

Models of atmospheric refractive index with reference to coastal stations of India

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ABSTRACT. Based on the radiosonde data during the years 1957-1960, the effective earth's radius and the exponential model of atmospheric refractive index over coastal stations of India have been deduced. The correlation coefficient and the least squares straight line fit between the gradient of refractive index in the lowest one kilometre, ΔN and that near the ground surface, N_s has been studied and compared with the earlier studies based on the data of one year. The radio climatology of refractive index at 900-mb level and that of ΔN and their diurnal variation have also been briefly described.

1. Introduction

Since the installation of weather radars over the coastal stations of the country, angel activity associated with sea breeze has attracted the attention of several workers (Rai 1960, 1961; Datar and Sikdar 1964). The propagation characteristics when compared with that over inland stations are found to be distinctly different. It appears therefore, useful to evolve models of radio refractive index over the coastal stations rather than attempting a single model over the entire country as done hitherto (Srivastava 1968). It may also be mentioned that besides this aspect, the earlier study was based on one year data for the four months recommended by International Radio Consultative Committee (CCIR) and thus the results reported may not be truly representative for radio climatic purpose.

The object of this study is, therefore, to derive the various models of atmospheric refractive index over the coastal stations of India and to examine the seasonal and diurnal variation of the gradient in the lowest one kilometre so that transmission loss and refraction effects may be estimated. The radio climatology at 900-mb level has also been briefly described in the paper.

2. Models of radio atmospheres over coastal stations of India

Effective earth's radius model—The effective earth's radius model assumes an earth larger than the actual earth to the extent that the curvature of the radio ray may be absorbed in the curvature of the effective earth so that radio rays may be drawn as straight line over this earth rather than curved rays over the true earth. The effective

earth's radius k is given by —

$$k = \frac{1}{1 + \frac{a}{n} \frac{dn}{dh} \cos \theta} \quad (1)$$

where a is the true radius of the earth, n is the refractive index, and dn/dh is the initial gradient with respect to height. The distance to the radio horizon, h of the radio ray leaving an antenna of height, h can be calculated from $\sqrt{2kh}$ for rays tangential to the earth ($\theta_0 = 0$) and assuming n to be unity, equation (1) is approximated by

$$k \approx \frac{1}{1 + a \frac{dn}{dh}} \quad (2)$$

The values of dn/dh computed from the gradient of radio refractive index in the lowest one kilometre (Table 1) was substituted in (2) for calculating the effective earth's radius k .

Exponential model of radio refractive index—The exponential model of radio refractive index is given by —

$$N(h) = A \exp(-Bh) \quad (3)$$

where $N(h)$ is the radio refractive index at height h and A and B are constants which can be determined by least squares method from the radio climatological data.

3. Estimation of the gradient of refractive index in the lowest 1 km from surface value

On account of availability of large number of surface weather observations several times a day, attempts are being made to estimate the gradient of radio refractive index ΔN from that near

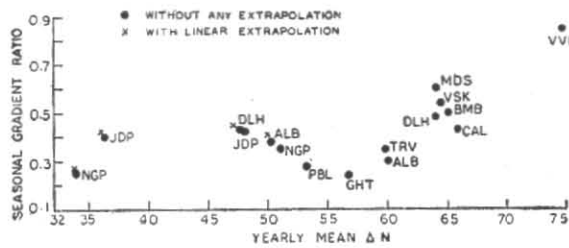


Fig. 1. Radio climatic classification of ΔN over India

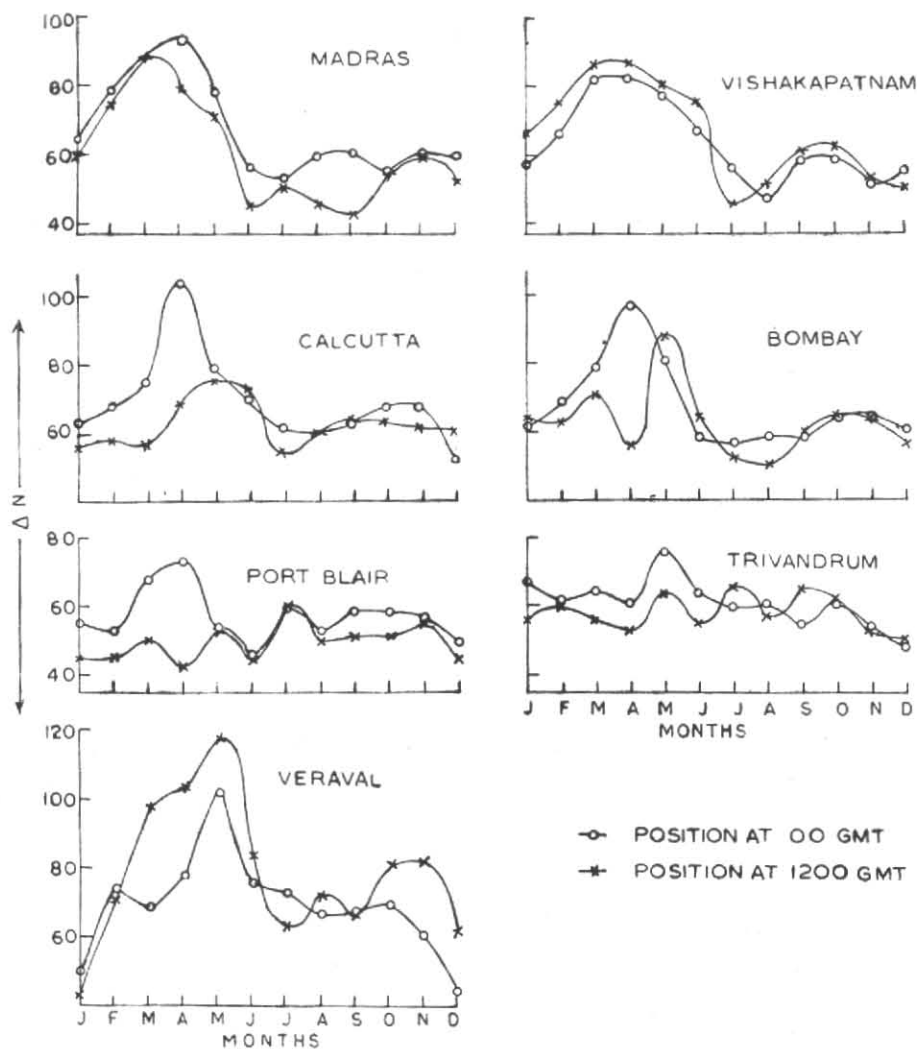


Fig. 2. Monthly variation of ΔN

TABLE 1

Gradient of Radio Refractive Index (ΔN) in the lowest 1 km over coastal stations of India (1957-1960)

| Station | Jan | Feb | Mar | Apr | May | Jun | Jul | Aug | Sep | Oct | Nov | Dec |
|---------------|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|
| Bombay | 63 | 66 | 75 | 76 | 84 | 62 | 55 | 54 | 59 | 64 | 64 | 58 |
| Calcutta | 59 | 62 | 65 | 85 | 82 | 71 | 58 | 60 | 63 | 65 | 64 | 56 |
| Visakhapatnam | 62 | 71 | 84 | 84 | 79 | 71 | 51 | 49 | 60 | 60 | 52 | 53 |
| Veraval | 47 | 77 | 84 | 91 | 110 | 80 | 68 | 69 | 67 | 76 | 72 | 54 |
| Madras | 63 | 77 | 89 | 87 | 75 | 51 | 52 | 52 | 51 | 60 | 60 | 53 |
| Trivandrum | 62 | 61 | 61 | 58 | 70 | 60 | 63 | 59 | 61 | 62 | 55 | 50 |
| Port Blair | 50 | 49 | 59 | 58 | 54 | 45 | 60 | 52 | 55 | 55 | 56 | 48 |

TABLE 2

Correlation Coefficient between ΔN versus N_s over India

| | $\Delta N, N$ (1957-1960) | $\Delta N, N_s^*$ (1965) |
|----------|------------------------------|-----------------------------|
| February | 0.71 | 0.97 |
| May | 0.66 | 0.83 |
| August | 0.77 | 0.43 |
| November | 0.61 | 0.83 |

*Srivastava and Chatterjee (1957)

the ground surface, N_s . The relation between N_s and ΔN is given by —

$$\Delta(N) = CN_s + D \quad (4)$$

where C and D are constants which may be determined by least square method and neglecting the sign of ΔN which is negative.

4. Data

The daily values of atmospheric pressure, dry bulb temperature and dew point as obtained by 12 Indian radiosonde stations (Table 4) at 00 and 12 GMT for the years 1957-1960 were converted into mean monthly values of refractive index near the ground surface and at 900 mb. It is noticed from *Indian Daily Weather Reports* that the height of 900-mb surface remains close to 1 km throughout the year at all the coastal stations. However, for the inland stations of the country, height of 900-mb level shows well marked seasonal variation and may be as low as 850 m or so, after subtracting the station elevation, for some months.

Bean and Dutton (1966) suggested that the climatic variation of ΔN may be represented by plotting the ratio $R(\Delta N)$ of the highest and lowest values of ΔN versus the mean value of ΔN . The results given in Fig. 1 show that the grouping of 12 radiosonde stations does not truly reflect the climatic characteristics at all the stations. In case ΔN values ($N_s - N_{900}$) at inland stations are linearly extrapolated for 1 km, a rather different type of classification is obtained for the same set of stations which again does not delineate the climatic characteristics besides showing unrealistically large values of ΔN for some inland stations. On this account, the correlation coefficient (CC) between ΔN versus N_s has been studied without any extrapolation but the limitations as mentioned above may be kept in view while interpreting the results.

5. Results and Discussion

The values of ΔN for seven coastal stations in India are given in Table 1.

TABLE 3
Effective earth's radius factor over coastal stations of India

| Station | Jan | Feb | Mar | Apr | May | Jun | Jul | Aug | Sep | Oct | Nov | Dec |
|---------------|------|------|------|------|------|------|------|------|------|------|------|------|
| Bombay | 1.65 | 1.73 | 1.91 | 1.96 | 2.15 | 1.65 | 1.52 | 1.52 | 1.60 | 1.69 | 1.69 | 1.59 |
| Calcutta | 1.60 | 1.64 | 1.72 | 2.15 | 2.09 | 1.83 | 1.59 | 1.62 | 1.65 | 1.71 | 1.69 | 1.54 |
| Madras | 1.67 | 1.96 | 2.16 | 2.14 | 1.90 | 1.49 | 1.50 | 1.50 | 1.49 | 1.62 | 1.62 | 1.51 |
| Trivandrum | 1.65 | 1.64 | 1.64 | 1.57 | 1.81 | 1.62 | 1.65 | 1.60 | 1.64 | 1.65 | 1.54 | 1.47 |
| Veraval | 1.43 | 1.97 | 2.15 | 2.39 | 3.34 | 2.02 | 1.77 | 1.81 | 1.77 | 1.95 | 1.85 | 1.52 |
| Visakhapatnam | 1.65 | 1.83 | 2.15 | 2.15 | 2.01 | 1.83 | 1.49 | 1.45 | 1.62 | 1.62 | 1.50 | 1.51 |

TABLE 4
Radio climatology at 900 mb over India (1957-1960)

| Station | Time (GMT) | Jan | Feb | Mar | Apr | May | Jun | Jul | Aug | Sep | Oct | Nov | Dec |
|---------------|------------|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|-----|
| Allahabad | 00 | 275 | 265 | 264 | 259 | 264 | 286 | 340 | 341 | 321 | 298 | 270 | 270 |
| | 12 | 273 | 266 | 261 | 258 | 263 | 292 | 326 | 338 | 324 | 290 | 268 | 260 |
| Bombay | 00 | 275 | 272 | 274 | 278 | 304 | 329 | 329 | 326 | 326 | 306 | 290 | 284 |
| | 12 | 277 | 275 | 276 | 280 | 290 | 320 | 334 | 330 | 322 | 304 | 289 | 286 |
| Calcutta | 00 | 278 | 280 | 286 | 286 | 308 | 327 | 335 | 335 | 335 | 321 | 292 | 286 |
| | 12 | 280 | 281 | 283 | 296 | 314 | 317 | 338 | 332 | 333 | 317 | 294 | 288 |
| Gauhati | 00 | 287 | 286 | 287 | 310 | 314 | 328 | 334 | 335 | 340 | 318 | 312 | 295 |
| | 12 | 290 | 289 | 293 | 316 | 332 | 344 | 340 | 340 | 338 | 324 | 304 | 296 |
| Jodhpur | 00 | 268 | 259 | 261 | 260 | 272 | 311 | 330 | 326 | 316 | 283 | 270 | 270 |
| | 12 | 269 | 252 | 260 | 258 | 271 | 300 | 324 | 320 | 306 | 282 | 265 | 268 |
| Madras | 00 | 302 | 291 | 279 | 294 | 310 | 316 | 313 | 313 | 318 | 330 | 313 | 309 |
| | 12 | 296 | 282 | 280 | 296 | 310 | 326 | 320 | 324 | 324 | 320 | 316 | 312 |
| Nagpur | 00 | 281 | 272 | 269 | 273 | 275 | 304 | 328 | 327 | 322 | 304 | 284 | 280 |
| | 12 | 282 | 268 | 265 | 268 | 268 | 302 | 324 | 332 | 319 | 300 | 284 | 278 |
| New Delhi | 00 | 273 | 268 | 270 | 266 | 270 | 293 | 333 | 334 | 320 | 288 | 268 | 282 |
| | 12 | 271 | 265 | 270 | 262 | 265 | 290 | 331 | 339 | 315 | 286 | 272 | 273 |
| Port Blair | 00 | 318 | 322 | 306 | 312 | 336 | 336 | 322 | 329 | 324 | 325 | 321 | 322 |
| | 12 | 322 | 322 | 322 | 336 | 335 | 336 | 320 | 330 | 330 | 331 | 327 | 328 |
| Trivandrum | 00 | 300 | 303 | 305 | 314 | 316 | 320 | 316 | 316 | 316 | 316 | 316 | 316 |
| | 12 | 308 | 304 | 312 | 323 | 322 | 325 | 314 | 315 | 314 | 318 | 319 | 314 |
| Veraval | 00 | 273 | 284 | 273 | 286 | 282 | 318 | 323 | 324 | 314 | 296 | 280 | 279 |
| | 12 | 280 | 270 | 274 | 280 | 282 | 315 | 330 | 322 | 318 | 296 | 284 | 281 |
| Visakhapatnam | 00 | 293 | 288 | 284 | 301 | 309 | 319 | 322 | 331 | 322 | 318 | 303 | 293 |
| | 12 | 292 | 285 | 284 | 299 | 310 | 315 | 325 | 326 | 322 | 316 | 302 | 296 |

Diurnal variation of ΔN at coastal stations — It may be seen from Fig. 2 that in general, ΔN falls in the afternoon as compared to that in the morning at all the coastal stations over the country, except over Veraval and Visakhapatnam where ΔN varies in the reverse order. The diurnal changes from month to month do not vary uniformly.

Monthly variation of ΔN at coastal stations — In general, ΔN is maximum during the summer and decreases in the remaining months of the year. The largest values of ΔN are found at Veraval. The well marked difference in ΔN between Veraval and Bombay is partly attributed to the differences in the performance characteristics of C and F-type radiosondes*. The larger gradient observed at Veraval, north of Bombay in summer indicates predominant ducting conditions over north of Arabian Sea as compared to south. This has been supported by Srivastava and Mehrotra (see Ref.), who have found that M-inversions are more frequent at Bombay as compared to Trivandrum or Cochin.

The value of ΔN is lowest at Madras during July to September which is in conformity with the climatic and geographical considerations. However, the monthly variation of ΔN does not show a significant decrease at all the coastal stations during southwest monsoon when the incidence of super-refracting conditions is found to be minimum. This indicates that ducting conditions for such stations will have to be studied in the light of stability parameters, eddy diffusivity and prevailing meteorological conditions over the oceanic areas for which instruments for measuring turbulence parameters in the lower layers of atmosphere and fine structure of radio refractive index as measured by microwave refractometers may be needed.

Correlation between ΔN versus N_s and regression equations — The seasonal variation of ΔN versus N_s as based on radio-climatic data is given in Table 2. It may be seen that a poor CC (0.40) between the surface value of refractive index and that in the lowest 1 km during southwest monsoon based on one year data may not be truly representative. It appears that the low value of CC in that year occurred partly due to lack of monsoon rainfall and consequent decrease of vapour pressure over the country.

The regression equations between ΔN vs N_s based on one year data for four representative

months recommended by CCIR as reported earlier (Srivastava and Chatterjee 1967) get modified as follows—

| | | |
|------------------------------|-----|-------|
| $\Delta N=0.4104 N_s-80.89$ | Feb | } (5) |
| $\Delta N=0.2870 N_s-42.32$ | May | |
| $\Delta N=0.6665 N_s-200.94$ | Aug | |
| $\Delta N=0.2861 N_s-45.28$ | Nov | |

Keeping in view the limitations of the correlation between N_s and ΔN over the inland stations of the country, it was desired to exclude them and study the CC between the quantities with reference to coastal stations of the country. It was rather surprising to find that except during southwest monsoon, the CC ($\Delta N, N_s$) was quite poor. It is thus difficult to estimate ΔN from N_s to a reliable degree from the regression equations for the coastal stations alone as given below —

| | | |
|------------------------------|--------------|-------|
| $\Delta N=0.083 N_s+88.30$ | Winter | } (6) |
| $\Delta N.N_s=0.24$ | | |
| $\Delta N=0.1645 N_s+15.32$ | Summer | |
| $\Delta N.N_s=0.30$ | | |
| $\Delta N=0.6776 N_s-200.53$ | SW monsoon | |
| $\Delta N.N_s=0.66$ | | } (6) |
| $\Delta N=0.1440 N_s-104.29$ | Post monsoon | |
| $\Delta N.N_s=0.24$ | | |

Effective earth's radius model over coastal stations — The values of effective earth's radius factor k which offer considerable simplification for radio propagation problems and prediction of ducting conditions over oceanic areas are given in Table 3. The results are self explanatory.

Exponential model over coastal stations —

| | | |
|---------------------------|--------------|-------|
| $N(h)=N_s \exp(-0.133 h)$ | Winter | } (7) |
| $N(h)=N_s \exp(-0.135 h)$ | Summer | |
| $N(h)=N_s \exp(-0.140 h)$ | SW monsoon | |
| $N(h)=N_s \exp(-0.136 h)$ | Post monsoon | |

This model has been found to hold good up to lowest 3 km as against the effective earth's radius model which holds good in general only up to 1 km above the ground surface.

Radio climatology at 900-mb level — The radio climatology and its diurnal variation are given in Table 4. It is seen that N_{900} is highest during

*On the performance characteristics of C and F-type Radiosondes Pt. I — Systematic C/F differences by R. Ananthkrishnan, R. Y. Mokashi and A. R. Ramakrishnan, Pre-publ. Sci. Rep. No. 21, India met. Dep., 1968

southwest monsoon and decreases in the other seasons. It is rather difficult to draw any definite conclusion regarding its diurnal variation from morning to evening hours. In general, however, the value of N_{900} falls in the afternoon as com-

pared to that in the morning. It is also worth mentioning that the diurnal variation is quite small being of the order of 2 to 8 N units as compared to N_s when the range may be of the order of 25 to 30 N units at inland stations.

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