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# Homogenizing monthly rainfall and temperature data series in Maharashtra & Goa

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सार – महाराष्ट्र और गोवा के सभी जलवाय केंद्रों की वार्षिक वर्षा और तापमान डेटा श्रृंखला का डेटा एकरूपता के लिए सांख्यिकीय रूप से परीक्षण किया जाता है। किसी स्टेशन की एकरूपता का निरीक्षण करने के लिए, दो-चरणीय दृष्टिकोण का पालन किया जाता है। सबसे पहले, चार समरूपता परीक्षण मानक सामान्य समरूपता परीक्षण, पेटिट का परीक्षण, ब्यूशैंड की रेंज परीक्षण और 5% महत्व के स्तर पर वॉन न्यूमैन राशन परीक्षण का उपयोग वार्षिक वर्षा और तापमान के परीक्षण प्राचलों में एकरूपता के लिए परीक्षण परिकल्पना निर्धारित करने के लिए किया जाता है। दूसरा, इन सभी चार परीक्षणों के परिणामों को 'उपयोगी', 'संदेहात्मक' और 'संदिग्ध' के रूप में तीन अलग-अलग श्रेणियों में एकत्रित किया गया। यहां 30 वर्षा, 29 अधिकतम और न्यूनतम तापमान जलवाय् स्टेशनों का परीक्षण किया गया। प्राप्त परिणामों ने 80% स्टेशनों को 'उपयोगी', 7% को 'संदिग्ध' और 13% स्टेशनों को वर्षा के लिए 'संदेहात्मक' के रूप में दिखाया है, अधिकतम तापमान श्रृंखला के लिए ये परिणाम 17% को 'उपयोगी', 7% को 'संदिग्ध' और 76% को दर्शाते हैं। 'संदिग्ध' के रूप में, जबकि न्यूनतम तापमान शृंखला के लिए ये परिणाम 21% 'उपयोगी' , 10% 'संदिग्ध' और 69% 'संदेहात्मक' हैं। इसके अलावा, इस अध्ययन में एकरूपता के लिए मासिक वर्षा और तापमान डेटा श्रृंखला को सही करने का भी प्रयास किया गया है। ंउपयोगी के रूप में वर्गीकृत स्टेशनों को संदिग्ध और संदेहात्मक स्टेशनों से असमानताओं को दूर करने के लिए संदर्भ श्रृंखला के रूप में उपयोग किया जाता है। वर्षा को सही करने के लिए श्रृंखला अनुपात विधि का उपयोग किया जाता है जबकि तापमान के लिए श्रृंखला जोड़ विधि का उपयोग किया जाता है। सुधार परिणामों ने संदिग्ध श्रेणी के स्टेशनों में महत्वपूर्ण सुधार को दिखाया गया है। असमान श्रंखला के सुधार के बाद, परिणाम दर्शाते हैं कि सभी 100% वर्षा स्टेशन और 65% से अधिक तापमान मापी स्टेशन अब 'उपयोगी' श्रेणी में हैं। संशोधित स्टेशनों को आगे के जलवाय् अनुसंधान अध्ययनों में शामिल किया जा सकता है।

ABSTRACT. Annual rainfall and temperature data series of all climate stations in Maharashtra & Goa are statistically tested for data homogeneity. To inspect homogeneity of a station, a two-step approach is followed. First, four homogeneity tests Standard normal homogeneity test, Pettit's test, Buishand's range test and Von Neumann ration test at 5% level of significance are used to determine test hypothesis for homogeneity on testing parameters of annual rainfall and temperature. Second, results from all these four tests aggregated together into three different classes as 'useful', 'doubtful' and 'suspect'. Here 30 rainfall, 29 maximum and minimum temperature climate stations were tested. The results showed 80% stations as 'useful', 7% as 'suspect' and 13% as 'doubtful' for rainfall, for maximum temperature series these results are 17% as 'useful', 7% as 'suspect' and 76% as 'doubtful', while for minimum temperature series these results are 21% as 'useful', 10% as 'suspect' and 69% as 'doubtful'. Further, in this study an attempt is also made to correct the monthly rainfall and temperature data series for homogeneity. Stations categorised as 'useful' are used as reference series to remove in homogeneities from 'suspect' and 'doubtful' stations. To correct rainfall series ratio's method is used while for temperature series addition method is used. Correction results showed significant improvement in 'suspect' category stations. After correction of inhomogeneous series, the results show all 100% of rainfall stations and more than 65% of temperature stations are now in 'useful' category. The corrected stations may be included in further climate research studies.

**Key words** – Homogeneity tests, Standard normal homogeneity test, Pettit's tests, Buishand's range test, Von neumann ration test, Rainfall series, Temperature series.

#### 1. Introduction

Data homogeneity is very important part of data quality control and historical data archival. Homogenizing

a climate data series is the procedure for removing nonclimatic changes from respective time series. A data series contains breaks or shifts in time series, which are result of change in climate itself or resulted from non-climatic

 $\label{table 1}$  Stations summary & data availability details

						То	tal Num	ber of Yea	rs	_
Index	District	Station	Latitude	Longitude	Period	Including Missing	Exc	luding Mis	ssing	Class of OBSY
						Missing	PRE	MAX	MIN	
42851	Jalgaon	Jalgaon	21.03	75.34	1937-2008	72	70	70	70	II b
42920	Nashik	Ozar	20.08	73.55	1965-2006	42	41	41	41	I
42921	Nashik	Nashik	19.58	73.49	1949-1983	35	31	32	28	I
42925	Nashik	Malegaon	20.33	74.32	1901-2008	108	107	101	107	II b
43075	Osmanabad	Osmanabad	18.07	76.01	1976-2008	33	26	26	24	II b
43001	Thane	Dahanu	19.58	72.43	1944-2017	74	68	67	67	I
43003	Greater Mumbai	Santacruz	19.07	72.51	1951-2017	67	67	67	67	I
43009	Ahmednagar	Ahmednagar	19.05	74.48	1901-2013	113	105	99	100	II b
43011	Beed	Beed	19.00	75.24	1961-1996	36	34	34	34	II b
43013	Aurangabad	Aurangabad	19.53	75.20	1902-1983	82	82	82	82	IIb
43014	Aurangabad	Chikalthana	19.51	75.24	1952-2017	66	66	66	66	I
43021	Nanded	Nanded	19.05	77.20	1960-2007	48	37	37	38	II c
43057	Mumbai	Colaba	18.54	72.48	1901-2017	117	117	117	117	I
43058	Raigad	Alibag	18.38	72.52	1933-2015	83	80	80	80	II b
43061	Pune	Khandala	18.46	73.22	1951-1973	23	21	0	0	Vb
43062	Raigad	Bhira	18.27	73.24	1963-2008	46	43	43	42	II b
43063	Pune	Pune	18.32	73.51	1901-2016	116	116	116	116	I
43069	Pune	Baramati	18.09	74.35	1954-1993	40	40	40	39	II b
43071	Solapur	Jeur	18.12	75.12	1951-2003	53	46	46	45	II b
43109	Ratnagiri	Harnai	17.49	73.06	1943-2014	72	69	69	69	II a
43110	Ratnagiri	Ratnagiri	16.59	73.20	1901-2017	117	111	111	111	I
43111	Satara	Mahabaleshwar	17.56	73.40	1932-2014	83	79	77	77	I
43113	Satara	Satara	17.42	74.01	1976-2015	40	38	38	38	II a
43117	Solapur	Solapur	17.40	75.54	1901-2017	117	117	117	117	I
43153	Sindhudurg	Devgad	16.23	73.21	1945-2006	62	49	49	49	II b
43157	Kolhapur	Kolhapur	16.42	74.14	1946-2016	71	69	68	69	II a
43158	Sangli	Sangli	16.51	74.36	1932-2015	84	69	69	69	II a
43192	Goa	Panjim	15.29	73.49	1964-2016	53	53	53	53	I
43193	Sindhudurg	Vegurla	15.52	73.38	1949-2008	60	51	52	50	ΠЬ
43196	Goa	Mormugao	15.25	73.47	1964-2007	44	44	44	44	I

jumps due to relocation or changes in recording instrumentation, urban heat island effects, changes in observing methods and surroundings of instrument etc. Various studies have shown that sometimes inhomogeneities can cause biased trends in raw data and inhomogeneities in historical data will affect the outcome of data analysis and forecasts. Therefore, homogenizing climate data is necessary in order to obtain reliable regional or global trends, which in result will improve the quality of forecasting. Time series quality control and homogenization is a part of World Meteorological Organization (WMO) guidelines for Global Framework for Climate Services (GFCS) & Climate Services Information System (CSIS), which is to be implemented by National Hydrological & Meteorological Services (NHMS) & Regional Climate Centre (RCC). Here, the concept is to test whether a given data is said to be homogenous over time. In other words, if there is a significant break in trend of particular time series it is classified as inhomogeneous. To identify these breaks in trend one should use metadata to detect any climatic breaks in time series. Upon unavailability of metadata, to identify these breaks there several statistical methods are available but very few methods are available to remove these breaks.

A study on European climate data by Wijngaard et al., (2003) shows a detailed approach for testing homogeneity using four different statistical tests namely standard normal homogeneity test, Pettit's tests, Buishand's range test and Von Neumann ration test. These statistical tests were applied on various testing parameters of temperature and rainfall and based on the number of tests rejecting the null hypothesis the results were categorised into three classes, as useful, doubtful and suspect. There is no study in India so far to test the data homogeneity let alone an attempt to correct it. Monthly rainfall and temperature data from all climate stations in state of Maharashtra and Goa is used in this study, which is provided by India Meteorological Department for all available period since 1901. In order to test data homogeneity, we have adopted the Wijngaard et al., (2003) approach. Annual rainfall amount and annual mean temperature are used as testing variables in this study. Stations with significant inhomogeneities are needed to be removed and adjustment for homogeneity should be made. Homogeneity correction methods for temperature and rainfall are different in approach. Hanssen and Forland (1994) had shown a method to correct inhomogeneous rainfall data series using a reference station and Morozova and Valente (2012) constructed a procedure to correct temperature data series, both of which we have implemented to correct monthly rainfall and temperature data series for all climate stations in state of Maharashtra and Goa.

#### 2. Data and methodology

Monthly rainfall and temperature data from all climate stations in states of Maharashtra and Goa is collected from India Meteorological Department (IMD) Pune. Table 1 shows stations with their geographical coordinates with data availability statistics. Stations with higher percentage of missing data are not considered for this study. Annual mean temperature and rainfall amount are used as testing variables.

#### 2.1. Statistical methods

Firstly, original series is visually analysed for detection of any possible breaks such as outliers, which could arrive due to typing or OCR errors. Further, all four homogeneity tests are applied to the data series. The break dates from homogeneity tests are compared with metadata to identify climatic and non-climatic jumps or breaks in series. It is possible that metadata could not provide all the information regarding changes in stations, observation procedure, changes in surroundings etc. and in this study, we could not find any significant breaks which are associated with metadata or any climate forcing data records. After the testing procedure is over, results from testing are collected and stations are categorised to be selected for correction procedure. For a station, break dates that are common for most tests are selected for correction and in case of different break dates for same station, the later one is selected for correction. For correction of data series with significant inhomogeneities, different methods are used for rainfall and temperature, respectively. Rainfall has more variability throughout the year as compared to temperature, because of that; a use of reference station is must for rainfall series correction. An adjustment factor is computed for each station using the method suggested by Hanssen and Forland (1994) and Morozova and Valente (2012) for rainfall and temperature, respectively. Accordingly, adjustment factor is then used to correct rainfall series by ratio's method and temperature series by additive method. Corrected stations again undergo the homogeneity testing procedure to check for any remaining inhomogeneities and same procedure repeated until most significant non-climatic breaks are removed.

# 2.2. Homogeneity testing

Four statistical tests are used to test the homogeneity of rainfall and temperature data. Standard normal homogeneity test (SNHT), Pettit's test (PT), Buishand's test (BT) and Von Neumann ratio test (VNR) are used to test the data homogeneity. There are several studies in the past by González-Rouco *et al.*, (2001), Tuomenvirta (2001), Kang and Yusof (2012), Ahmad and Deni (2013),

TABLE 2  $1\% \ and \ 5\% \ critical \ values \ for \ \it{X}_m \ of \ the \ Pettit's \ test \ as \ a \ function \\ of \ \it{n}; \ values \ are \ based \ on \ simulation.$ 

n	20	30	40	50	70	100
1%	71	133	208	293	488	841
5%	57	107	167	235	393	677

Taxak et al., (2014), Guhathakurta et al., (2015), Agha et al., (2017) and Byakatonda et al., (2018) where these tests are used for testing homogeneity of data series. These tests have different sensitivities in different parts of series based on their characteristics as parametric, nonparametric and maximum likelihood, which would help in obtaining results that are more significant. The null hypothesis for these tests is that annual values  $Y_i$  of the testing variable Y are independent and identically distributed while the alternate hypothesis for SNHT, PT and BT assume that there is a shift in the mean is present, this significant break in series is considered as inhomogeneous. These tests will identify the year where the break is significant which would be the year where inhomogeneity first occurred. On the other hand, VNR test assume that series is not randomly distributed under alternate hypothesis and the VNR tests does not provide information about a year where break occurs.

The SNHT easily detects breaks near the beginning and at the end of series, whereas to detect breaks in the middle of the series, PT and BR tests are more sensitive (Hawkins, 1977). Pettit's test is based on ranks of the elements in the series, which makes it less sensitive to outliers than other tests like SNHT and Buishand's test, which assume that  $Y_i$  values are normally distributed. The VNR test has sensitivity to departures of homogeneity having characteristics other than stepwise shifts. Annual rainfall amount and mean temperature are testing variables given that  $Y_i$ , (i = 1, 2, ..., n) is the testing variable with  $\bar{Y}$  is the mean and s is the standard deviation.

#### 2.2.1. Standard normal homogeneity test

SNHT is the most referred homogeneity test in climate studies. A test statistic T(d) is used to compare the mean of the first y years with the last of (n - y) years can be written as,

$$T_d = d \bar{z}_1^2 + (n - d) \bar{z}_2^2, d = 1, 2, ..., n$$

where,

$$\overline{z}_1 = \frac{1}{d} \sum_{i=1}^{d} \frac{(Y_i - \overline{Y})}{s} \text{ and } \overline{z}_2 = \frac{1}{n-d} \sum_{i=d+1}^{d} \frac{(Y_i - \overline{Y})}{s}$$

TABLE 3

1% critical values for the statistic  $T_0$  of the single shift SNHT as a function of n [calculated from simulations carried out by Jaruskova (1996) and the 5% critical value (Alexandersson and Moberg, 1997)

n	20	30	40	50	70	100
1%	9.56	10.45	11.01	11.38	11.89	12.32
5%	6.95	7.65	8.10	8.45	8.80	9.15

 $\bar{z}_1$  and  $\bar{z}_2$  are the mean values of  $z_i$  during the first d years and last (n-d) years, respectively. If a break is located at year D then T(d) reaches a maximum near the year d = D. The test statistic  $T_0$  is defines as

$$T_0 = \max_{1 \le d \le n} T(d)$$

The probability of rejecting the null hypothesis when  $T_0$  exceeds a certain critical value, which depends on sample size. Critical values for  $T_0$  are given in Table 3 (Khaliq and Quarda, 2007).

## 2.2.2. Buishand's range test

The homogeneity test can be based on the cumulative deviations from the mean or adjusted partial sum which are define as,

$$S^* = 0$$
 and  $S_d^* = 2\sum_{i=1}^d (Y_i - \overline{Y}), d = 1, 2, ..., n$ 

For a homogeneous series, there is no systematic deviations of the  $Y_i$  values with respect to their mean will appear, therefore the values of  $S_d^*$  will fluctuate around zero. Contrarily if a break is present at year D, then  $S_d^*$  reaches a maximum (negative shift) or minimum (positive shift) near the year d = D. Rescaled adjusted partial sums are obtained by dividing the values of  $S_d^*$  by the sample standard deviation s. The values are not influenced by any linear transformation; therefore, it is suitable to use the homogeneity test by Q statistics (Buishand, 1982).

$$Q = \max_{0 \le d \le n} \left| \frac{S_d^*}{s} \right|$$

In addition, the difference between the maximum and minimum value of the rescaled adjusted partial sums can also be computed by range statistics or R statistics.

$$R = \frac{\left(\max_{0 \le d \le n} S_d^* - \min_{0 \le d \le n} S_d^*\right)}{s}$$

TABLE 4 1% and 5% critical values for Q  $\sqrt{n}$  of the Buishand's range test as a function of n (Buishand, 1982); the value of n=70 is simulated

n	20	30	40	50	70	100
1%	1.60	1.70	1.74	1.78	1.81	1.86
5%	1.43	1.50	1.53	1.55	1.59	1.62

TABLE 5  $1\% \ and \ 5\% \ critical \ values \ for \ R \ \sqrt{n} \ of \ the \ Buishand's \ range \ test}$  as a function of n (Buishand, 1982)

n	10	20	30	40	50	100
1%	1.29	1.42	1.46	1.50	1.52	1.55
5%	1.14	1.22	1.24	1.26	1.27	1.29

Critical values table by Buishand for both homogeneity tests  $Q\sqrt{n}$  and  $R\sqrt{n}$  are given in Tables 4&5 respectively.

#### 2.2.3. Pettit's test

This test is non-parametric rank test, which does not require any assumption for normality. The test is based on the ranking order and the ranks  $r_1$ ,  $r_2$ ,...  $r_n$  of the  $Y_1$ ,  $Y_2$  ...  $Y_n$  are used to calculate the statistics (*Pettit*, 1979).

$$X_d = 2\sum_{i=1}^d r_i - d(n+1) \ d = 1, 2, ..., n$$

If the break occurs in year m, then the statistics is maximal or minimal near the year d = m.

$$X_m = \max_{1 \le d \le n} |X_d|$$

Critical values for  $X_{\rm m}$  are given in Table 2.

#### 2.2.4. Von Neumann's ratio test

The Von Neumann ratio is defined as the ratio of the mean square successive difference between the years to the variance (von Neumann 1941). The test statistic is given as,

$$N = \frac{\sum_{i=1}^{n-1} (Y_i - Y_{i+1})^2}{\sum_{i=1}^{n} (Y_i - \overline{Y})^2}$$

TABLE 6

1% and 5% critical values of N of the Von Neumann ratio test as a function of n. For N  $\leq$  50 these values are taken from Owen (1962); for N=70 and N=100 the critical values are based on asymptotic normal distribution of N (Buishand, 1981)

n	20	30	40	50	70	100
1%	1.04	1.20	1.29	1.36	1.45	1.54
5%	1.30	1.42	1.49	1.54	1.61	1.67

This statistic is often used to test the independence of the random variables  $Y_{i}$ , which are assumed to be successive observations on a stationary Gaussian time series, and sometimes used to test that stationarity of the mean for uncorrelated time. However, this statistic can also be applied in detecting the inhomogeneous series. The series can be considered as homogeneous series if expected value, N = 2. However, for inhomogeneous series, the value of N tends to be lower than 2 (Buishand, 1982). Values greater than 2 implies that the series has rapid variations or oscillations in the mean (Suhaila *et al.*, 2008). This test gives no information about the location of the shift. Table 6 gives critical values for N.

#### 2.3. Critical values table

The homogeneity tests are applied to the rainfall and temperature series at 5% level of significance. Critical values for all four tests (Wijngaard *et al.*, 2003) for rainfall, maximum temperature & minimum temperature are given in Tables 7, 8&9 respectively, followed by critical values after homogeneity correction in Tables 10, 11 & 12.

# 2.4. Classification of homogeneity tests

Classification is the second part of homogeneity tests where results from all these four tests are evaluated and combined together. Classification is made into three different categories based on how many tests rejected the null hypothesis (Schönwiese *Rapp*, 1997). These categories are Useful, Doubtful and Suspect and these can be interpreted as follows,

Class A: Useful - The series that rejects one or none null hypothesis under four tests at 5% level of significance are considered. The stations under this class can be considered as homogeneous and can be used for further analysis.

Class B: Doubtful - The series that rejects two null hypothesis under four tests at 5% level of significance are considered. The stations under this class have

TABLE 7  $\label{eq:table_to_table}$  Homogeneity test statistics (p value approximation) for rainfall

		Pettitt	's Test			SNI	НТ		I	Buishan	ıd's Tes	st		VNR		
Station	p	K	Year	Trend	p	$T_0$	Year	Trend	p	Q	Year	Trend	p	N	Trend	- Class
Aurangabad	0.2243	421	1930	Н0	0.2499	5.79	1925	Н0	0.1066	10.29	1930	Н0	0.0167	1.54	На	Useful
Ahmednagar	0.2754	583	1962	Н0	0.6075	3.62	1962	H0	0.2715	9.63	1962	H0	0.003	1.48	Ha	Useful
Chikalthana	0.0098	475	1986	Ha	0.0701	9.18	1986	Н0	0.0099	12.38	1986	Ha	0.2265	1.82	Н0	Doubtful
Beed	0.7107	71	1986	Н0	0.7472	2.47	1986	H0	0.5698	4.11	1986	H0	0.3899	1.91	H0	Useful
Dahanu	0.0779	393	1964	Н0	0.0828	8.42	1964	H0	0.0338	11.14	1964	Ha	0.057	1.63	Н0	Useful
Mormugao	0.6863	109	2000	H0	0.7403	2.77	2000	H0	0.7378	4.08	2000	H0	0.0803	1.59	H0	Useful
Panjim	0.4772	174	1993	Н0	0.8637	2.14	1994	H0	0.5533	5.32	1993	Н0	0.2246	1.80	Н0	Useful
Jalgaon	0.4426	272	1970	Н0	0.1965	5.93	2003	H0	0.459	6.61	1970	H0	0.2414	1.83	H0	Useful
Kolhapur	0.0933	386	1967	Н0	0.207	6.27	1967	Н0	0.0886	9.76	1967	Н0	0.0343	1.56	Ha	Useful
Malegaon	0.8689	354	1926	Н0	0.3798	4.87	2004	H0	0.8568	5.80	1926	H0	0.109	1.76	H0	Useful
Colaba	0.0012	1390	1952	На	0.0013	15.83	1941	Ha	0.0003	21.23	1952	Ha	0.1356	1.80	Н0	Suspect
Santacruz	0.312	286	2004	Н0	0.2423	5.77	1959	H0	0.2649	7.70	2004	Н0	0.1496	1.75	H0	Useful
Nanded	0.7452	78	1963	Н0	0.2832	4.92	1963	H0	0.5864	4.25	1963	H0	0.7545	2.22	Н0	Useful
Nashik	0.8826	50	1964	Н0	0.8083	2.15	1952	H0	0.7915	3.19	1957	H0	0.4861	1.98	Н0	Useful
Ozar	0.0153	222	1996	Ha	<0.0001	19.89	2002	Ha	0.0093	9.59	1997	На	0.0016	1.05	На	Suspect
Osmanabad	0.4522	59	1997	Н0	0.6879	2.55	1976	H0	0.4542	3.96	1997	Н0	0.9491	2.62	H0	Useful
Pune	0.0089	1170	1972	Ha	0.0075	13.49	2003	На	0.0077	17.59	1972	На	0.0006	1.44	Ha	Suspect
Baramati	0.2152	143	1972	Н0	0.5195	3.76	1972	H0	0.2113	6.20	1972	H0	0.1875	1.72	Н0	Useful
Khandala	0.1263	58	1963	Н0	0.0448	7.16	1971	Ha	0.0789	5.20	1963	Н0	0.074	1.39	Н0	Useful
Alibag	0.1427	452	1952	Н0	0.3359	5.32	1952	Н0	0.1982	8.99	1952	Н0	0.5054	2.00	Н0	Useful
Bhira	0.4498	128	1980	Н0	0.7807	2.22	2003	Н0	0.8286	3.73	1980	Н0	0.0064	1.32	Ha	Useful
Ratnagiri	0.0003	1426	1952	Ha	0.0004	19.15	1930	Ha	0.0004	21.94	1952	Ha	0.0198	1.61	Ha	Suspect
Harnai	0.0751	402	1964	Н0	0.12	6.92	1964	Н0	0.0655	10.26	1964	Н0	0.294	1.87	Н0	Useful
Sangli	0.9006	171	1965	Н0	0.2578	5.64	1932	Н0	0.905	4.23	2004	Н0	0.175	1.78	Н0	Useful
Satara	0.3907	112	1981	Н0	0.3851	4.30	1981	Н0	0.4893	4.73	1981	Н0	0.0591	1.51	Н0	Useful
Mahabaleshwa	0.0562	516	1965	Н0	0.1774	6.27	1963	Н0	0.0643	11.04	1965	Н0	0.0938	1.71	Н0	Useful
Devgad	0.4894	152	1990	Н0	0.4973	3.95	1947	Н0	0.6049	4.87	1989	Н0	0.0674	1.58	Н0	Useful
Vengurla	0.0209	298	1985	Ha	0.0811	7.42	1985	Н0	0.0321	9.67	1985	Ha	0.2993	1.85	Н0	Doubtful
Solapur	0.1939	754	1947	Н0	0.6258	3.63	1947	Н0	0.2825	10.15	1947	Н0	0.4089	1.96	Н0	Useful
Jeur	0.1692	186	1976	Н0	0.2138	5.74	2000	Н0	0.2567	6.32	1976	Н0	0.012	1.34	Ha	Useful

 ${\bf TABLE~8}$  Homogeneity test statistics (p value approximation) for maximum temperature

G:	I	Pettitt's	Test			SNE	IT		Ві	iishand	's Tes	t	,	VNR		CI.
Station	p	K	Year	Trend	p	$T_0$	Year	Trend	p	Q	Year	Trend	p	N	Trend	- Class
Aurangabad	<0.0001	1090	1949	Ha	<0.0001	25.19	1963	Ha	<0.0001	22.12	1949	Ha	<0.0001	1.01	Ha	Suspect
Ahmednagar	<0.0001	1699	1957	На	<0.0001	26.03	1957	Ha	<0.0001	25.22	1957	Ha	<0.0001	0.74	Ha	Suspect
Chikalthana	0.2534	297	2000	Н0	0.3342	4.87	2000	Н0	0.2286	7.90	2000	Н0	0.0063	1.42	Ha	Useful
Beed	0.0142	166	1981	На	0.0437	8.11	1982	Ha	0.0185	8.19	1981	Ha	0.0018	1.05	Ha	Suspect
Dahanu	<0.0001	914	1984	Ha	<0.0001	29.35	1984	Ha	<0.0001	22.04	1984	Ha	<0.0001	0.77	Ha	Suspect
Mormugao	0.0005	326	1976	Ha	<0.0001	28.25	1971	Ha	<0.0001	14.62	1975	Ha	<0.0001	0.39	Ha	Suspect
Panjim	<0.0001	558	2001	Ha	<0.0001	29.93	2001	Ha	<0.0001	18.11	2001	Ha	<0.0001	0.56	Ha	Suspect
Jalgaon	0.0295	470	1948	Ha	0.0496	11.49	1948	Ha	0.0451	10.76	1948	Ha	0.0008	1.24	Ha	Suspect
Kolhapur	0.1315	357	1989	Н0	0.0314	9.65	1956	Ha	0.1047	9.50	1956	Н0	0.0002	1.11	Ha	Doubtful
Malegaon	0.4236	488	1950	Н0	0.0001	20.76	2002	Ha	0.2213	9.98	2002	Н0	<0.0001	1.08	Ha	Doubtful
Colaba	<0.0001	3049	1950	Ha	<0.0001	68.01	1940	Ha	<0.0001	44.05	1950	Ha	<0.0001	0.41	Ha	Suspect
Santacruz	<0.0001	900	1985	Ha	<0.0001	30.05	1985	Ha	<0.0001	22.58	1985	Ha	<0.0001	0.96	Ha	Suspect
Nanded	0.5785	92	1963	Н0	0.0034	16.53	2006	Ha	0.6597	4.07	2006	Н0	0.2122	1.75	Н0	Useful
Nashik	0.001	186	1964	Ha	0.0232	9.04	1964	Ha	0.0033	8.57	1964	Ha	0.0084	1.21	Ha	Suspect
Ozar	0.3983	125	1984	Н0	0.6484	2.85	1965	Н0	0.3892	5.34	1984	Н0	0.7254	2.19	Н0	Useful
Osmanabad	0.0001	153	1985	Ha	<0.0001	17.95	1985	Ha	<0.0001	10.48	1985	Ha	<0.0001	0.56	Ha	Suspect
Pune	0.0958	855	1969	Н0	0.1893	6.08	1904	Н0	0.0755	13.25	1972	Н0	<0.0001	1.28	Ha	Useful
Baramati	0.0022	249	1974	Ha	0.0089	10.73	1975	Ha	0.0017	10.46	1974	Ha	<0.0001	0.86	Н0	Suspect
Alibag	<0.0001	1484	1975	Ha	<0.0001	53.25	1985	Ha	<0.0001	31.87	1975	Ha	<0.0001	0.31	Ha	Suspect
Bhira	0.0881	190	1973	Н0	0.2273	5.49	1973	Н0	0.1308	7.10	1991	Н0	<0.0001	0.58	Ha	Useful
Ratnagiri	<0.0001	2875	1964	Ha	<0.0001	65.42	1964	Ha	<0.0001	42.76	1964	Ha	<0.0001	0.38	Ha	Suspect
Harnai	<0.0001	962	1985	На	<0.0001	29.95	1986	Ha	<0.0001	22.17	1975	Ha	<0.0001	1.00	Ha	Suspect
Sangli	<0.0001	803	1971	Ha	<0.0001	21.13	1971	Ha	<0.0001	19.18	1971	Ha	<0.0001	0.94	Ha	Suspect
Satara	<0.0001	333	1990	Ha	<0.0001	26.91	1990	Ha	<0.0001	15.84	1990	Ha	<0.0001	0.35	Ha	Suspect
Mahabaleshwai	<0.0001	1342	1964	Ha	<0.0001	49.94	1963	Ha	<0.0001	30.76	1963	Ha	<0.0001	0.54	Ha	Suspect
Devgad	<0.0001	460	1975	На	<0.0001	20.91	1975	Ha	<0.0001	15.59	1975	Ha	<0.0001	0.81	Ha	Suspect
Vengurla	<0.0001	596	1970	Ha	<0.0001	26.83	1970	Ha	<0.0001	18.63	1970	Ha	<0.0001	0.82	Ha	Suspect
Solapur	<0.0001	2001	1964	Ha	<0.0001	27.55	2000	Ha	<0.0001	27.48	1964	Ha	<0.0001	0.93	Ha	Suspect
Solapur Jeur	<0.0001	491	1971	Ha	<0.0001	26.75	1971	Ha	<0.0001	17.67	1971	Ha	<0.0001	0.56	Ha	Suspect

 ${\bf TABLE~9}$  Homogeneity test statistics (p value approximation) for minimum temperature

a:	I	Pettitt's	Test			SNH	Т		Ви	iishand	's Tes	t	,	VNR		CI.
Station -	p	K	Year	Trend	p	$T_0$	Year	Trend	p	Q	Year	Trend	p	N	Trend	- Class
Aurangabad	0.0006	805	1917	На	<0.0001	34.79	1910	Ha	<0.0001	19.68	1917	Ha	<0.0001	1.03	Ha	Suspect
Ahmednagar	0.0001	1132	1930	Ha	0.0073	13.75	1930	Ha	0.0019	17.08	1930	Ha	0.0839	1.73	H0	Suspect
Chikalthana	<0.0001	922	1977	Ha	<0.0001	35.36	2001	Ha	<0.0001	22.36	1977	Ha	<0.0001	0.29	Ha	Suspect
Beed	0.0001	233	1975	Ha	<0.0001	15.98	1975	Ha	<0.0001	11.75	1975	Ha	0.0017	1.02	Ha	Suspect
Dahanu	<0.0001	988	1975	Ha	<0.0001	29.83	1975	Ha	<0.0001	22.40	1975	Ha	0.0011	1.25	Ha	Suspect
Mormugao	0.0111	260	1976	Ha	0.0052	13.94	1975	Ha	0.0031	11.16	1975	Ha	0.0143	1.34	Ha	Suspect
Panjim	0.0608	278	1976	Н0	0.037	10.63	1964	Ha	0.0285	9.86	1976	Ha	0.0061	1.32	Ha	Suspect
Jalgaon	0.1754	351	1992	H0	0.0521	8.66	1992	H0	0.0824	9.92	1992	H0	0.0004	1.18	Ha	Useful
Kolhapur	<0.0001	858	1976	Ha	<0.0001	27.09	2008	Ha	<0.0001	17.20	1976	Ha	<0.0001	0.58	Ha	Suspect
Malegaon	0.4704	514	1988	H0	0.0701	10.35	1989	Н0	0.0693	12.58	1966	Н0	<0.0001	0.92	Ha	Useful
Colaba	0.0118	1134	1927	Ha	0.0003	17.98	2008	Ha	0.0161	16.06	1997	Ha	<0.0001	1.17	Ha	Suspect
Santacruz	<0.0001	901	1976	Ha	<0.0001	31.87	1975	Ha	<0.0001	22.57	1976	Ha	<0.0001	0.88	Ha	Suspect
Nanded	0.0173	193	1993	На	0.0815	8.45	1993	Н0	0.0456	7.72	1993	Ha	0.0152	1.35	Ha	Suspect
Nashik	0.158	88	1956	H0	0.2869	4.71	1956	H0	0.0872	5.28	1956	H0	0.1731	1.77	H0	Useful
Ozar	0.0402	198	1999	На	0.0002	17.03	1999	Ha	0.0094	9.46	1999	Ha	0.0393	1.48	Ha	Suspect
Osmanabad	0.0058	106	2002	Ha	0.2049	5.30	2002	H0	0.1398	4.99	2002	H0	0.0041	0.93	Ha	Doubtful
Pune	0.057	930	1962	H0	0.083	8.50	2005	Н0	0.0674	13.44	1962	Н0	<0.0001	1.10	Ha	Useful
Baramati	0.7612	84	1988	H0	0.5235	3.62	1991	H0	0.7023	3.93	1970	Н0	0.1987	1.74	H0	Useful
Alibag	0.0215	592	1994	Ha	0.002	14.22	2001	Ha	0.0161	13.25	1995	Ha	0.0008	1.29	Ha	Suspect
Bhira	0.0005	300	1978	Ha	0.03	13.84	2007	Ha	0.0039	10.25	1975	Ha	<0.0001	0.51	Ha	Suspect
Ratnagiri	0.0008	1271	1960	Ha	0.0007	26.32	2010	Ha	0.0003	19.86	1960	Ha	<0.0001	0.93	Ha	Suspect
Harnai	0.2534	317	1974	Н0	0.0132	13.39	2008	Ha	0.1558	8.82	1974	Н0	0.0021	1.29	Ha	Doubtful
Sangli	0.0494	428	1989	Ha	0.0022	17.69	2014	Ha	0.0607	10.18	1990	Н0	<0.0001	0.95	Ha	Suspect
Satara	0.0116	200	1987	Ha	0.0017	13.27	1982	Ha	0.0075	9.20	1987	Ha	0.0055	1.21	Ha	Suspect
Mahabaleshwar	0.4481	315	1997	H0	0.5813	3.51	1972	H0	0.2533	8.26	1972	Н0	0.1549	1.78	Н0	Useful
Devgad	0.0136	293	1966	Ha	0.0037	14.09	1981	Ha	0.0028	11.48	1980	Ha	<0.0001	0.80	Ha	Suspect
Vengurla	<0.0001	487	1981	Ha	0.0001	19.00	1981	Ha	0.0001	15.55	1981	Ha	0.0023	1.23	Ha	Suspect
Solapur	<0.0001	1928	1956	На	0.0432	15.21	1996	Ha	0.0051	17.35	1956	Ha	<0.0001	0.99	Ha	Suspect
Jeur	0.261	162	1987	Н0	0.0543	11.70	1993	Н0	0.0108	9.48	1987	Ha	0.001	1.02	Ha	Doubtful

 ${\bf TABLE~10}$  Homogeneity Test Statistics (p value approximation) for Rainfall after Correction

Station	I	Pettitt's	s Test			SNI	łΤ		Ві	iishand	l's Tes	t		VNR		Class
Station	p	K	Year	Trend	p	$T_0$	Year	Trend	p	Q	Year	Trend	p	N	Trend	Class
Chikalthana	0.2048	312	1993	НО	0.6473	3.13	1993	Н0	0.3612	6.96	1993	НО	0.7406	2.16	Н0	Useful
Colaba	0.2556	710	1929	Н0	0.1941	6.64	1925	Н0	0.1474	11.83	1941	Н0	0.542	2.02	H0	Useful
Ozar	0.7036	96	2002	Н0	0.2763	4.97	2004	Н0	0.6841	4.15	2002	Н0	0.3929	1.91	H0	Useful
Pune	0.8446	418	2003	Н0	0.3894	4.79	2003	Н0	0.6045	7.72	2003	Н0	0.0172	1.61	H0	Useful
Ratnagiri	0.0397	906	1972	Ha	0.1854	6.57	1972	Н0	0.0603	13.32	1972	Н0	0.2482	1.87	H0	Useful
Vengurla	0.48	162	1964	Н0	0.7046	2.65	1963	Н0	0.5409	5.24	1963	Н0	0.5695	2.05	H0	Useful

 ${\bf TABLE~11}$  Homogeneity test statistics (p value approximation) for maximum temperature after correction

G:	I	Pettitt's	Test			SNH	ΙΤ		Ві	iishand	's Tes	t		VNR		CI
Station -	p	K	Year	Trend	p	$T_0$	Year	Trend	p	Q	Year	Trend	p	N	Trend	- Class
Aurangabad	0.3184	389	1963	Н0	0.3191	5.37	1971	НО	0.3886	7.73	1963	Н0	0.0055	1.45	Ha	Useful
Ahmednagar	0.1703	607	1940	Н0	0.2393	6.25	1981	Н0	0.2427	9.64	1981	H0	<0.0001	1.02	Ha	Useful
Beed	0.487	86	1990	Н0	0.0868	7.15	1961	H0	0.5114	4.30	1990	H0	0.0358	1.41	H0	Useful
Dahanu	0.1444	346	1964	H0	0.2227	6.17	1961	H0	0.1158	9.27	1964	H0	0.005	1.40	Ha	Useful
Mormugao	0.0923	196	1969	Н0	0.0047	11.68	1968	Ha	0.0771	7.87	1969	Н0	<0.0001	0.85	Ha	Doubtful
Panjim	<0.0001	570	2001	Ha	<0.0001	47.20	2001	Ha	<0.0001	22.75	2001	Ha	<0.0001	0.17	Ha	Suspect
Jalgaon	0.1481	366	1988	H0	0.124	7.85	1989	H0	0.0664	10.20	1988	H0	0.0099	1.46	Ha	Useful
Kolhapur	0.0054	521	1989	Ha	0.0267	9.88	1989	Ha	0.0102	12.59	1989	Ha	0.0008	1.21	Ha	Suspect
Colaba	0.0008	1434	1985	Ha	0.0002	19.28	1986	Ha	0.0009	21.05	1986	Ha	<0.0001	0.97	Ha	Suspect
Santacruz	0.2332	309	1999	H0	0.5241	3.84	2001	H0	0.3883	6.89	2001	H0	0.1546	1.75	H0	Useful
Nashik	0.0573	122	1958	Н0	0.11	7.53	1982	Н0	0.3719	4.57	1958	H0	0.0427	1.43	Ha	Useful
Osmanabad	0.443	59	1981	H0	0.1782	5.12	1978	H0	0.4096	4.06	1981	H0	0.1731	1.64	H0	Useful
Baramati	0.1527	154	1963	H0	0.0686	7.54	1990	H0	0.153	6.60	1963	H0	0.0041	1.19	Ha	Useful
Alibag	0.0101	646	1994	Ha	0.0018	14.75	1997	Ha	0.0091	14.12	1994	Ha	<0.0001	0.85	Ha	Suspect
Ratnagiri	0.0019	1244	2000	Ha	<0.0001	32.63	2001	Ha	0.0006	21.30	2000	Ha	<0.0001	0.92	Ha	Suspect
Harnai	0.1782	347	1959	H0	0.13	7.55	1957	H0	0.0859	9.81	1959	H0	0.1567	1.76	H0	Useful
Sangli	0.1704	346	1939	H0	0.0155	11.05	1939	Ha	0.1477	8.90	1939	H0	0.0008	1.27	Ha	Doubtful
Satara	0.0775	161	1999	H0	0.1514	5.86	2008	H0	0.2592	5.68	2006	H0	0.0004	0.99	Ha	Useful
Mahabaleshwar	0.0324	530	1948	Ha	0.0159	11.10	1940	Ha	0.0525	11.14	1948	H0	0.0296	1.58	Ha	Suspect
Devgad	0.3327	174	1993	H0	0.0155	10.00	1993	Ha	0.3388	6.13	1993	H0	0.0129	1.39	Ha	Doubtful
Vengurla	0.0877	254	2000	Н0	0.0155	10.10	2001	Ha	0.1183	7.99	2000	Н0	0.0527	1.56	H0	Useful
Solapur	0.0632	918	2000	H0	0.0109	12.09	2001	Ha	0.0734	13.20	2000	H0	<0.0001	1.19	Ha	Doubtful
Jeur	0.1731	185	1981	Н0	0.2896	5.13	1981	Н0	0.1385	7.28	1981	Н0	0.0053	1.27	Ha	Useful

 ${\bf TABLE~12}$  Homogeneity Test Statistics (p value approximation) for Minimum Temperature after Correction

Station	]	Pettitt's	s Test			SNF	łΤ		Ві	uishand	's Tes	t	,	VNR		Class
Station	p	K	Year	Trend	p	$T_0$	Year	Trend	p	Q	Year	Trend	p	N	Trend	- Class
Aurangabad	0.4013	357	1973	Н0	0.0437	8.72	1980	На	0.4195	7.45	1973	Н0	0.0192	1.55	На	Useful
Ahmednagar	0.1493	632	1981	Н0	0.0293	13.34	2003	Ha	0.4511	8.00	2003	Н0	0.4619	1.98	H0	Useful
Chikalthana	0.0001	616	2001	Ha	<0.0001	29.65	2006	Н0	0.0001	16.97	2005	Ha	<0.0001	0.58	Ha	Suspect
Beed	0.9289	54	1994	Н0	0.3463	4.17	1961	Н0	0.9766	2.38	1994	Н0	0.3164	1.83	H0	Useful
Dahanu	0.3288	281	1968	Н0	0.4173	4.22	2005	Н0	0.6569	5.51	1968	НО	0.8619	2.25	H0	Useful
Mormugao	0.7359	104	1989	Н0	0.1675	5.63	1964	Н0	0.9249	3.19	1987	Н0	0.3711	1.90	H0	Useful
Panjim	0.8276	125	2003	Н0	0.0322	9.41	1964	Ha	0.7295	4.53	1976	Н0	0.0395	1.53	Ha	Doubtful
Kolhapur	0.0532	418	2007	Н0	0.0009	22.16	2008	Ha	0.0189	11.59	2007	Ha	<0.0001	0.76	Ha	Suspect
Colaba	0.0012	1354	1995	Ha	<0.0001	26.87	2000	Ha	0.0004	21.48	1995	Ha	<0.0001	1.30	Ha	Suspect
Santacruz	0.5856	227	2010	Н0	0.1898	6.15	1954	H0	0.6102	5.74	2010	H0	0.0513	1.62	H0	Useful
Nanded	0.3324	117	1993	НО	0.2123	5.70	2006	Н0	0.6209	4.16	1993	Н0	0.1056	1.60	H0	Useful
Ozar	0.4618	118	1975	Н0	0.2064	5.68	1968	Н0	0.2325	6.07	1975	Н0	0.5784	2.06	H0	Useful
Osmanabad	0.2867	60	2002	Н0	0.4379	3.30	1997	Н0	0.2385	4.53	1997	Н0	<0.0001	8.18	Ha	Useful
Alibag	0.7735	251	1982	Н0	0.5231	3.98	2014	Н0	0.6567	6.02	1982	Н0	0.0205	1.55	Ha	Useful
Bhira	0.6026	109	1975	НО	0.0083	21.55	2007	Ha	0.1231	6.89	2005	Н0	<0.0001	0.60	Ha	Doubtful
Ratnagiri	0.3593	594	2010	Н0	0.0013	21.74	2010	Ha	0.112	12.00	2010	Н0	<0.0001	1.06	Ha	Doubtful
Harnai	0.5106	251	1974	Н0	0.607	3.48	1950	Н0	0.4699	6.50	1974	НО	0.0494	1.61	Ha	Useful
Sangli	0.4666	263	1969	Н0	0.0191	15.98	2014	Ha	0.5527	6.08	1969	Н0	<0.0001	1.04	Ha	Doubtful
Satara	0.3116	120	2000	Н0	0.1522	6.54	1980	Н0	0.1979	6.10	1982	Н0	0.069	1.53	Н0	Useful
Devgad	0.3609	168	1985	Н0	0.3742	4.37	2003	НО	0.4739	5.41	1985	Н0	0.0004	1.08	На	Useful
Vengurla	0.6855	132	1978	Н0	0.9053	1.82	1950	Н0	0.8124	4.04	1978	Н0	0.2017	1.77	H0	Useful
Solapur	0.0081	1178	1975	Ha	0.2681	6.14	2000	Н0	0.2526	10.24	1996	Н0	<0.0001	1.08	Ha	Doubtful

inhomogeneities and can be used for further analysis after critical inspection.

Class C: Suspect - The series that rejects three or all null hypothesis under four tests at 5% level of significance are considered. The stations under this class have significant inhomogeneities and cannot be used for further analysis.

In this study stations with class 'doubtful' and 'suspect' are excluded from further climate trend analysis until they are thoroughly inspected for inhomogeneities and necessary corrections are made to adjust the data for homogeneity.

## 2.5. Homogeneity correction

The results from homogeneity tests are analysed and based on categorisation, stations with class 'doubtful' and 'suspect' are selected for correction procedure. Stations with category 'useful' are used as reference station. Adjustment factor is calculated for both rainfall and temperature, which is needed to correct the original data series.

# 2.5.1. Selection of reference station

For selection of reference station, influence of spatial variation and eventual inhomogeneities in the reference

 ${\bf TABLE~13}$  Homogeneity tests before and after correction summary

Index			Class	of Homogeneity	(Before)	Class of Homogeneity (After)			
	District	Station	Rainfall	Maximum Temperature	Minimum Temperature	Rainfall	Maximum Temperature	Minimum Temperature	
42851	Ahmednagar	Ahmednagar	A	С	С		A	A	
43013	Aurangabad	Aurangabad	A	C	C		A	A	
42921	Aurangabad	Chikalthana	В	A	С	A		С	
42925	Beed	Beed	A	С	С		A	A	
43075	Dahanu	Dahanu	A	С	С		A	A	
43001	Goa	Mormugao	A	С	С		В	A	
43003	Goa	Panjim	A	С	С		C	В	
43009	Jalgaon	Jalgaon	A	С	A		A		
43011	Kolhapur	Kolhapur	A	В	С		C	C	
43014	Malegaon	Malegaon	A	В	A				
43021	Mumbai	Colaba	C	С	С	A	С	C	
43057	Mumbai	Santacruz	A	С	С		A	A	
43058	Nanded	Nanded	A	A	С			A	
43060	Nashik	Nashik	A	С	A		A		
42920	Nashik	Ozar	C	A	С	A		A	
43062	Osmanabad	Osmanabad	A	С	В		A	A	
43063	Pune	Pune	C	A	A	A			
43069	Pune	Baramati	A	С	A		A		
43061	Pune	Khandala	A						
43071	Raigad	Alibag	A	С	С		С	A	
43109	Raigad	Bhira	A	A	С			В	
43111	Ratnagiri	Ratnagiri	С	С	С	A	С	В	
43113	Ratnagiri	Harnai	A	С	В		A	A	
43117	Sangli	Sangli	A	С	С		В	В	
43153	Satara	Satara	A	С	С		A	A	
43157	Satara	Mahabaleshwar	A	С	A		С		
43158	Sindhudurg	Devgad	A	С	С		В	A	
43192	Sindhudurg	Vengurla	В	С	С	A	A	A	
43193	Solapur	Solapur	A	C	C		В	В	
43196	Solapur	Jeur	A	С	В		A		

TABLE 14

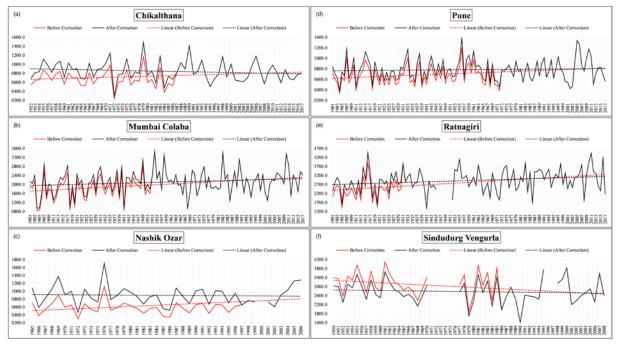
Homogeneity test results

Parameter	Total	Station	Class (Before Co	rrection)	Station Class (After Correction)				
Farameter	Stations	Class A:Useful	Class B:Doubtful	Class C:Suspect	Class A:Useful C	Class B:Doubtfu	l Class C:Suspect		
Precipitation	30	24 (80%)	2 (7%)	4 (13%)	30 (100%)	0 (0%)	0 (0%)		
Maximum Temperature	29	5 (17%)	2 (7%)	22 (76%)	18 (62%)	5 (17%)	6 (21%)		
Minimum Temperature	29	6 (21%)	3 (10%)	20 (69%)	20 (69%)	6 (21%)	3 (10%)		

TABLE 15

Homogeneity Test Regression Statistics Results

Station	Rainfall				Maximum Temperature				Minimum Temperature			
	Before		After		Before		After		Before		After	
	R Square	Slope	R Square	Slope	R Square	Slope	R Square	Slope	R Square	Slope	R Square	Slope
Aurangabad					0.245	0.011	0.001	0.001	0.163	0.008	0.001	0.000
Ahmednagar					0.187	0.013	0.000	0.000	0.038	0.004	0.002	-0.001
Chikalthana	0.054	2.472	0.026	-1.786					0.594	0.041	0.074	0.010
Beed					0.090	-0.026	0.015	0.009	0.386	0.034	0.007	0.003
Dahanu					0.507	0.020	0.023	0.003	0.391	0.019	0.005	0.002
Mormugao					0.548	0.030	0.064	0.007	0.173	0.011	0.001	-0.001
Panjim					0.619	0.023	0.696	0.052	0.064	0.005	0.005	0.001
Jalgaon					0.007	0.002	0.034	-0.005				
Kolhapur					0.003	0.001	0.040	-0.004	0.253	0.015	0.009	0.002
Colaba	0.119	5.310	0.031	2.674	0.714	0.015	0.089	0.004	0.002	0.000	0.105	0.004
Santacruz					0.412	0.014	0.006	0.001	0.349	0.017	0.000	0.000
Nanded									0.088	0.021	0.012	0.007
Nashik					0.155	0.035	0.007	-0.006				
Ozar	0.162	7.277	0.003	-1.035					0.211	0.022	0.022	0.005
Osmanabad					0.617	0.053	0.051	0.008	0.002	-0.005	0.025	-0.017
Pune	0.087	1.655	0.004	0.353								
Baramati					0.048	-0.012	0.084	0.014				
Alibag					0.788	0.024	0.110	0.005	0.074	0.004	0.001	0.000
Bhira									0.078	0.028	0.018	-0.012
Ratnagiri	0.169	7.536	0.050	3.892	0.654	0.014	0.050	0.002	0.114	-0.005	0.009	-0.001
Harnai					0.500	0.018	0.035	0.004	0.041	-0.006	0.002	-0.001
Sangli					0.313	0.014	0.001	0.001	0.041	0.006	0.001	-0.001
Satara					0.714	0.066	0.044	0.009	0.085	-0.016	0.001	-0.001
Mahabaleshwa	r				0.664	0.032	0.066	0.006				
Devgad					0.419	0.018	0.035	0.004	0.248	-0.018	0.001	-0.001
Vengurla	0.103	-10.379	0.008	-2.500	0.551	0.015	0.013	0.002	0.252	-0.012	0.004	0.001
Solapur					0.180	0.007	0.001	0.001	0.082	0.006	0.000	0.000
Jeur					0.521	0.047	0.020	0.006				



Figs. 1(a-f). Rainfall series before and after homogeneity correction

series will be reduced by using more than one reference series. The criteria for selecting reference series stations must satisfy conditions such as, it must be at minimum distance from test station, it should have same climate properties as test station, the availability of data must be of same or more period as test station and it the data of reference series should highly correlate with test series data. Based on our testing results, we have decided to use only single reference station because the aforementioned selection criteria was difficult to achieve for multiple reference station as stations with 'useful' categories were at far distance from each other having different data availability period and difference in climate patterns.

# 2.5.2. Adjustment factor

After identifying the break dates in test series, an adjustment for homogeneity is made by calculating adjustment factor. This adjustment factor is then used to correct the data before the break date by multiplying original rainfall data or by adding in original temperature data. As mentioned earlier, rainfall and temperature series have different correction procedure; the details of which are given below.

#### 2.5.3. Rainfall series correction

Generally, ratio in test and reference series q in a specific year i is given as,

$$q_i = \frac{f(P_i)}{g(Q_{ij}, j = 1, K_i)}$$

Here, i is the function of precipitation  $P_i$  at the test station,  $Q_{ij}$  is the precipitation at the  $j^{th}$  reference station, and g is the function of precipitation at all the  $k_i k_i$  reference stations. Functions of f and g are defined as,

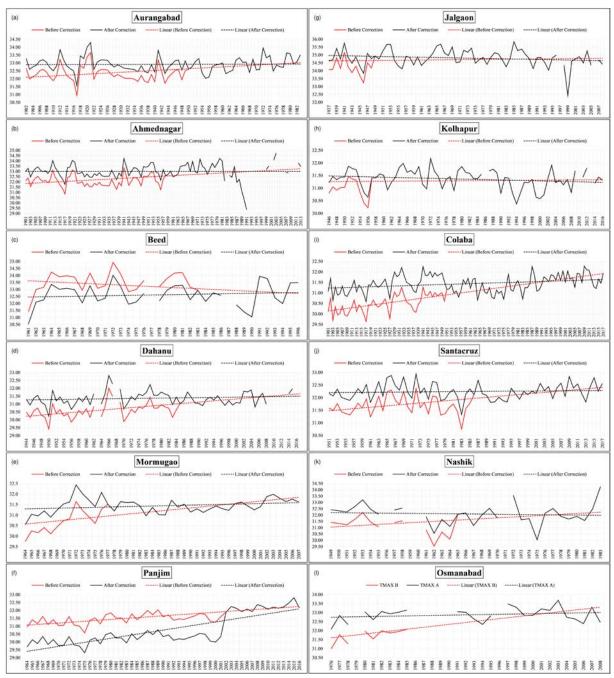
$$f(P_i) = \frac{P_i}{\overline{P}}$$
 and  $g(Q_{ij}, j = 1, K_i) = \frac{\sum_{j=1}^{k_i} V_j(Q_{ij} / \overline{Q}_j)}{\sum_{j=1}^{k_i} V_j}$ 

where,  $\bar{P}$  and  $\bar{Q}_j$  are mean precipitation throughout the observation period for the test station and the  $j^{\text{th}}$  reference station, respectively, and  $k_i$  is the number of reference stations in the ith year of observation.  $V_j$  is a weight factor for reference station j which is defined by the correlation coefficients between the test series and  $j^{\text{th}}$  reference series (Peterson et al., 1998).

Inhomogeneous series is adjusted by multiplying the precipitation values for the period before the break date with an adjustment factor,

$$AF = \frac{q_a}{q_b}$$

Here,  $q_a$  and  $q_b$  are mean values of  $q_i$  after and before break year, respectively.

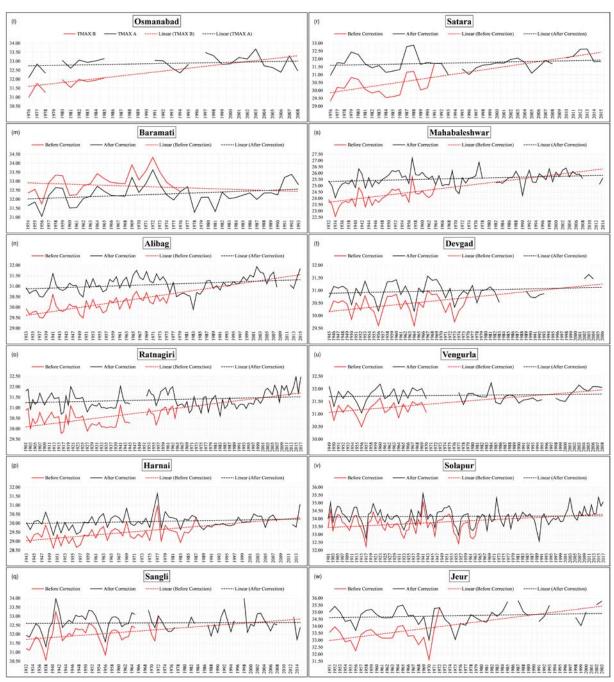


Figs. 2(a-l). Maximum temperature series before and after homogeneity correction

# 2.5.4. Temperature series correction

First, select a time interval  $\Delta t$  before and after break year  $t_{\text{break}}$ . Next, for each interval  $\Delta t$ , take average of each individual month separately and let us call it  $t_{\text{break}}$  and  $t_{\text{after}}$ . Now, calculate the correction dT for each month separately by taking difference of these averages  $t_{\text{before}}$  and  $t_{\text{after}}$ .

After calculating dT, make smoothing 12 monthly dT by taking 3-month adjacent running average to achieve reasonable variation of dT throughout the year. By doing this for each of 12 months individually, we will get our adjustments factor AF for each month. Corrected series can be achieved by adding this adjustment factor in original temperature series for the period before the break year.

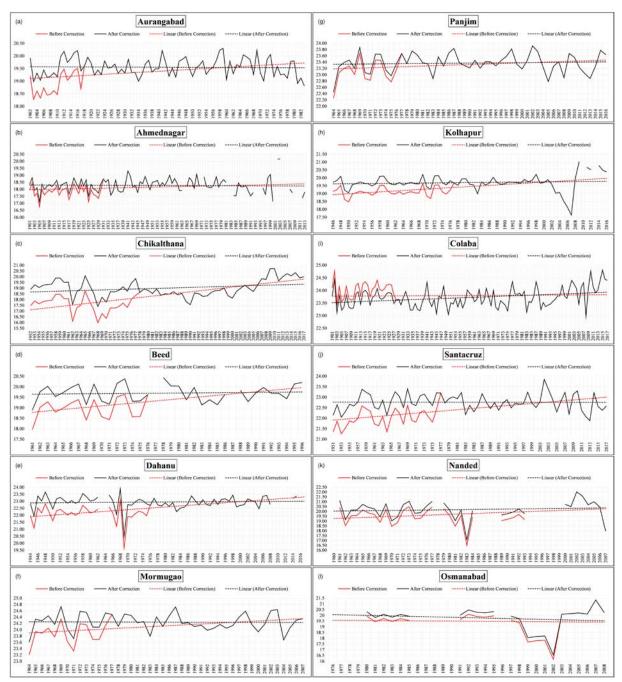


Figs. 3(I-w). Maximum temperature series before and after homogeneity correction

# 3. Results and discussions

Total 30 rainfall and temperature stations were tested for data homogeneity. For rainfall series, homogeneity test results show 80% as useful, 7% as doubtful and 13% as suspect (Table 7), For maximum temperature series, homogeneity test result shows 17% useful, 7% doubtful and 76% suspect stations (Table 8) and for minimum temperature series, homogeneity test result shows 21%

useful, 10% doubtful and 69% suspect stations (Table 9). Use of various test found to be useful in detecting break year because of their characteristics as parametric, non-parametric, rank, *etc*. It is also observed that, break year detected by most tests were same for a respective station. The stations with doubtful and suspect category were taken up for correction by using useful category stations as reference series for rainfall. After applying correction procedures, the results show all 100% rainfall stations as

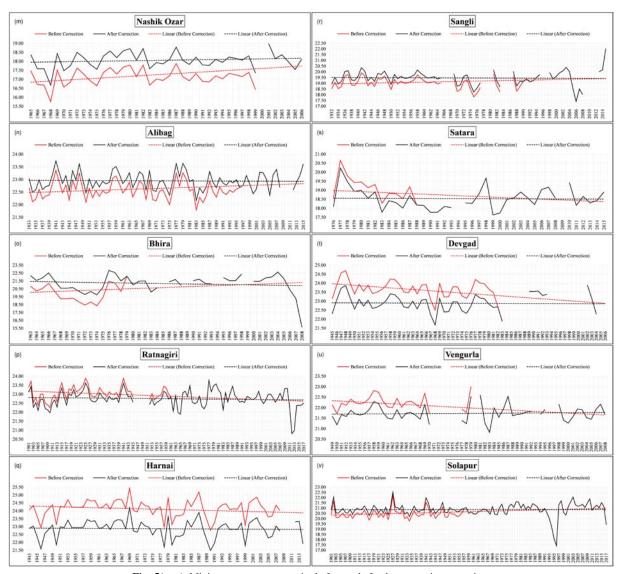


Figs. 4(a-l). Minimum temperature series before and after homogeneity correction

'useful' (Table 10), for maximum temperature series the results are 62% as 'useful', 17% as 'doubtful' and 21% as 'suspect (Table 11), while for minimum temperature series these results are 69% as 'useful', 21% as 'suspect' and 10% as 'doubtful' (Table 12). The 'useful' category stations may be used further in climate and trend analysis studies. Table 13 shows the summary of homogeneity tests before & after correction for all three parameters,

whereas Table 14 shows the consolidated homogeneity test results.

Table 15 shows the linear regression statistics (change in trend) for all three parameters before and after applying the homogeneity correction. Fig. 1 shows the linear trend line for rainfall and it shows that 33% stations show-decreasing trend after applying homogeneity



Figs 5(m-v). Minimum temperature series before and after homogeneity correction

correction. Figs. 2&3 shows the linear trend line for maximum temperature and it shows that 9% of the stations show the increasing trend whereas 13% of the total stations show the decreasing trend after applying homogeneity correction. Figs. 4&5 shows the linear trend line for minimum temperature, which shows that 5% of the stations show the increasing trend whereas 18% of total stations show the decreasing Trend. For majority of stations, it has been seen then there is an overall reduction in magnitude of a trained to applying the homogeneity correction for all the rainfall and temperature series.

One of the biggest consequences of biased trends are reflected in the climate change studies. Several climate studies show the increasing trend in temperature and decreasing trend in rainfall over last century, concluding as an effect of global warming. Such studies have failed to incorporate effects of non-climatic jumps in data and the inhomogeneities arrived due to various external factors are also not considered. Here, the trends in series before and after correction will also help to determine if there is an effect of climate change or not.

#### 4. Conclusions

In the present study, homogeneity tests are proven most important part of climate and trend analysis studies. Different characteristics of these tests not only helped in identifying significant non-climatic jump in a time series but also location of a break period. The correction methods used are proved to be very effective in removing non-climatic jump from time series and making it homogeneous.

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