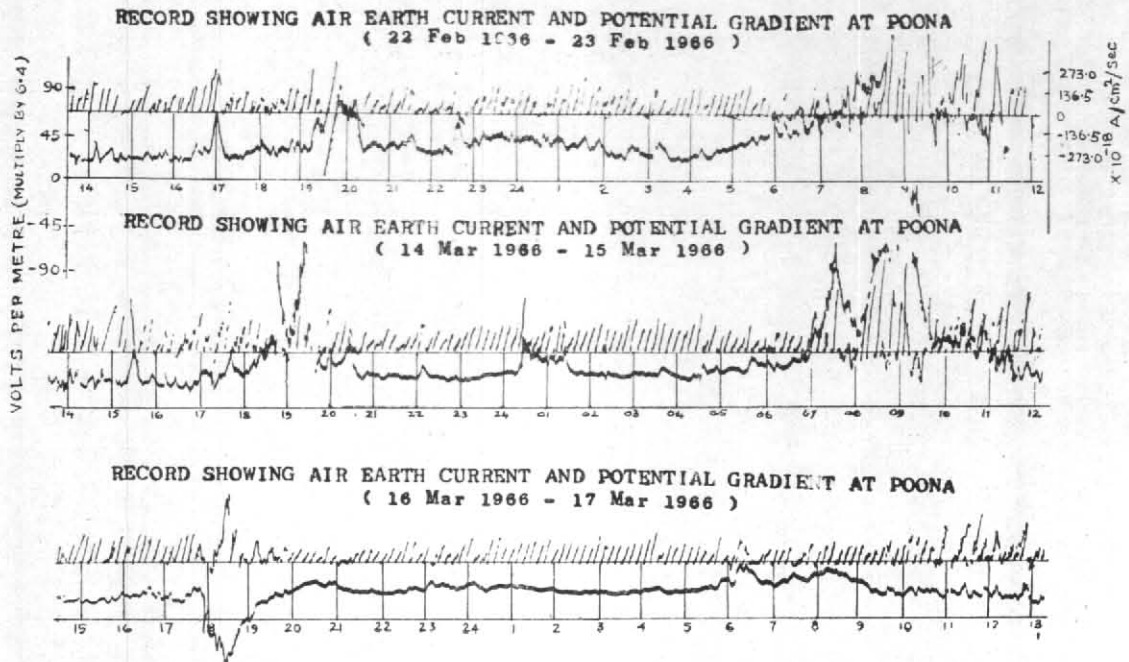


Fig. 5

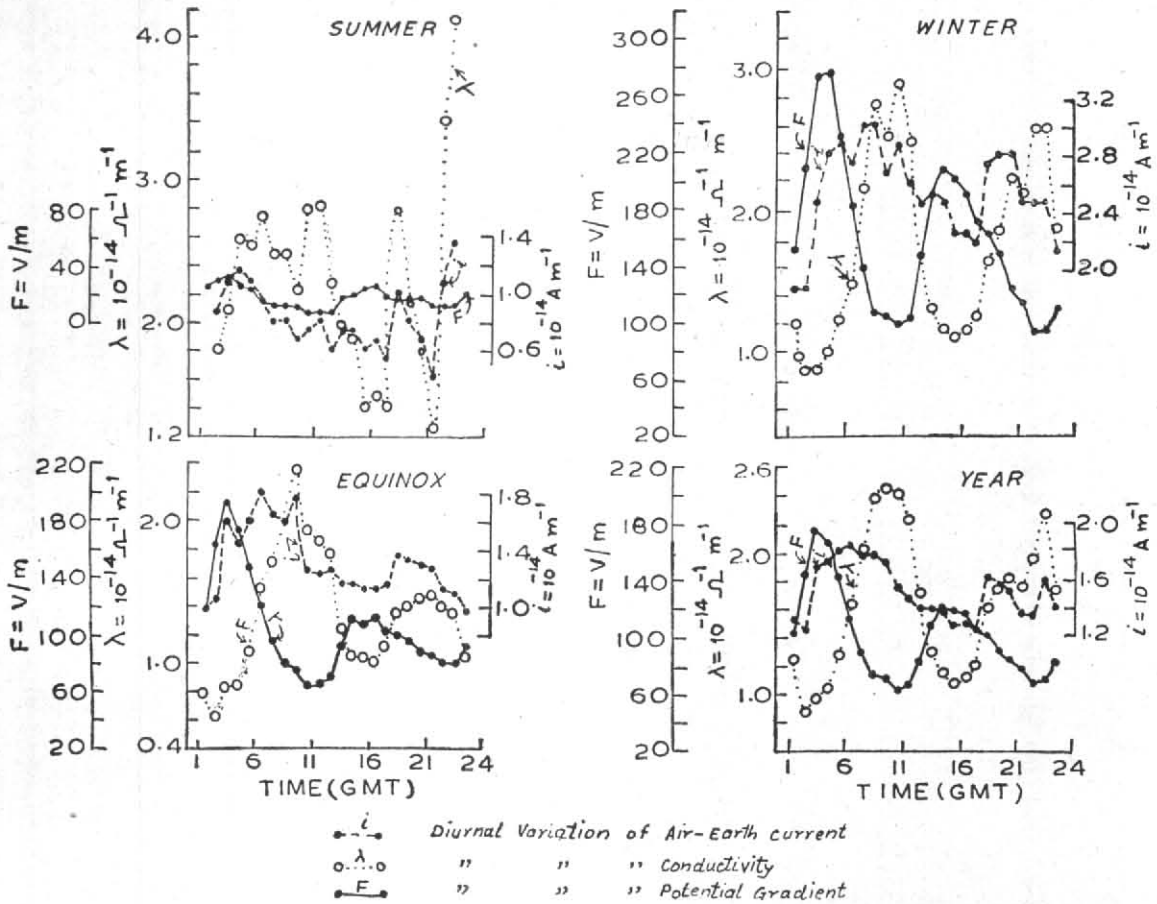


Fig. 6. Diurnal variation of the three parameters for respective seasons of 1966 at Poona

TABLE 4

Seasonal and annual means of potential gradient (F), conductivity (λ) and air earth current (i) at Poona during January to December 1967

Time (GMT)	Potential gradient (F) (volts/m)				Conductivity (λ) ($10^{-14} \Omega^{-1} m^{-1}$)				Air earth current (i) ($\times 10^{-12}$ amp/m ² /sec)			
	Summer	Equinox	Winter	Year	Summer	Equinox	Winter	Year	Summer	Equinox	Winter	Year
0030	55	83	112	82	-0.5	0.5	-4.0	-1.0	-0.3	0.4	-4.5	-0.8
0130	63	132	171	119	-0.8	0.2	-4.0	-1.3	-0.5	0.3	-6.8	-1.5
0230	75	150	222	143	-1.3	0.2	-3.9	-1.4	-1.0	0.3	-8.6	-2.0
0330	75	176	325	180	-0.4	0.2	-2.5	-0.8	-0.3	0.4	-8.1	-1.5
0430	68	176	323	176	0.1	0.6	-0.5	0.3	0.1	1.1	-0.7	0.4
0530	56	142	285	150	-0.5	1.3	-1.7	0.1	-0.3	1.9	-4.8	0.1
0630	58	128	188	119	0.9	1.6	-3.6	-0.3	0.5	2.0	-6.8	-0.3
0730	55	98	131	91	2.0	1.5	-4.7	-0.4	1.1	1.5	-6.1	-0.4
0830	45	76	98	70	1.8	3.4	-3.6	1.0	0.8	2.6	-3.5	0.7
0930	52	67	97	70	0.6	3.7	-1.5	1.1	0.3	2.5	-1.5	0.8
1030	43	66	88	65	0	3.0	-1.7	0.8	0	2.0	-1.5	0.5
1130	38	69	98	65	-1.3	3.3	-2.0	0.5	-0.5	2.3	-2.0	0.3
1230	45	82	146	85	-2.2	1.8	-1.8	-0.4	-1.0	1.5	-2.6	-0.3
1330	54	100	175	104	-4.6	0.8	-2.3	-1.4	-2.5	0.8	-4.0	-1.4
1430	54	105	221	118	-6.1	0.1	-2.3	-1.9	-3.3	0.1	-5.1	-2.3
1530	56	92	191	106	-3.6	0.4	-2.4	-1.4	-2.0	0.4	-4.5	-1.5
1630	55	92	174	101	-2.9	0.3	-3.1	-1.6	-1.6	0.3	-5.5	-1.6
1730	52	83	169	96	-2.7	0.5	-2.9	-1.2	-1.4	0.4	-4.9	-1.5
1830	47	85	167	94	-2.1	0.6	-4.2	-1.6	-1.0	0.5	-7.0	-1.5
1930	50	81	152	89	-0.8	1.4	-4.0	-1.2	-0.4	1.1	-6.1	-1.1
2030	47	76	125	79	0.9	1.3	-4.2	-0.6	0.4	1.0	-5.3	-0.5
2130	43	73	94	68	1.6	1.4	-4.0	-0.1	0.7	1.0	-3.8	-0.1
2230	43	67	81	73	1.2	1.8	-3.6	0.1	0.5	1.2	-2.9	0.1
2330	50	70	84	66	1.0	1.4	-3.9	0	0.5	1.0	-3.3	0

months March and February 1966 are reproduced in Fig. 5.

5. Measurement of the two components of air earth current during 1967

During the measurement of air earth current in 1966, the two test plates are brought to earth potential every nine minutes. This means that each plate will receive positive charge when the potential gradient is positive. But, according to Chalmers (private communication) the plate should be brought to the potential of their surroundings each time the record is to start. As the radioactive collector is in the same equipotential as the plate, two radioactive collectors, one for each plate, were used in 1967. With these two collectors, the apparatus is corrected for displacement current, i.e., change of potential gradient.

Chalmers (1957) has mentioned that one of the main problems in atmospheric electricity to be solved is the measurement of the two components of air earth current corresponding to the positive and negative conductivities. So the equation connecting the two components of air earth current, potential gradient and conductivity can be written as—

$$i_+ + i_- = F(\lambda_+ + \lambda_-)$$

$$\lambda_+ = n_+ .e. k_+$$

$$\lambda_- = n_- .e. k_-$$

where, n_+ = No. of positive ions

n_- = No. of negative ions

e = Charge of ion

k_+ = Mobility of positive ions

k_- = Mobility of negative ions

TABLE 5

Hourly values recorded at Poona, of potential gradient (F), two components of air earth current (i_+ , i_-), two components of conductivity (λ_+ , λ_-) and number of positive and negative ions expressed as (n_+ , n_-)
 F in V/m, (i_+ , i_-) in 10^{-12} Am $^{-2}$, in 10^{-14} Ω^{-1} m $^{-1}$ = (λ_+ , λ_-)

Time (IST)	Quantity				Quantity				Quantity			
	F	i_+ , i_-	λ_+ , λ_-	n_+ , n_-	F	i_+ , i_-	λ_+ , λ_-	n_+ , n_-	F	i_+ , i_-	λ_+ , λ_-	n_+ , n_-
	2 to 3 Jan 1967				3 to 4 Jan 1967				21 to 22 Jan 1967			
14-15	61	3.00	4.92	2050	—	—	—	—	61	2.05	3.36	1400
15-16	69	1.90	2.76	1150	84	2.74	3.26	1358	38	1.64	4.32	1800
16-17	69	—	—	—	77	2.74	3.56	1484	46	1.50	3.26	1358
17-18	84	3.00	3.58	1491	69	2.74	3.98	1659	0	1.37	—	—
18-19	123	3.82	3.10	1292	69	3.54	5.14	2142	69	1.64	2.38	992
19-20	115	3.82	3.32	1383	100	3.54	3.54	2142	61	1.77	2.90	1299
20-21	107	3.00	2.80	1167	169	5.18	3.06	1275	46	1.77	3.85	1604
21-22	115	3.54	3.08	1283	184	3.28	1.78	742	0	-0.41	—	—
22-23	146	3.54	2.42	1008	169	7.38	4.36	1817	-38	-3.00	7.89	-3288
23-24	13	3.54	2.72	1133	115	1.36	1.18	492	-8	-0.82	-10.25	-4271
24- 1	84	1.94	2.26	942	115	2.46	2.14	892	-8	0.14	1.75	729
1-2	84	2.18	2.60	1083	130	2.46	1.90	792	-8	-0.27	-3.37	-1404
2-3	84	-2.46	-2.92	-1217	92	1.36	1.48	617	0	-0.27	—	—
3-4	84	-11.74	-13.98	-5822	77	0.82	1.06	442	15	-0.14	0.64	-267
4-5	77	-21.02	-27.30	-11380	69	-1.10	-1.60	-667	23	-1.23	-5.35	-2229
5-6	84	-26.48	-31.52	-13130	84	-2.18	-2.60	-1083	23	-2.59	-11.26	-4691
6-7	115	-28.66	-24.92	-10390	115	-2.18	-1.90	-792	69	-7.92	-11.48	-4784
7-8	153	-29.76	-19.46	-8108	138	3.28	2.38	992	23	-0.82	-3.57	-1488
8-9	146	-14.46	-9.90	-4125	199	9.56	4.20	2000	107	0.82	0.77	321
9-10	123	4.64	3.78	1575	238	4.10	1.72	717	161	1.37	0.85	354
10-11	161	2.46	1.42	592	192	1.36	0.70	292	169	0.27	0.16	667
11-12	130	3.00	2.30	958	207	—	—	—	169	-1.09	-0.64	-267
12-13	123	2.18	1.78	742	146	2.18	1.50	625	130	1.09	0.84	350
13-14	100	1.64	1.64	683	—	—	—	—	84	1.64	1.95	812
	31 Mar to 1 Apr 1967				3 May to 4 May 1967				21 May to 22 May 1967			
14-15	69	3.00	4.35	1812	23	3.82	16.61	6921	46	2.59	5.63	2346
15-16	54	3.14	5.81	2421	38	3.69	9.71	4046	—	2.59	5.63	2346
16-17	100	3.14	3.14	1307	46	4.37	9.50	3959	54	2.46	4.55	1895
17-18	100	3.55	3.55	1477	84	-3.69	-4.40	-1833	38	2.46	6.47	2696
18-19	92	3.69	4.01	1670	100	-7.92	-7.92	-3300	46	2.32	5.00	2084
19-20	107	3.28	3.07	1279	115	-8.74	-7.60	-3168	38	2.18	5.74	2392
20-21	107	3.82	3.57	1488	115	-10.92	-9.50	-3959	23	2.18	9.48	3950
21-22	130	3.41	3.62	1091	138	-13.24	-9.60	-4000	15	1.77	11.80	4917
23-24	84	3.14	3.74	1559	115	-9.55	-8.30	-3458	23	1.77	7.7	3206
23-24	100	2.87	2.87	1196	107	-7.51	-7.02	-2925	15	1.37	9.13	3805
24- 1	69	3.00	4.35	1812	107	-9.28	-8.67	-3613	31	1.91	6.16	2566
1-2	54	2.73	5.06	2109	84	-3.28	-3.90	-1625	38	2.05	5.40	2250
2-3	61	3.10	4.92	2050	69	3.14	4.55	1895	31	1.91	6.16	2566
3-4	69	2.87	4.16	1734	54	3.69	6.83	2845	38	1.77	4.66	1942
4-5	54	2.87	5.31	2213	54	2.87	5.31	2213	38	1.77	4.66	1942
5-6	77	3.00	3.90	1625	61	3.14	5.15	2146	54	1.50	2.78	1159
6-7	100	3.28	3.28	1367	54	3.14	5.81	2421	38	2.05	5.04	2250
7-8	115	3.00	2.61	1087	92	-3.90	-3.26	-1358	61	2.32	3.8	1583
8-9	77	3.28	4.26	1775	92	-3.14	-3.41	-1419	54	2.46	4.55	1895
9-10	84	3.14	3.74	1559	69	-0.14	-0.59	-246	54	2.32	4.30	1792
10-11	115	3.14	2.73	1138	38	2.46	6.47	2696	46	2.32	5.00	2084
11-12	77	3.28	4.26	1775	31	3.28	10.60	4417	46	2.59	5.63	2346
12-13	69	3.14	4.55	1895	23	3.55	15.43	6430	46	2.46	5.35	2229
13-14	69	3.10	4.35	1812	0	3.55	—	—	46	2.46	5.35	2229

TABLE 6

Time (IST)	Quantity											
	<i>F</i>	<i>i</i> ₊ , <i>i</i> ₋	<i>λ</i> ₊ , <i>λ</i> ₋	<i>n</i> ₊ , <i>n</i> ₋	<i>F</i>	<i>i</i> ₊ , <i>i</i> ₋	<i>λ</i> ₊ , <i>λ</i> ₋	<i>n</i> ₊ , <i>n</i> ₋	<i>F</i>	<i>i</i> ₊ , <i>i</i> ₋	<i>λ</i> ₊ , <i>λ</i> ₋	<i>n</i> ₊ , <i>n</i> ₋
	5 to 6 April 1967				1 to 2 April 1967				27 to 28 June 1967			
14-15	—	—	—	—	61	3.28	5.38	2242	—	—	—	—
15-16	54	2.73	5.06	2109	54	3.55	6.57	2738	—	—	—	—
16-17	46	2.87	6.24	2600	15	1.09	7.27	3029	—	—	—	—
17-18	115	4.09	3.56	1484	100	3.55	3.55	1479	32	0.68	2.13	889
18-19	161	3.41	2.12	883	138	0.82	0.59	246	45	-2.32	-5.16	-2150
19-20	238	-1.23	-3.04	-1267	161	-2.05	-1.27	-529	38	-3.14	-8.28	-3449
20-21	153	2.87	1.88	783	146	3.69	2.53	1055	-7	-0.27	-3.86	-1608
21-22	146	4.09	2.80	1167	153	2.87	1.87	779	25	-1.37	-5.48	-2283
22-23	130	4.09	3.15	1312	100	3.14	3.15	1307	45	-4.09	-9.11	-3796
23-24	146	2.96	2.71	1129	100	3.41	3.41	1419	32	-2.87	-9.00	-3750
0-1	130	3.28	2.52	1050	92	3.55	3.86	1608	45	-2.59	5.76	-2400
1-2	123	4.09	2.33	1388	77	4.23	5.49	2288	45	-2.32	-5.16	-2150
2-3	100	3.69	3.65	1538	84	3.69	4.39	1829	32	-1.91	-5.94	-2475
3-4	115	4.23	3.68	1533	92	3.69	4.01	1670	20	-0.27	-1.35	-562
4-5	107	3.96	3.70	1542	84	2.87	3.42	1425	32	-0.17	-0.53	-221
5-6	199	3.69	1.85	771	107	3.41	3.19	1329	77	-0.27	-0.35	-146
6-7	245	-9.55	-3.90	-1625	138	2.87	2.08	867	70	-0.27	-0.39	-163
7-8	346	-13.24	-3.83	-1596	176	-1.77	-1.08	-417	103	0.55	0.53	221
8-9	384	-13.24	-3.44	-1433	199	-3.82	-1.92	-800	96	—	—	—
9-10	222	-13.24	-5.96	-2483	245	-1.91	-0.78	-325	77	0.55	0.71	296
10-11	184	-9.01	-5.46	-2275	291	-13.24	-4.55	-1895	33	—	—	—
11-12	146	-11.06	-7.58	-3159	153	1.09	0.71	296	-33	0.68	2.06	858
12-13	84	-1.09	-1.30	-542	84	3.82	4.55	1895	58	0.27	0.47	196
13-14	—	—	—	—	77	3.69	4.79	1995	—	—	—	—
	3 to 4 June 1967				1 to 2 July 1967				2 to 3 July 1967			
14-15	15	2.59	17.27	7201	32	4.5	14.06	5859	—	—	—	—
15-16	15	1.23	8.20	3417	32	3.69	11.53	4804	—	—	—	—
16-17	15	2.46	16.40	6832	25	2.87	11.48	4784	—	—	—	—
17-18	15	3.41	22.73	9471	38	1.64	4.32	1800	70	1.91	2.73	1138
18-19	38	4.23	11.13	4637	45	1.09	2.42	1008	45	0.68	1.51	629
19-20	54	4.37	8.09	3370	38	-3.96	-10.42	-4342	38	-0.55	-1.45	-604
20-21	54	4.37	8.09	3370	13	-0.55	-4.23	-1762	58	0	0	0
21-22	54	4.37	8.09	3370	45	-2.18	-4.84	-2017	58	-0.41	-0.71	-296
22-23	46	4.37	9.50	3959	25	-1.91	-7.64	-3184	51	0.27	0.53	221
23-24	38	4.37	11.50	4792	20	-3.69	-18.45	-7688	38	0.14	0.37	154
0-1	31	3.96	12.77	5321	38	-1.09	-2.87	-1196	58	1.37	2.36	983
1-2	38	3.96	10.42	4342	45	0.68	1.51	629	64	2.05	3.20	1333
2-3	31	3.55	11.45	4772	7	0.55	7.86	3275	64	2.40	3.75	1563
3-4	31	3.55	11.45	4772	-218	2.32	1.06	4417	58	1.77	3.05	1271
4-5	31	3.69	11.90	4958	—	—	—	—	58	-0.14	-0.24	-100
5-6	38	4.09	10.76	4483	-223	0.68	0.30	126	64	2.59	4.05	1688
6-7	46	4.23	9.20	3833	-13	0	0	0	77	3.00	3.89	1621
7-8	46	4.23	9.20	3833	20	0.41	-2.05	-854	77	3.41	4.43	1846
8-9	38	4.09	10.76	4483	25	-0.14	-0.56	-233	58	3.41	5.88	2450
9-10	31	4.09	13.19	5495	13	0.14	1.08	450	64	1.77	2.77	1154
10-11	31	4.09	13.19	5495	45	-1.91	-4.24	-1762	58	—	—	—
11-12	23	4.23	18.39	7663	25	-0.95	-3.80	-1767	77	—	—	—
12-13	23	1.91	8.30	3459	32	-2.18	-6.81	-2837	—	—	—	—
13-14	23	0.41	1.78	742	13	1.09	8.38	3491	—	—	—	—

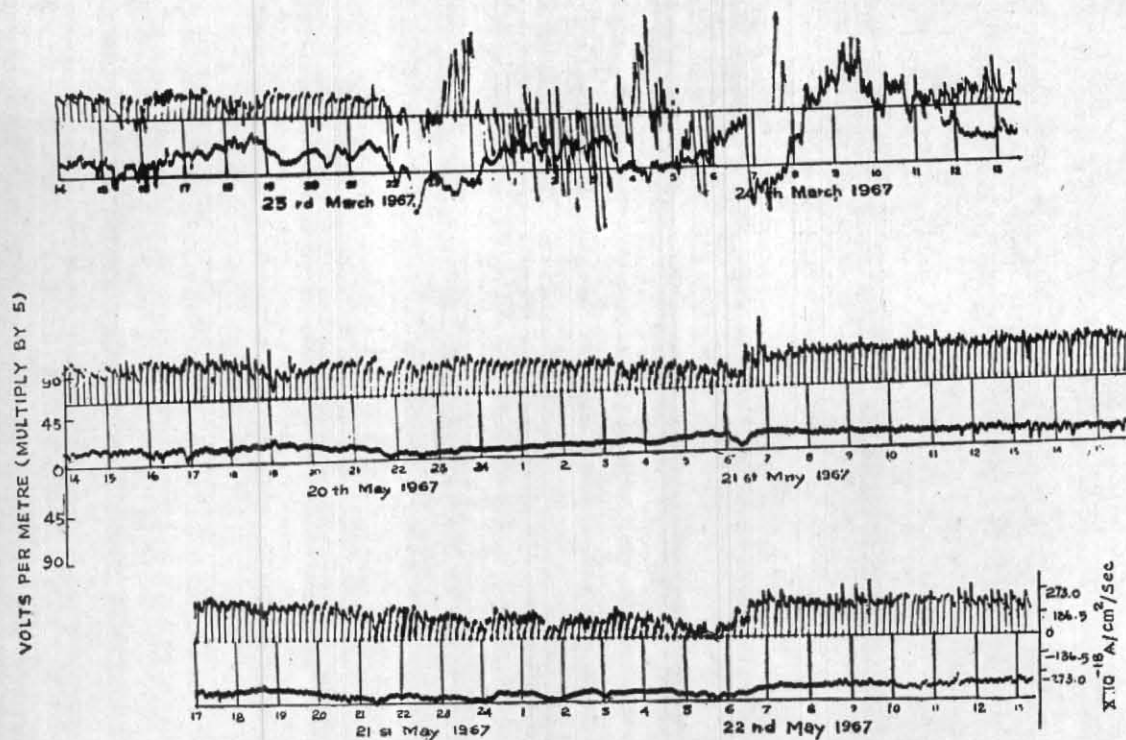


Fig. 7

Record showing air earth current and positive potential gradient at Poona on 23 March (Equinox Time) 20 May, 21 May 1967 (Summer Time). The positive and negative components of air earth current are clearly seen on 23 March 1967

Assuming the values of e , k_+ , k_- and from the measurement of i_+ , i_- and F , it is possible to calculate the number of positive ions and negative ions, which are predominant. Table 4 gives the mean hourly values of potential gradient, conductivity and the air earth current for the various seasons and year as a whole for the year 1967.

Tables 5 to 7 give the values of the two components of air earth current, potential gradient, positive conductivity and negative conductivity, positive and negative number of ions for selected days in winter, equinox and summer (Figs. 8 to 11) show the records of the two components of air earth current and potential gradient. It is interesting to see from Fig. 11 (4-5 September 1967, 9-10 September 1967 and 19-20 September 1967) that the positive air earth current becomes negative between 1800-1900 hrs and again becomes positive between 0800-0900 hrs. This means that after sunset, during night, negative ions predominate for these days. Tables 5 to 7 give the values of positive ions and negative ions also for each hour for the various dates.

It may be mentioned, however, that on some fine weather days we get only positive air earth currents throughout the day (quiet days) and on some disturbed days due to fog, mist, haze, we get

only negative currents throughout the day. Fig. 7 (20 to 21 May 1967, 21 to 22 May 1967) shows only positive currents. Figs. 8 to 11 show both positive and negative currents.

6. Discussion

1. Significance of the air earth current

The general constancy of the air earth current with height above the ground and in different parts of the world, suggests that at any rate to a first approximation, it has its origin in a constant potential difference between the conducting layers of the upper atmosphere and the earth and that its actual value depends upon the resistance of the air between the two regions. It has been found from simultaneous determination of potential gradient and atmospheric conductivity during balloon ascents (Wigand 1925) that the rapid decrease in F with height above the ground is accompanied by an increase in λ of such magnitude as to keep the conduction currents approximately constant. Table 8 shows the results approximately constant. Table 8 shows the results of an ascent by Wigand (1925) from which it will be seen that when the field had fallen to 1/15th of its value at the ground, the conductivity had risen elevenfold.

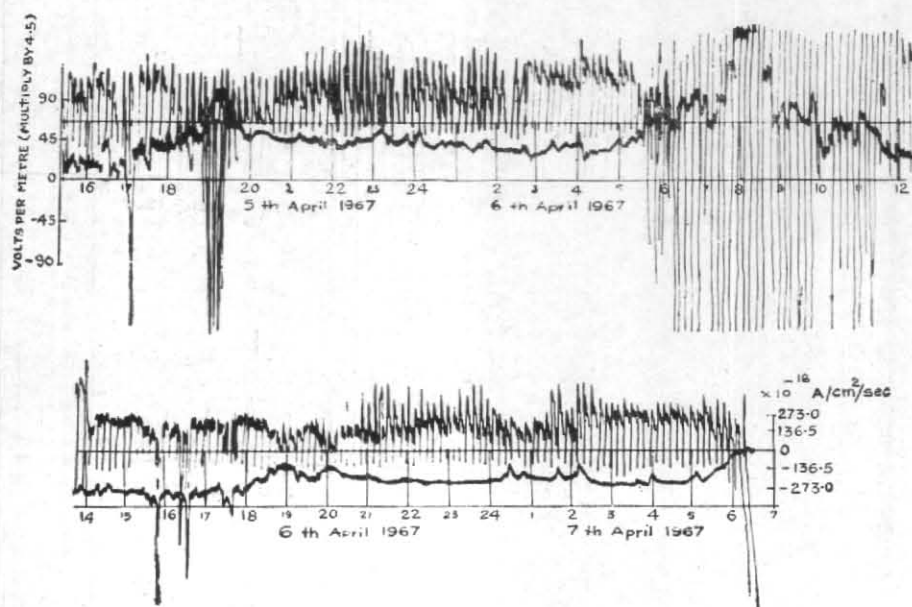


Fig. 8

Records of the positive and negative components of air earth current and positive potential gradient on 5-6 and 6-7 April 1967

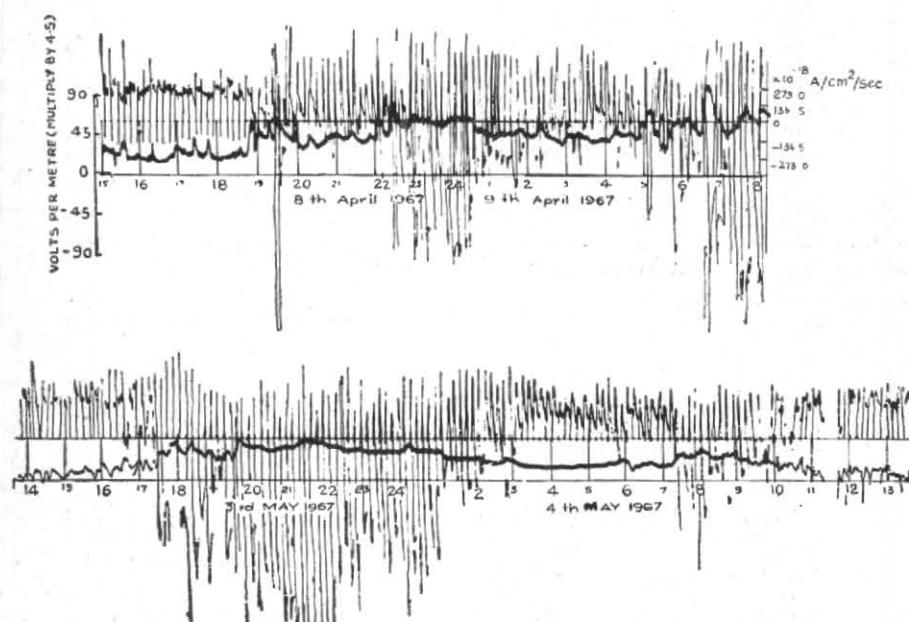


Fig. 9

Records of positive and negative components of air earth current and positive potential gradient on 8-9 April and 3-4 May 1967

These results have been confirmed and extended by conductivity measurements during the flight of Explorer II in 1935. They showed that the conductivity at 18 km was about 100 times the usual value at the ground.

The conductivity of the upper portion of the air path is entirely due to cosmic radiation and is unlikely to alter appreciably with time of locality or season. But below a height of about 2 km over land areas, an additional ionising influence is

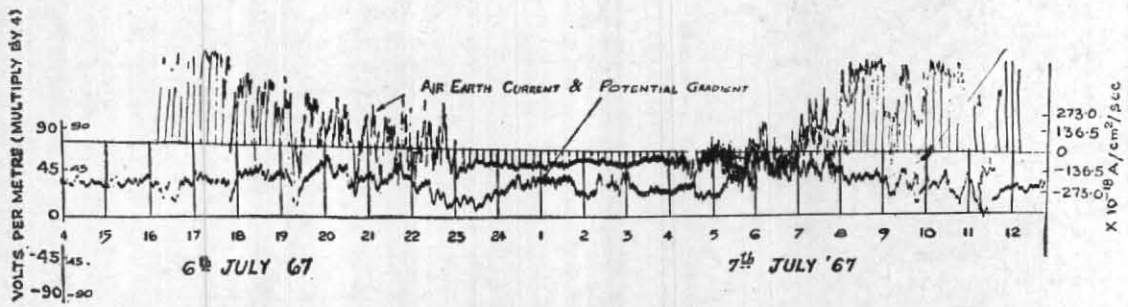
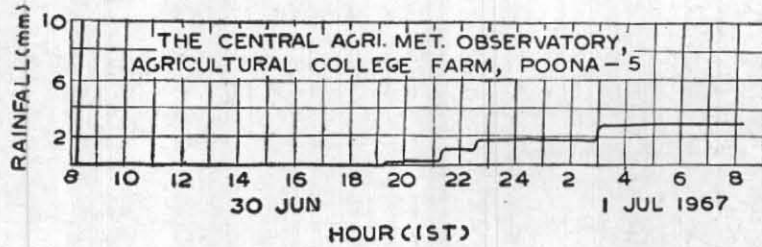
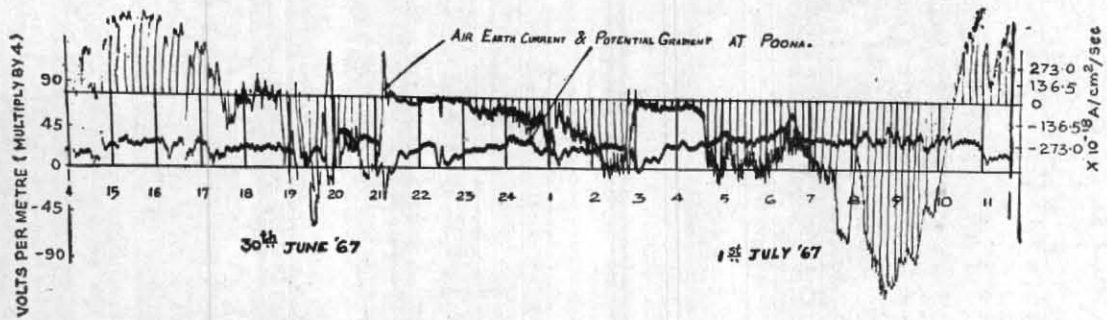


Fig. 10. Records of +ve and -ve components of air earth current and positive potential gradient

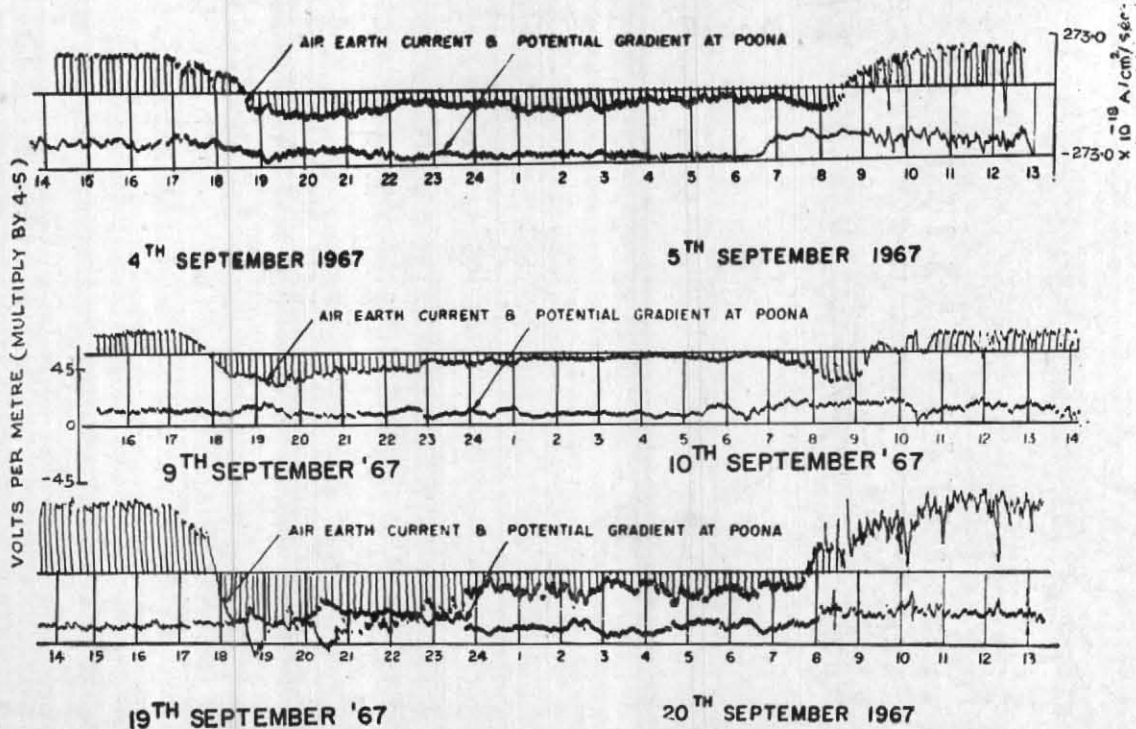


Fig. 11. Records of +ve and -ve components of air earth current and positive potential gradient

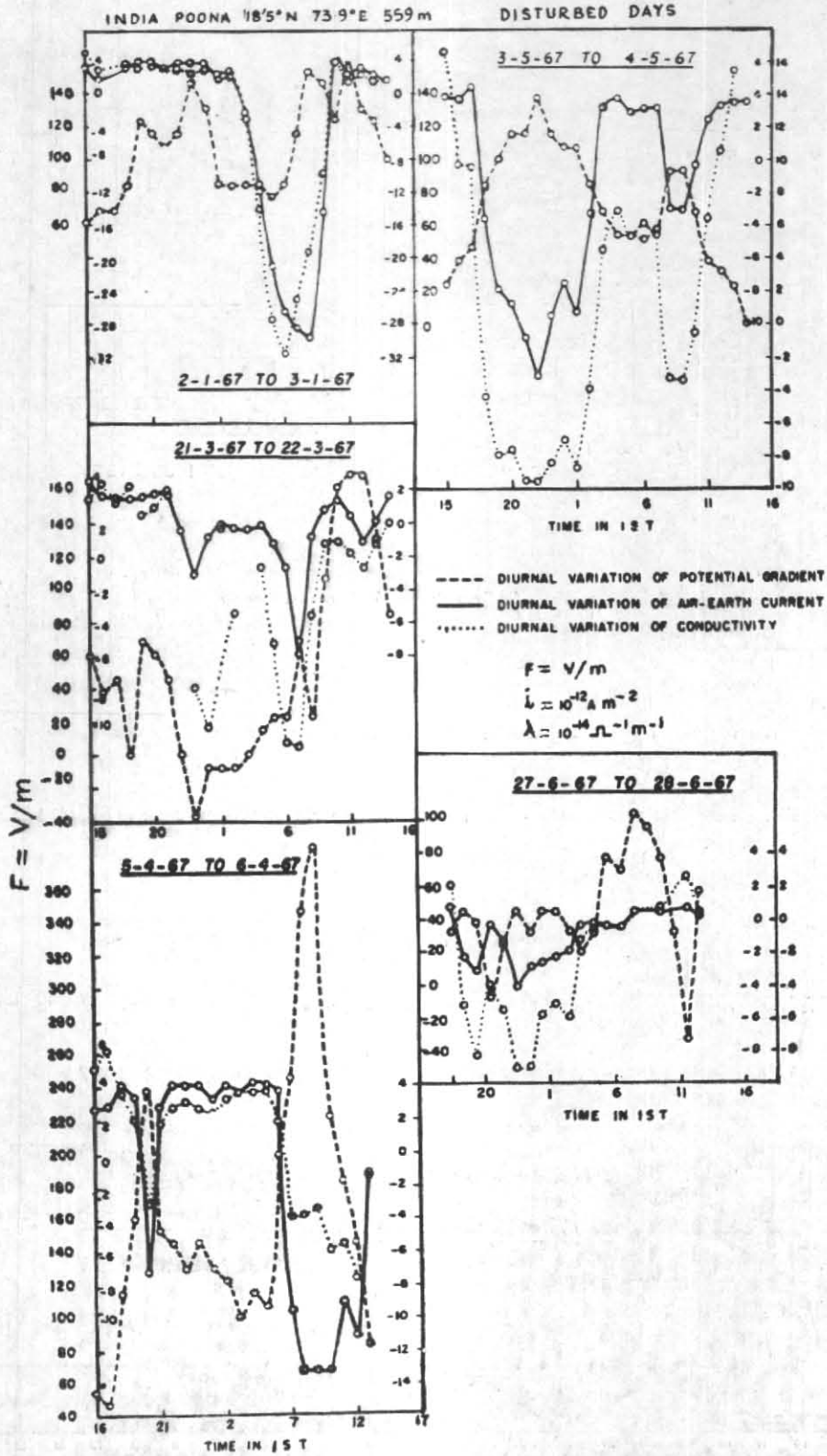


Fig. 12. Diurnal variations of air earth current, conductivity and potential gradient

exercised by radioactive matter in the air and earth and this varies very much. The current bearing column of air above the point of observation thus consists of an upper constant resistance and a lower variable one in series with each other. The extent to which the current carried by the air column alters with local conditions will depend on the ratio which the lower variable resistance bears to the total resistance.

Law (1963) made measurements at Cambridge on ion densities and of space charge and came to the conclusion that the conductivity varies with the height so that the air earth conduction current cannot be the same at all levels thus necessitating a convection current also varying with height. The results of Higazi and Chalmers (1966) lead to the same conclusion.

Dolezalek (1960) discussed a number of measurements of potential gradient, air earth current and conductivity. With Ohm's Law it is assumed that there is no convection current; then deviations from Ohm's Law would appear if a convection current actually exists; Dolezalek discussed the various other factors that would give apparent deviations from Ohm's Law and came to the conclusion that a comparison of results by the two methods (direct and indirect) of determining the air earth current is not simply sufficient to give the convection current, if any. It appears probably that the extent to which the convection current is important may vary from one place to another.

2: *The universal diurnal variations in potential gradient and air earth current*

The original Wilson hypothesis that thunderstorms and shower clouds act as the generators that maintain the fair weather aspects of atmospheric electricity is now almost universally accepted. A feature claimed to support the hypothesis is that the diurnal variations, related to U.T. (G.M.T.) of world thunderstorms activity and of the fair weather elements of atmospheric electricity at oceanic stations are in phase and of similar amplitude. In Fig. 14 the world-wide variations of thunder area is plotted from the data of Whipple and Scrase (1936). In deriving the curve for the whole world, an estimate for oceanic thunderstorms is also included. This has been omitted by several authors who have considered variations over land areas alone and thereby obtained curves whose amplitude about the mean is rather greater than it should be. Fig. 15 shows the diurnal variation with respect to U.T. of potential gradient and air earth current for observations at sea. The curves in Fig. 15 (b) are a mean derived from the results (1955) of several oceanic cruises

(Carnegie, Horizon, etc). It is seen that the two world curves of Figs. 15 (a) and (b) are broadly parallel. Fig. 16 shows the annual diurnal variation of thunderstorm for various stations in India (Tropics), indicating that the maximum frequency of thunderstorms take place at about 12 GMT. The major thunderstorm activity occurs in tropical regions where the storms are mainly of the convective type and reach their maximum activity from about mid-day to the late afternoon hours. This is the reason why the maximum of the curves for Asia, Africa and America follow in sequence at 0800, 1400 and 2000 GMT. When simultaneous measurements are made of any two of three quantities F , i and λ , it is possible to make useful comparison of the results at different places and to distinguish local and worldwide effects.

3. *Electrical balance sheet*

If the negative charge on the earth is to be maintained there must be for the earth as a whole a balance between different processes bringing charge to the earth. Wormell (1930) first discussed the possibility of working out such a balance in a small area of the earth and used available data to give results for Cambridge. Since then, more accurate data have become available for some more places, but so far no attempt has been made to calculate the electrical balance sheet for the tropics. Now it has been found from the air earth current observations at Poona that there are two components of air earth current corresponding to the positive and negative conductivities. These have to be taken into account in estimating the annual loss and gain of charge by the earth. Full details of the values by the different processes (conduction currents, point discharge current, precipitation current and lightning discharges) for Poona are being discussed in another paper.

4. *Calculation of the number of positive and negative ions from the continuous records of air earth current and potential gradient*

The specific conductivity λ_+ of air for positive ions is given by $\lambda_+ = i_+/F$ where, i_+ is current density of positive ions and F is the potential gradient. Similarly for specific conductivity λ_- for negative ions is given by $\lambda_- = i_-/F$. If n_+ and n_- are the number of positive and negative ions per c.c. with mobilities k_+ and k_- and charge $\pm e$,

$$\begin{aligned}\lambda_+ &= n_+ \cdot e \cdot k_+ = i_+/F \\ \lambda_- &= n_- \cdot e \cdot k_- = i_-/F\end{aligned}$$

Assuming the values of ionic charge e and mobility of ion as 1.5 cm in a field of 1 volt/cm, the values of n_+ , n_- are calculated from the known values of i_+ , i_- and F . It is seen from Tables 5-7 that the values of n_+ and n_- range from 100 to 13,000.

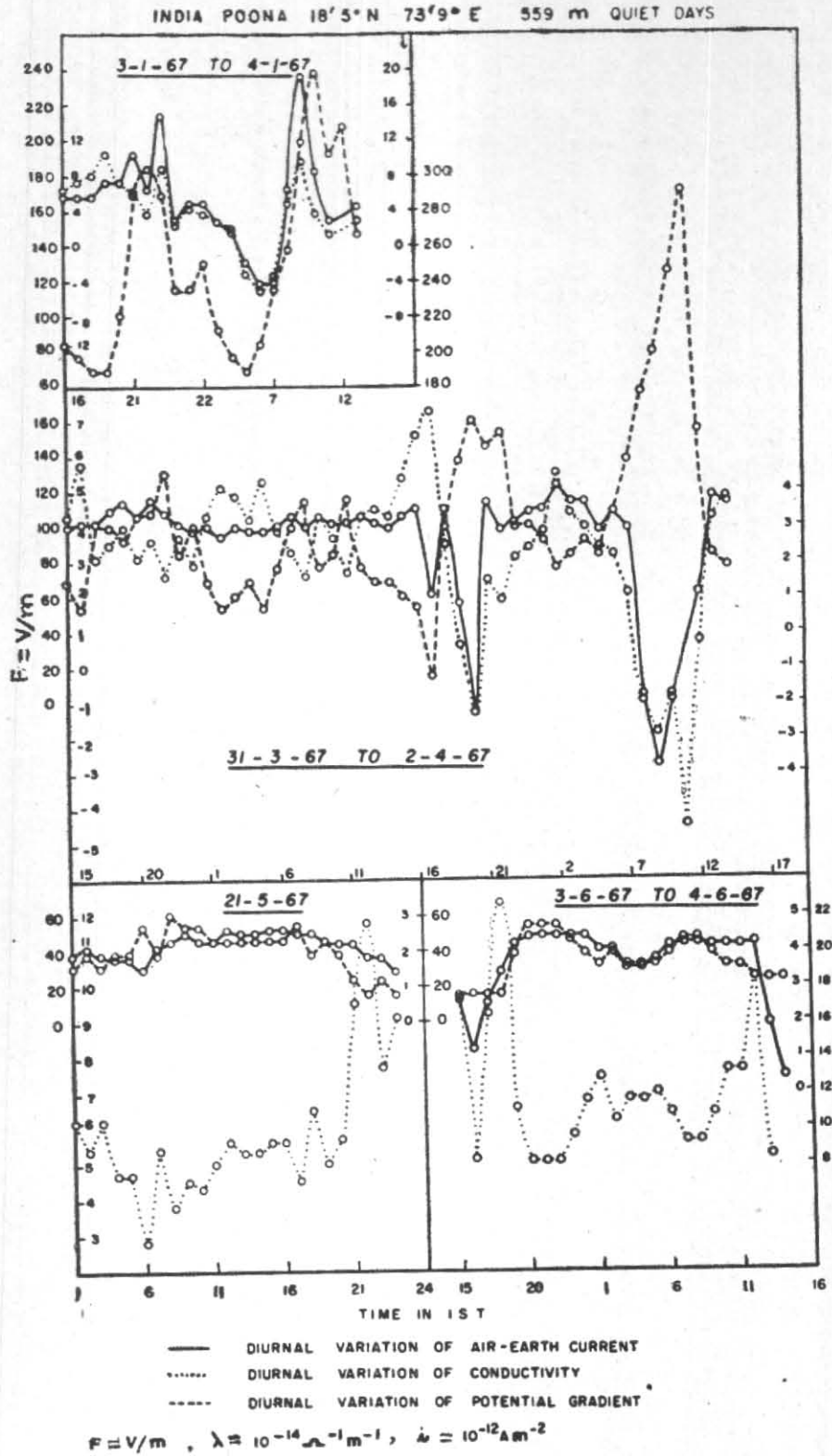


Fig. 13. Diurnal variation of air earth current, conductivity and potential gradient

Days are classified in Tables 5-7 as quiet days and disturbed days. Quiet days are days during which the air earth current throughout the day is +ve. Disturbed days are days during which the air earth current is showing both +ve and -ve currents. Generally the number of small ions per c.c. in oceanic air is 700, in country air is 600 and in city air is 100. The number of positive and negative ions are not very different but there is necessarily a small excess of positive ions to make up the positive space charge of the lower atmosphere. Measurements of the number of small ions (positive and negative) using Ebert's ion counter at Poona are given in Table 9. From this, we see that the number of positive and negative ions range from 900 to 6000. But the measurements of the number of positive and negative ions from the continuous records of air earth current and potential gradient generally range from 100 to 13,000. It is difficult to understand how the average values of small ions, positive and negative or the various hours range from a low value to a very high value, some times.

At first sight, it might appear that an ion carrying an electric charge and situated in an electric field, would experience a definite force and would therefore be accelerated acquiring a velocity which would not be constant but which would continuously increase. However, this neglects the fact that the ion experiences incessant impacts with the molecules of the air and so loses at each impact some or all of the momentum it has acquired since the previous impact. There will thus be an average velocity of travel of the ion in the direction of the potential gradient and proportional to the potential gradient; clearly the longer the mean free path of the ion, the greater is the mobility.

Actually the existence of two main types of ion (large and small), leads to the flow of four different ion streams, two up and two down if the field is vertical, in each of which the velocity is different, being given by $F \times k$, where, k is the corresponding mobility. The total current is thus —

$$i_+ + i_- = F e (n_+ k_+ + n_- k_- + N_+ K_+ + N_- K_-)$$

where, k_+ , k_- , K_+ , K_- are the mobilities of the positive and the negative ions, small and large respectively, and e is the elementary charge. In clear country air over land n_+ and n_- number about 600 per c.c. while N_+ and N_- are about 2000 per c.c., k_+ and k_- are 1.5 approximately and K_+ and K_- are 0.0004. Here, therefore, the ratio of the first to the second term in the bracket given above is 1100 : 1 and therefore most of the conductivity of the air is due to small ions. Thus

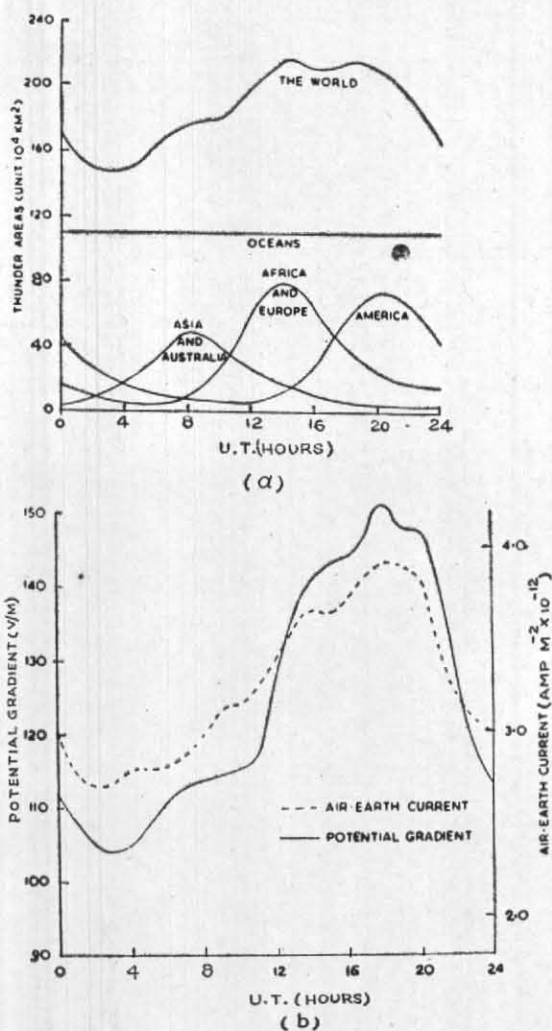


Fig. 14

Diurnal variation with respect to U.T. of (a) global thunder areas (After Ruthenberg, S. and Holzer, R.E. 1955) and (b) of air earth current and potential gradient at sea (After Pierce, E.T. 1958)

TABLE 8

Conduction current during balloon ascent (Wigand)

Height (m)	F (V/m)	λ (e.s.u.)	1 amp per sq. cm
0	136	1.1×10^{-4}	1.7×10^{-10}
2500	27	4.8×10^{-4}	1.4×10^{-10}
4400	18	8.2×10^{-4}	1.6×10^{-10}
6500	8	12.6×10^{-4}	1.2×10^{-10}

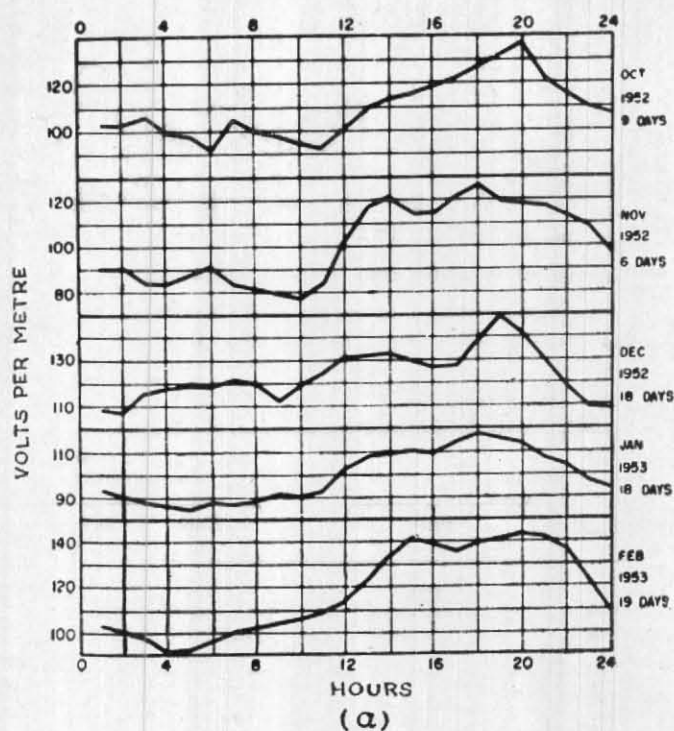
TABLE 9
Measurement of +ve and -ve ions using Ebert's Ion Counter

Date 1967	Time (IST)	Ions		Date 1967	Time (IST)	Ions		Date 1967	Time (IST)	Ions		Date 1967	Time (IST)	Ions	
		+ve	-ve			+ve	-ve			+ve	-ve			+ve	-ve
12 Jun	1740	2855	3295	24 Jun	0930	2915	3375	5 Jul	0900	2621	2139	13 Sep	0800	1443	1489
13 "	0800	2648	2994		1200	2276	2169		1204	2429	2594		1200	1627	1781
	1200	2585	2277		1608	2955	2957		1600	—	—		1600	1992	1878
	2000	2096	2011		1905	1529	1585		2000	—	—	14 "	0800	1569	1408
14 "	0805	4392	4387	25 "	0900	1765	1956	6 "	0800	—	—		1200	2164	1607
	1205	4875	3960		1215	2648	2345		1200	3115	2990		1600	1599	1379
	1605	4463	4192		1610	2404	2075		1600	—	—	15 "	0800	1383	1134
	1955	4566	4804		2030	1817	1711		2000	—	—		1200	1865	1159
15 "	0800	5774	6486	26 "	0900	1765	1956	7 "	0800	2055	1714		1600	1785	1242
	1200	5839	5436		1215	2648	2345		1200	—	—		2000	1214	486
	1600	3880	3959		1610	2404	2075		1600	2082	2171	16 "	0800	788	1059
	2000	2845	2698		2042	1427	1419		2000	—	—		1200	1449	1147
16 "	0800	—	—	27 "	0820	2559	2256	1 Sep	0800	1920	1686		1600	1335	1261
	1206	2401	3116		1210	2788	2872		1200	2250	1776		2000	1254	1241
	1605	3781	2413		1600	1855	1609		1600	1319	1314	17 "	0800	1340	836
	2010	2620	2422		2015	1500	1092		2000	1069	743				
17 "	0810	3437	4260	28 "	0845	1226	1425	2 "	0800	1209	1306	18 "	0800	1447	1246
	1205	3051	4349		1206	2970	2781		1200	1608	1468		1200	1886	1832
	1605	2026	1728		1605	2680	2524		1600	1293	1271		1600	1443	1275
	2010	2464	2207		2005	—	1866	3 "	1200	1540	1525		2000	1175	1533
18 "	0825	2401	2226	29 "	0905	2433	2374		1600	1434	1369	19 "	0800	1165	929
	1220	4275	2792		1200	3738	3362		2000	1160	877		1200	1421	1392
	1540	3531	3033		1610	2803	2867	4 "	0800	1271	1361		1600	1269	1250
	2000	—	—		2100	3111	1191		1200	1912	1596		2000	865	628
19 "	0830	3000	3291	30 "	0900	1406	3036		1600	1308	1375	20 "	0800	1218	978
	1210	4669	4215		1200	2853	2985		2000	890	1615		1200	2255	2232
	1605	4894	3604		1605	3208	2729	5 "	0800	1065	1064		1600	1480	1490
	2000	4894	3604		2030	1372	1874		1200	1386	1481		2000	1520	1275
20 "	0850	2207	1484	1 Jul	0845	1125	2217		1600	1389	1353	21 "	1200	2252	1747
	1200	2915	2649		1210	2375	2041		2000	1216	1379		1600	1480	1007
	1600	5061	1047		1625	2584	2391	6 "	0800	1366	1356	22 "	1600	1349	1231
	2000	—	—		2050	1888	2685		1200	1499	1922				
21 "	0900	1577	1147	2 "	0800	—	—		1600	—	1070	23 "	1200	1994	2009
	1200	2713	2038		1200	—	—	7 "	0800	1195	1235	25 "	1200	1782	1540
	1600	2013	2140		1500	2685	2186		1200	1336	1335		1600	1913	1578
	2000	1429	1115		2000	—	—		1600	936	973				
22 "	0800	2199	2707	3 "	0800	—	—		2000	1070	1202	26 "	1200	1994	1421
	1200	2783	2993		1205	2985	3163	8 "	0810	1471	1399	27 "	1200	1662	1711
	1600	1864	1958		1600	2472	2013		1200	1833	1695				
	2000	2159	1533		2015	1674	1726		1600	1244	1427	28 "	1200	1970	1583
23 "	0810	2373	1663	4 "	0930	—	—	9 "	0800	1417	1070	29 "	1200	12339	1488
	1225	3192	3244		1215	2082	2540		1200	1391	850		1600	2040	1177
	1605	2898	2969		1605	2374	1792	12 "	1200	1929	1734	30 "	1200	2579	1782
	2000	—	—		2005	1867	1864		1600	1372	1698		1600	1258	1114

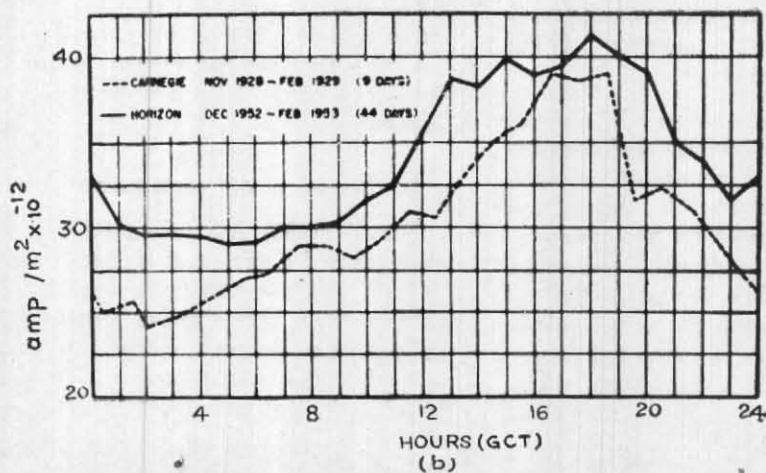
Monthly mean values of positive conductivity, potential gradient and air earth current density

Date	Positive conductivity (ohm ⁻¹ m ⁻¹) × 10 ⁻¹⁴	Potential gradient (Volts/m)	Air earth current density (amp/m ²) × 10 ⁻¹²
7 Oct 1952	1.58	108.4	2.95
3-7 Nov 1952	2.08	100.4	4.18
22 Dec 1952	1.68	123.1	4.18
18 Jan 1953	1.45	98.9	2.95
19 Feb 1953	1.19	107.8	2.55

The monthly means of the diurnal variation of the potential gradient are shown as a function of Greenwich Civil Time in figure below. All of the monthly curves exhibit a single daily oscillation



(a) Reduced values of the potential gradient in volts/metre taken aboard the R/V *Horizon* in the Central Pacific Ocean during Oct 1952 to Feb 1953



(b) Computed values of air earth current density in the Pacific Ocean (*Carnegie* 1928-29 and *Horizon* 1952-53)

Fig. 15

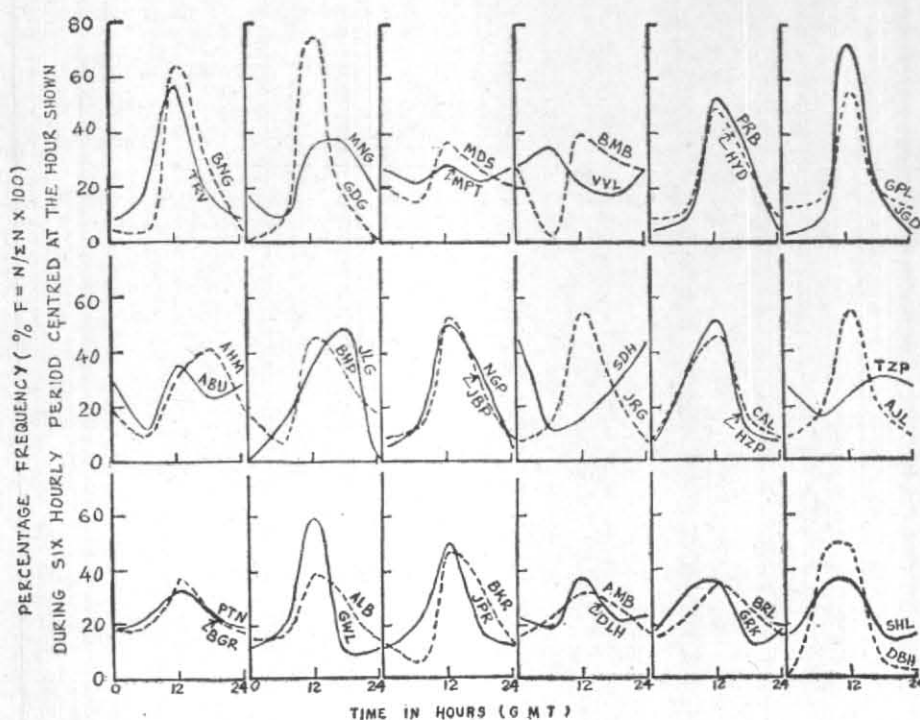


Fig. 16. Diurnal variation of thunderstorms (1941-45) annual total

N = No of occurrences for each six hourly period of a season

ΣN = Total No. of occurrences during all the diurnal periods

TABLE 10

Quantity	Units	Poona		Other stations (as reported by Chalmers 1967)		
		1966	1967	Kew	Average land station	Ocean
Potential gradient	V/m	101	100	365	130	126
Air earth current	Δ/m^2	1.51×10^{-12}	1.50×10^{-12}	1.12×10^{-12}	214×10^{-12}	3.7×10^{-12}
Conductivity	$\Omega^{-1}m^{-1}$	1.61×10^{-14}	1.53×10^{-14}	3.0×10^{-15}	1.8×10^{-14}	2.8×10^{-14}

the part played by large ions is usually negligible, as far as ionic current is concerned although they play an important part in the question of the number of small ions present.

The above factors have to be taken into account in calculating the number of positive and negative ions from the continuous records of air earth current and potential gradient. In the normal use of a quadrant electrometer for the measurement of currents one pair of quadrants is connected to the current collector while the other pair is earthed. In the present method, the second pair is not earthed. The second pair is connected to the bottom plate. The top plate is connected to the other pair. If now both pairs of quadrants are first earthed and released, the top plate would receive one component of the current (+ve)

and the bottom would receive the (-ve) component of the current. As compensation for displacement currents due to field changes has been arranged, the predominant current (+ve or -ve) will be shown by the deflections of the quadrant electrometer.

5. Changes of air earth current during fog, mist

Scruse (1933) has observed increase of field strength and decrease of air earth current density in fog. He has also noted that during fog convection, with high relative humidity, very often the measured air earth current seems to be negative while the field is positive. Measurements by Israel and Kasemir (1952) of the potential gradient and conduction current during widespread fog showed a considerable increase in potential

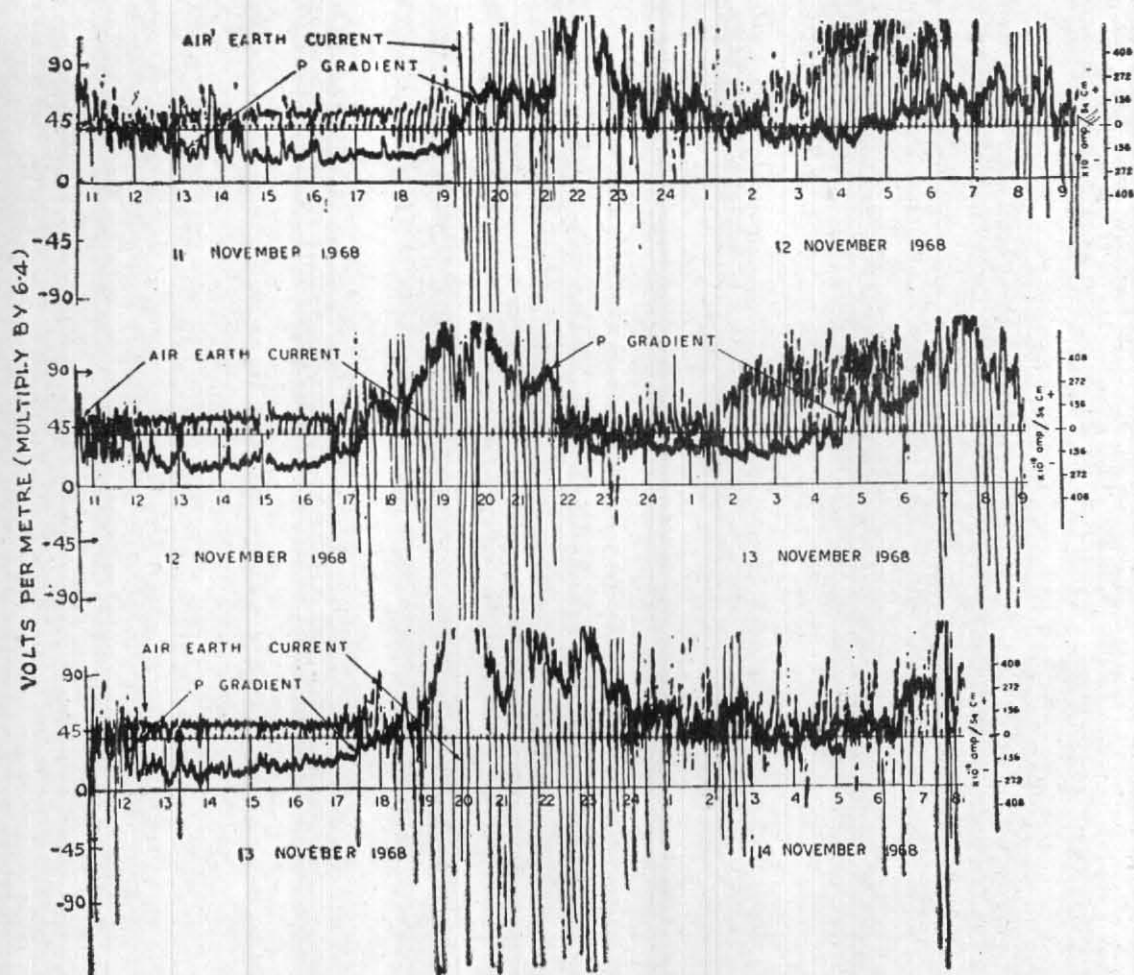


Fig. 17. A good example of effect of fog, mist on potential gradient and air earth current at Poona

The normal position of air earth current started decreasing before the onset of fog with corresponding increase in potential gradient from 19 IST on 11, 17 IST on 12 and 18 IST on 13 Nov 1968

gradient and a slight decrease in conduction current as compared with clear air, giving on the average, a conductivity about one-third of its value in fine weather. The same facts have been observed by a number of other workers in various places and a very thorough discussion has been given by Dolezalek (1962). Serbu and Trent (1958) at Argentina in New Foundland found that the decrease in conductivity and consequent increase in potential gradient became evident 1 to 2 hours before the onset of fog and the reverse changes appeared $\frac{1}{2}$ to $1\frac{1}{2}$ hours before the dissipation of the fog. Dolezalek (1962) had stated that the above results, while not 100 per cent reliable for forecasting gave around 80 per cent reliability.

Venkiteswaran (1958) at Poona had observed a strong increase of the electric field about 10 hr before the occurrence of fog as measured by a potential gradient radiosonde in October 1953. This increase of potential gradient had occurred in altitudes between 900 and 680 mb.

For the first time at Poona, surface observations of air earth current and potential gradient during 1966, 1967 and 1968 have shown increase of potential gradient and decrease of air earth current density during fog and mist in agreement with the results of Scrase (1953) and Israel and Kasemir (1952).

TABLE 11

Hourly values recorded at Poona of potential gradient (F), two components of air earth current (i_+ , i_-), two components of conductivity (λ_+ , λ_-) and number of positive and negative ions per c.c. (n_+ , n_-) during fog, mist day in the year 1968

Time (IST)	Quantity											
	F	i_+ , i_-	λ_+ , λ_-	n_+ , n_-	F	i_+ , i_-	λ_+ , λ_-	n_+ , n_-	F	i_+ , i_-	λ_+ , λ_-	n_+ , n_-
	11 to 12 Nov 1968				12 to 13 Nov 1968				13 to 14 Nov 1968			
16-17	121	1.1	0.9	375	173	—	—	—	138	1.0	0.7	292
17-18	150	1.2	0.8	333	357	2.0	0.6	250	184	1.8	1.0	417
18-19	150	1.6	1.1	131	576	-7.2	-1.3	-541	311	0.5	0.2	83
19-20	277	-2.6	-0.9	-375	680	-15.0	-2.2	-917	588	-10.9	-1.9	-791
20-21	461	-12.7	-2.8	-1167	621	-13.7	-2.2	-917	621	-10.9	-1.8	750
21-22	—	—	—	—	518	-10.1	-1.9	-791	657	-12.2	-1.9	-791
22-23	—	—	—	—	242	1.4	0.6	250	621	-12.2	-2.1	-875
23-24	—	—	—	—	219	-2.3	-1.1	-458	610	-10.9	-1.8	-750
0-1	369	5.4	1.5	624	196	2.6	1.3	541	403	3.3	0.8	333
1-2	265	3.0	1.1	458	207	-0.1	0	0	311	2.3	0.7	292
2-3	253	2.9	1.1	458	161	2.2	1.4	583	357	-2.6	-0.7	-292
3-4	253	4.4	1.7	708	219	5.0	2.3	958	265	1.0	0.4	167
4-5	—	—	—	—	334	5.2	1.6	668	230	1.4	0.7	292
6-5	—	—	—	—	415	6.1	1.5	624	277	1.9	0.7	292
6-7	—	—	—	—	565	1.0	0.2	83	392	-0.7	-0.2	-83
7-8	—	—	—	—	726	-10.9	-1.5	-624	610	6.1	1.0	417
8-9	—	—	—	—	578	-10.7	-2.1	-875	—	—	—	—
	9 to 10 Dec 1968				18 to 19 Dec 1968				19 to 20 Dec 1968			
15-16	—	—	—	—	230	2.9	1.3	541	196	3.8	1.9	791
16-17	150	2.6	1.7	708	253	4.1	1.6	668	357	3.8	1.1	458
17-18	150	2.4	1.6	668	438	-1.5	-0.3	-125	300	1.0	0.3	125
18-19	323	-0.8	-0.2	-83	668	4.4	0.7	292	403	-2.9	-0.7	-292
19-20	703	5.7	0.8	333	657	-1.6	-0.2	-83	323	3.3	1.0	417
20-21	541	-0.5	-0.1	-42	691	-13.3	-1.9	-791	357	-9.5	-2.7	1126
21-22	334	4.6	1.4	583	530	-2.7	-0.5	-208	346	2.0	0.6	250
22-23	311	4.7	1.5	624	369	0.3	0.1	42	219	2.3	1.1	458
23-24	230	2.9	1.3	541	392	-1.9	-0.5	-208	357	4.1	1.1	458
0-1	161	4.8	1.1	458	369	-1.4	-0.4	-167	161	2.0	0.8	333
1-2	161	1.5	0.9	375	438	-5.3	-1.2	-500	392	3.8	1.0	417
2-3	115	1.6	1.4	583	565	1.8	0.5	292	323	-2.7	-0.8	-333
3-4	150	2.6	1.7	708	472	6.1	1.3	541	150	-1.0	-0.7	-292
4-5	104	1.5	1.4	583	311	1.0	0.3	125	196	1.5	0.7	292
5-6	92	1.0	1.1	458	461	-9.5	-2.1	-875	461	-2.6	-0.6	-250
6-7	127	0.7	0.6	250	496	4.5	0.9	375	621	-2.2	-2.0	-833
7-8	138	1.4	1.0	417	703	-13.7	-1.9	-791	588	-2.2	-2.1	-875
8-9	—	—	—	—	—	—	—	—	392	5.7	1.5	624
9-10	—	—	—	—	—	—	—	—	472	5.7	1.2	500

F is expressed in V/m

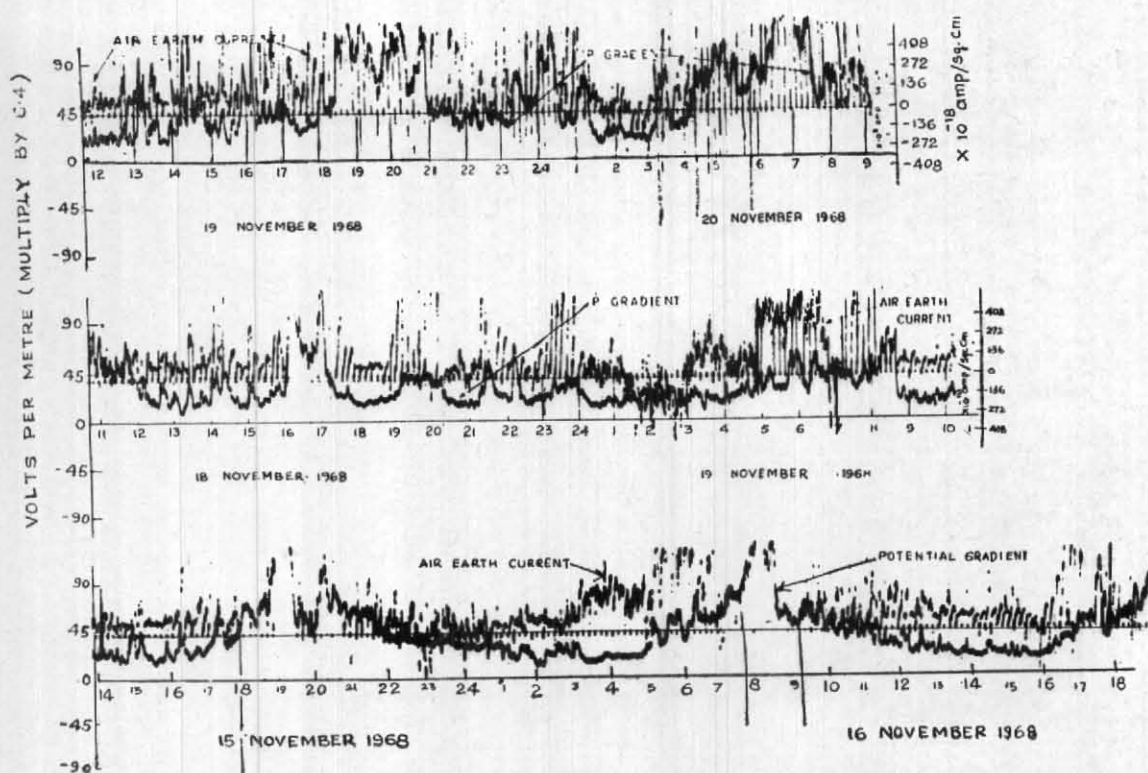


Fig. 18. Records of mainly positive air earth current and positive potential gradient

A few records showing the increase of potential gradient and decrease of air earth current during fog, mist at Poona are reproduced in Figs. 17, 18, 19, 20 and Table 11. Scrase (1933) has explained the general effect of fog on electrical conditions as follows —

The resistance of the air is increased owing to the diminution in the number of small ions. This diminution is brought about by the absorption of the small ions by the nuclei of condensation, the numbers of which are increased in fog. The increased resistance causes the potential gradient to assume high positive values as a rule, whilst the air earth current is zero and becomes negative. The surprising fact is that there is a reversal in the sign of the charge, entering the ground, and the reversal persists for several hours. But this is not accompanied by a corresponding change in the sign of the potential gradient. The occurrence of opposite signs for field and current is explained by Chalmers and Little (1941) by Maxwellian displacement currents generated by field changes. Basically during fog, the conductivity is decreased to about one-third. The decrease in conductivity may be reflected by a decreased air earth current density or by an increased field strength or by both. In addition, the ratio of negative to

positive polar conductivities may undergo some variations in connection with fog.

7. Summary and conclusions

(i) During 1966, in undisturbed conditions, the annual mean of the air earth current at Poona is 151.1×10^{-18} amp per sq. cm per sec. The charge due to above conduction current reaching 1 km^2 of earth in a year at Poona is 48 Coulombs as compared with 35.3 Coulombs per km^2 at Kew (Scrase).

(ii) The annual variation of the air earth current is the same as that of the potential gradient, there being a maximum in winter and minimum in summer.

(iii) In winter the daily fluctuations of current and conductivity are very much alike. In summer on the other hand, there is more resemblance between the diurnal variation of the air earth current and that of potential gradient.

(iv) During 1967, with the introduction of two radioactive collectors, one for each test plate, both the components of air earth current corresponding to the positive and negative conductivities are being recorded in the air earth current records.

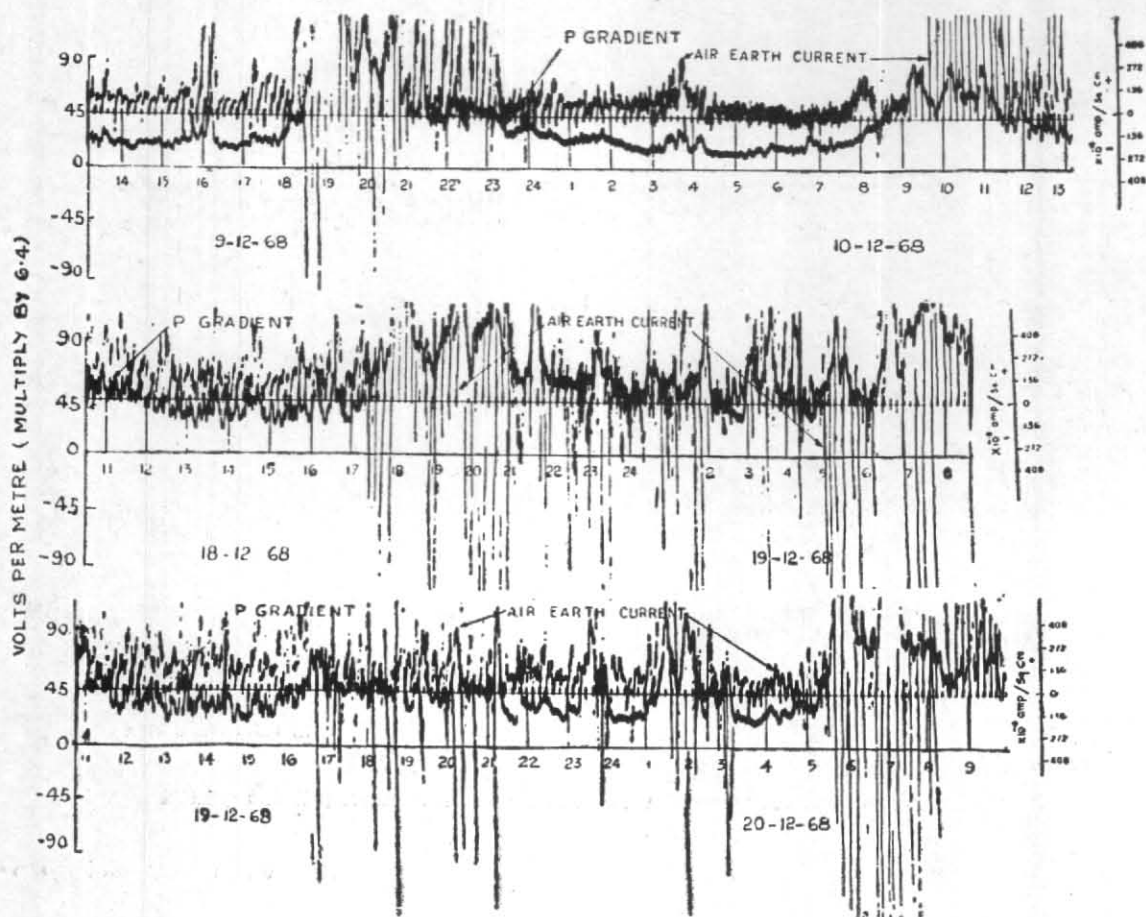


Fig. 19

Another example of the effect of fog, mist on potential gradient and air earth current on 9 and 18 Dec 1968 at Poona.

There was widespread fog commencing from 19 IST on 18 Dec 68 to 08 IST on 19 Dec 68. On 19 Dec 68 there was fog again commencing from 1700 IST on 19 Dec 68. On 9 Dec 68 the potential gradient started increasing from 18 IST with corresponding decreasing in air earth current changing the direction from positive to negative. On 18 Dec 68, at 1700 IST the potential gradient started increasing with corresponding decrease in air earth current and changing the direction from positive to negative at 1730 IST. On 19 Dec 68, at 1630 IST the potential gradient started increasing with corresponding decrease in air earth current and changing direction from positive to negative at 1630 IST.

(v) The annual total mean value of positive air earth current and negative air earth current for the year 1967 is 158.1×10^{-18} amp/sq. cm/sec. The charge due to above conduction current reaching 1 km^2 of earth in a year at Poona is + 50 Coulombs as compared with +60 Coulombs at Cambridge (Wormell 1953) and +60 Coulombs at Durham (Chalmers and Little 1947) and +90 Coulombs over the whole world (Israel 1953).

(vi) Table 10 gives a comparison of average values of atmospheric electric parameters at Poona

(1966-67) and other stations.

(vii) During fog, mist at Poona, the potential gradient increases, but the air earth current decreases and becomes negative in sign. The potential gradient increases about 1 to 2 hours before the onset of fog with corresponding decrease in air earth current. The reverse changes in both potential gradient and air earth current appear $\frac{1}{2}$ to $1\frac{1}{2}$ hours before the dissipation of fog. This fact is useful for forecasting of fog formation and dissipation.

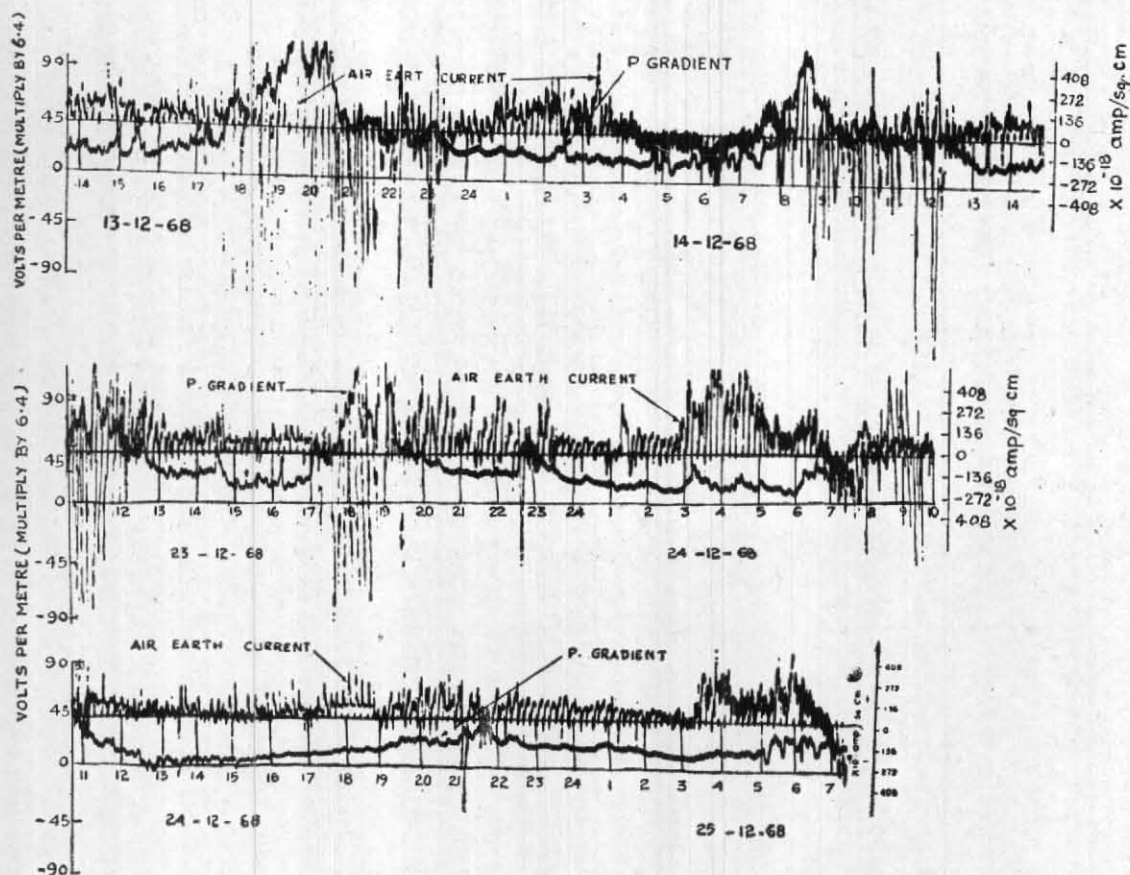


Fig. 20

The figure shows another example of fog, mist effect on potential gradient and air earth current on 13 Dec 68 with increase in potential gradient from 1700 IST and corresponding decrease in air earth current changing the direction from positive to negative at 1730 IST on 13 Dec 68. On 23 Dec 1968 the potential gradient started increasing at 17 IST with corresponding decrease in air earth current at 1715 IST. On 24-25 Dec 1968, the whole day, the air earth current and potential gradient are both positive in direction without the effect of any fog, mist.

Acknowledgements — Almost from the commencement of this investigation the Late Dr. J. A. Chalmers of Durham University has been helping us with his valuable suggestions for this work and we therefore take this opportunity

to express our heart felt thanks for his valuable help. We are also thankful to Miss A. Mani, Director (Instruments), Meteorological Office, Poona for supplying us the necessary equipment for this investigation.

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