MAUSAM, **75**, 1 (January 2024), 17-34

# MAUSAM

DOI : https://doi.org/10.54302/mausam.v75i1.5886 Homepage: https://mausamjournal.imd.gov.in/index.php/MAUSAM



UDC No. 551.58 : 551.577 : 551.524 (540)

## **Homogenizing monthly rainfall and temperature data series in Maharashtra & Goa**

NILESH WAGH and PULAK GUHATHAKURTA

*India Meteorological Department, Pune, Maharashtra – 411 005, India* (*Received 27 April 2022, Accepted 16 February 2023*) **e mail : nileshwagh.imd@gmail.com**

**सार** — **महारा� और गोवा के सभी जलवायुक� द् क� वा�षर् वषार और तापमान डेटा श्रृंख का डेटा एकरूपत** के लिए सांख्यिकीय रूप से परीक्षण किया जाता है। किसी स्टेशन की एकरूपता का निरीक्षण करने के लिए , दो-**चरणीय दृ�को का पालन �कया जाता है। सबसे पहले**, **चार समरूपत पर��ण मानक सामान् समरूपत पर��ण**, **पे�टट का पर��ण**, **ब्ूय श क� र�ज पर��ण और** 5% **महत् के स्त पर वॉन न्ूमैय राशन पर��ण का उपयोग वा�षर् वषार और तापमान के पर��ण प्राच म� एकरूपत के �लए पर��ण प�रकल्पन �नधार्र करने के �लए �कया जाता है। दसरा ू** , **इन सभी चार पर��ण के प�रणाम को** '**उपयोगी**', '**संदेहात्म**' **और** '**सं�दग्**' **के रू म� तीन अलग**-**अलग श्रेिण म� एक�त् �कया गया। यहां** 30 **वषार**, 29 **अ�धकतम और न्ूनत यतापमान जलवायु स्टेशन का पर��ण �कया गया। प्र प�रणाम ने** 80% **स्टेशन को** '**उपयोगी**', 7% **को** '**सं�दग्**' **और** 13% **स्टेशन को वषार के �लए** '**संदेहात्म**' **के रू म� �दखाया है**, **अ�धकतम तापमान श्रृंख के �लए ये प�रणाम** 17% **को** '**उपयोगी**', 7% **को** '**सं�दग्**' **और** 76% **को दशार्त ह।** '**सं�दग्**' **के रू म�**, **जब�क न्ूनत य तापमान श्रृंख के �लए ये प�रणाम** 21% '**उपयोगी**', 10% '**सं�दग्**' **और** 69% '**संदेहात्म**' **ह। इसके अलावा**, **इस** अध्ययन में एकरूपता के लिए मासिक वर्षा और तापमान डेटा श्रृंखला को सही करने का भी प्रयास किया गया है। '**उपयोगी**' **के रू म� वग�कृत स्टेशन को** '**सं�दग्**' **और** '**संदेहात्म**' **स्टेशन से असमानताओं को दरू करने के** लिए संदर्भ श्रृंखला के रूप में उपयोग किया जाता है। वर्षा को सही करने के लिए श्रृंखला अनुपात विधि का उपयोग किया जाता है जबकि तापमान के लिए श्रृंखला जोड़ विधि का उपयोग किया जाता है। सुधार परिणामों ने पांदिग्ध श्रेणी के स्टेशनों में महत्वपूर्ण सुधार को दिखाया गया है। असमान श्रृंखला के सुधार के बाद, परिणाम दर्शाते हैं कि **सभी** 100% **वषार स्टेश और** 65% **से अ�धक तापमान मापी स्टेश अब** '**उपयोगी**' **श्रे म� ह। संशो�धत स्टेशन को आगे के जलवायुअनुसधान ं अध्ययन म� शा�मल �कया जा सकता है।**

**ABSTRACT.** Annual rainfall and temperature data series of all climate stations in Maharashtra & Goa are statistically tested for data homogeneity. To inspect homogeneity of a station, a two-step approach is followed. First, four homogeneity tests Standard normal homogeneity test, Pettit's test, Buishand's range test and Von Neumann ration test at 5% level of significance are used to determine test hypothesis for homogeneity on testing parameters of annual rainfall and temperature. Second, results from all these four tests aggregated together into three different classes as 'useful', 'doubtful' and 'suspect'. Here 30 rainfall, 29 maximum and minimum temperature climate stations were tested. The results showed 80% stations as 'useful', 7% as 'suspect' and 13% as 'doubtful' for rainfall, for maximum temperature series these results are 17% as 'useful', 7% as 'suspect' and 76% as 'doubtful', while for minimum temperature series these results are 21% as 'useful', 10% as 'suspect' and 69% as 'doubtful'. Further, in this study an attempt is also made to correct the monthly rainfall and temperature data series for homogeneity. Stations categorised as 'useful' are used as reference series to remove in homogeneities from 'suspect' and 'doubtful' stations. To correct rainfall series ratio's method is used while for temperature series addition method is used. Correction results showed significant improvement in 'suspect' category stations. After correction of inhomogeneous series, the results show all 100% of rainfall stations and more than 65% of temperature stations are now in 'useful' category. The corrected stations may be included in further climate research studies.

**Key words** – Homogeneity tests, Standard normal homogeneity test, Pettit's tests, Buishand's range test, Von neumann ration test, Rainfall series, Temperature series.

#### **1. Introduction**

Data homogeneity is very important part of data quality control and historical data archival. Homogenizing

a climate data series is the procedure for removing nonclimatic changes from respective time series. A data series contains breaks or shifts in time series, which are result of change in climate itself or resulted from non-climatic

## **Stations summary & data availability details**



jumps due to relocation or changes in recording instrumentation, urban heat island effects, changes in observing methods and surroundings of instrument *etc*. Various studies have shown that sometimes inhomogeneities can cause biased trends in raw data and inhomogeneities in historical data will affect the outcome of data analysis and forecasts. Therefore, homogenizing climate data is necessary in order to obtain reliable regional or global trends, which in result will improve the quality of forecasting. Time series quality control and homogenization is a part of World Meteorological Organization (WMO) guidelines for Global Framework for Climate Services (GFCS) & Climate Services Information System (CSIS), which is to be implemented by National Hydrological & Meteorological Services (NHMS) & Regional Climate Centre (RCC). Here, the concept is to test whether a given data is said to be homogenous over time. In other words, if there is a significant break in trend of particular time series it is classified as inhomogeneous. To identify these breaks in trend one should use metadata to detect any climatic breaks in time series. Upon unavailability of metadata, to identify these breaks there several statistical methods are available but very few methods are available to remove these breaks.

A study on European climate data by Wijngaard *et al*., (2003) shows a detailed approach for testing homogeneity using four different statistical tests namely standard normal homogeneity test, Pettit's tests, Buishand's range test and Von Neumann ration test. These statistical tests were applied on various testing parameters of temperature and rainfall and based on the number of tests rejecting the null hypothesis the results were categorised into three classes, as useful, doubtful and suspect. There is no study in India so far to test the data homogeneity let alone an attempt to correct it. Monthly rainfall and temperature data from all climate stations in state of Maharashtra and Goa is used in this study, which is provided by India Meteorological Department for all available period since 1901. In order to test data homogeneity, we have adopted the Wijngaard *et al*., (2003) approach. Annual rainfall amount and annual mean temperature are used as testing variables in this study. Stations with significant inhomogeneities are needed to be removed and adjustment for homogeneity should be made. Homogeneity correction methods for temperature and rainfall are different in approach. Hanssen and Forland (1994) had shown a method to correct inhomogeneous rainfall data series using a reference station and Morozova and Valente (2012) constructed a procedure to correct temperature data series, both of which we have implemented to correct monthly rainfall and temperature data series for all climate stations in state of Maharashtra and Goa.

## **2. Data and methodology**

Monthly rainfall and temperature data from all climate stations in states of Maharashtra and Goa is collected from India Meteorological Department (IMD) Pune. Table 1 shows stations with their geographical coordinates with data availability statistics. Stations with higher percentage of missing data are not considered for this study. Annual mean temperature and rainfall amount are used as testing variables.

## 2.1. *Statistical methods*

Firstly, original series is visually analysed for detection of any possible breaks such as outliers, which could arrive due to typing or OCR errors. Further, all four homogeneity tests are applied to the data series. The break dates from homogeneity tests are compared with metadata to identify climatic and non-climatic jumps or breaks in series. It is possible that metadata could not provide all the information regarding changes in stations, observation procedure, changes in surroundings etc. and in this study, we could not find any significant breaks which are associated with metadata or any climate forcing data records. After the testing procedure is over, results from testing are collected and stations are categorised to be selected for correction procedure. For a station, break dates that are common for most tests are selected for correction and in case of different break dates for same station, the later one is selected for correction. For correction of data series with significant inhomogeneities, different methods are used for rainfall and temperature, respectively. Rainfall has more variability throughout the year as compared to temperature, because of that; a use of reference station is must for rainfall series correction. An adjustment factor is computed for each station using the method suggested by Hanssen and Forland (1994) and Morozova and Valente (2012) for rainfall and temperature, respectively. Accordingly, adjustment factor is then used to correct rainfall series by ratio's method and temperature series by additive method. Corrected stations again undergo the homogeneity testing procedure to check for any remaining inhomogeneities and same procedure repeated until most significant non-climatic breaks are removed.

## 2.2. *Homogeneity testing*

Four statistical tests are used to test the homogeneity of rainfall and temperature data. Standard normal homogeneity test (SNHT), Pettit's test (PT), Buishand's test (BT) and Von Neumann ratio test (VNR) are used to test the data homogeneity. There are several studies in the past by González-Rouco *et al*., (2001), Tuomenvirta (2001), Kang and Yusof (2012), Ahmad and Deni (2013),

**1% and 5% critical values for** *X***<sup>m</sup> of the Pettit's test as a function of** *n***; values are based on simulation.**

n	20	30	40	50	70	100
1%	-71	133	208	293	488	841
5%	57	107	167	235	393	677

Taxak *et al*., (2014), Guhathakurta *et al*., (2015), Agha *et al*., (2017) and Byakatonda *et al*., (2018) where these tests are used for testing homogeneity of data series. These tests have different sensitivities in different parts of series based on their characteristics as parametric, nonparametric and maximum likelihood, which would help in obtaining results that are more significant. The null hypothesis for these tests is that annual values  $Y_i$  of the testing variable *Y* are independent and identically distributed while the alternate hypothesis for SNHT, PT and BT assume that there is a shift in the mean is present, this significant break in series is considered as inhomogeneous. These tests will identify the year where the break is significant which would be the year where inhomogeneity first occurred. On the other hand, VNR test assume that series is not randomly distributed under alternate hypothesis and the VNR tests does not provide information about a year where break occurs.

The SNHT easily detects breaks near the beginning and at the end of series, whereas to detect breaks in the middle of the series, PT and BR tests are more sensitive (Hawkins, 1977). Pettit's test is based on ranks of the elements in the series, which makes it less sensitive to outliers than other tests like SNHT and Buishand's test, which assume that *Yi* values are normally distributed. The VNR test has sensitivity to departures of homogeneity having characteristics other than stepwise shifts. Annual rainfall amount and mean temperature are testing variables given that  $Y_{i}$ ,  $(i = 1,2, \ldots, n)$  is the testing variable with  $\overline{Y}$  is the mean and s is the standard deviation.

#### 2.2.1*. Standard normal homogeneity test*

SNHT is the most referred homogeneity test in climate studies. A test statistic  $T(d)$  is used to compare the mean of the first *y* years with the last of  $(n - y)$  years can be written as,

$$
T_d = d\,\overline{z}_1^{\,2} + (n-d)\,\overline{z}_2^{\,2},\,d = 1,2,...,n
$$

where,

$$
\overline{z}_1 = \frac{1}{d} \sum_{i=1}^d \frac{(Y_i - \overline{Y})}{s}
$$
 and  $\overline{z}_2 = \frac{1}{n - d} \sum_{i=d+1}^d \frac{(Y_i - \overline{Y})}{s}$ 

#### **TABLE 3**

**1% critical values for the statistic**  $T_0$  **of the single shift SNHT as a function of [calculated from simulations carried out by Jaruskova (1996) and the 5% critical value (Alexandersson and Moberg, 1997)**

n	20	30	40	50	70	100
$1\%$				9.56 10.45 11.01 11.38 11.89		12.32
5%	6.95	7.65	8.10	8.45	8.80	9.15

 $\bar{z}_1$  and  $\bar{z}_2$  are the mean values of  $z_i$  during the first *d* years and last (*n-d*) years, respectively. If a break is located at year D then *T*(*d*) reaches a maximum near the year  $d = D$ . The test statistic  $T_0$  is defines as

$$
T_0 = \max_{1 \le d \le n} T(d)
$$

The probability of rejecting the null hypothesis when  $T_0$  exceeds a certain critical value, which depends on sample size. Critical values for  $T_0$  are given in Table 3 (Khaliq and Quarda, 2007).

#### 2.2.2. *Buishand's range test*

The homogeneity test can be based on the cumulative deviations from the mean or adjusted partial sum which are define as,

$$
S^* = 0 \text{ and } S_d^* = 2 \sum_{i=1}^d (Y_i - \overline{Y}), d = 1, 2, \dots, n
$$

For a homogeneous series, there is no systematic deviations of the  $Y_i$  values with respect to their mean will appear, therefore the values of  $S_d^*$  will fluctuate around zero. Contrarily if a break is present at year *D*, then  $S_d^*$ reaches a maximum (negative shift) or minimum (positive shift) near the year  $d = D$ . Rescaled adjusted partial sums are obtained by dividing the values of  $S_d^*$  by the sample standard deviation *s*. The values are not influenced by any linear transformation; therefore, it is suitable to use the homogeneity test by Q statistics (Buishand, 1982).

$$
Q = \max_{0 \le d \le n} \left| \frac{S_d^*}{s} \right|
$$

In addition, the difference between the maximum and minimum value of the rescaled adjusted partial sums can also be computed by range statistics or R statistics.

$$
R = \frac{\left(\max_{0 \le d \le n} S_d^* - \min_{0 \le d \le n} S_d^*\right)}{s}
$$

*R*



**1% and 5% critical values for Q√n of the Buishand's range test as a function of n (Buishand, 1982); the value of**  $n = 70$  **is simulated** 



#### **TABLE 5**

**1% and 5% critical values for R√n of the Buishand's range test as a function of n (Buishand, 1982)**

n	10	20	30	40	50	-- 100
1%			1.29 1.42 1.46 1.50 1.52			1.55
5%	1.14 1.22 1.24 1.26 1.27 1.29					

Critical values table by Buishand for both homogeneity tests Q√n and R√n are given in Tables  $4&5$ respectively.

## 2.2.3. *Pettit's test*

This test is non-parametric rank test, which does not require any assumption for normality. The test is based on the ranking order and the ranks  $r_1$ ,  $r_2$ ,...  $r_n$  of the  $Y_1$ ,  $Y_2$ ... *Y*<sup>n</sup> are used to calculate the statistics *(Pettit*, *1979)*.

$$
X_d = 2\sum_{i=1}^d r_i - d(n+1) \ d = 1, 2, \dots, n
$$

If the break occurs in year m, then the statistics is maximal or minimal near the year  $d = m$ .

$$
X_m = \max_{1 \le d \le n} |X_d|
$$

Critical values for  $X<sub>m</sub>$  are given in Table 2.

#### 2.2.4. *Von Neumann's ratio test*

The Von Neumann ratio is defined as the ratio of the mean square successive difference between the years to the variance (von Neumann 1941). The test statistic is given as,

$$
N = \frac{\sum_{i=1}^{n-1} (Y_i - Y_{i+1})^2}{\sum_{i=1}^{n} (Y_i - \overline{Y})^2}
$$

#### **TABLE 6**

**1% and 5% critical values of N of the Von Neumann ratio test as a function of n. For N ≤ 50 these values are taken from Owen (1962); for N=70 and N=100 the critical values are based on asymptotic normal distribution of N (Buishand, 1981)**

n 20 30 40 50 70 100			
1\% 1.04 1.20 1.29 1.36 1.45 1.54			
5\% 1.30 1.42 1.49 1.54 1.61 1.67			

This statistic is often used to test the independence of the random variables  $Y_i$ , which are assumed to be successive observations on a stationary Gaussian time series, and sometimes used to test that stationarity of the mean for uncorrelated time. However, this statistic can also be applied in detecting the inhomogeneous series. The series can be considered as homogeneous series if expected value,  $N = 2$ . However, for inhomogeneous series, the value of *N* tends to be lower than 2 (Buishand, 1982). Values greater than 2 implies that the series has rapid variations or oscillations in the mean (Suhaila *et al.*, 2008). This test gives no information about the location of the shift. Table 6 gives critical values for N.

#### 2.3. *Critical values table*

The homogeneity tests are applied to the rainfall and temperature series at 5% level of significance. Critical values for all four tests (Wijngaard *et al.*, 2003) for rainfall, maximum temperature & minimum temperature are given in Tables 7, 8&9 respectively, followed by critical values after homogeneity correction in Tables 10, 11 & 12.

#### 2.4. *Classification of homogeneity tests*

Classification is the second part of homogeneity tests where results from all these four tests are evaluated and combined together. Classification is made into three different categories based on how many tests rejected the null hypothesis (Schönwiese *Rapp,* 1997). These categories are Useful, Doubtful and Suspect and these can be interpreted as follows,

*Class A: Useful -* The series that rejects one or none null hypothesis under four tests at 5% level of significance are considered. The stations under this class can be considered as homogeneous and can be used for further analysis.

*Class B: Doubtful -* The series that rejects two null hypothesis under four tests at 5% level of significance are considered. The stations under this class have

## **Homogeneity test statistics (***p* **value approximation) for rainfall**





## **Homogeneity test statistics (***p* **value approximation) for maximum temperature**

## **Homogeneity test statistics (***p* **value approximation) for minimum temperature**





**Homogeneity Test Statistics (***p* **value approximation) for Rainfall after Correction**

## **TABLE 11**

**Homogeneity test statistics (***p* **value approximation) for maximum temperature after correction**

		Pettitt's Test				<b>SNHT</b>			Buishand's Test				<b>VNR</b>			Class
Station	$\boldsymbol{p}$	K		Year Trend	$\boldsymbol{p}$	T <sub>0</sub>		Year Trend	$\boldsymbol{p}$	$\overline{Q}$		Year Trend	$\boldsymbol{p}$	Ν	<b>Trend</b>	
Aurangabad	0.3184	389	1963	H <sub>0</sub>	0.3191	5.37	1971	H <sub>0</sub>	0.3886	7.73	1963	H <sub>0</sub>	$0.0055$ 1.45		Ha	Useful
Ahmednagar	0.1703	607	1940	H <sub>0</sub>	0.2393	6.25	1981	H <sub>0</sub>	0.2427	9.64	1981	H <sub>0</sub>	$< 0.0001$ 1.02		Ha	Useful
<b>Beed</b>	0.487	86	1990	H <sub>0</sub>	0.0868	7.15	1961	H <sub>0</sub>	0.5114	4.30	1990	H <sub>0</sub>	0.0358	1.41	H <sub>0</sub>	Useful
Dahanu	0.1444	346	1964	H <sub>0</sub>	0.2227	6.17	1961	H <sub>0</sub>	0.1158	9.27	1964	H <sub>0</sub>	0.005	1.40	Ha	Useful
Mormugao	0.0923	196	1969	H <sub>0</sub>	0.0047	11.68 1968		Ha	0.0771	7.87	1969	H <sub>0</sub>	$< 0.0001$ 0.85		Ha	Doubtful
Panjim	< 0.0001	570	2001	Ha	$< 0.0001$ 47.20 2001			Ha	$< 0.0001$ 22.75 2001			Ha	$< 0.0001$ 0.17		Ha	Suspect
Jalgaon	0.1481	366	1988	H <sub>0</sub>	0.124	7.85	1989	H <sub>0</sub>	0.0664 10.20 1988			H <sub>0</sub>	$0.0099$ 1.46		Ha	Useful
Kolhapur	0.0054	521	1989	Ha	0.0267	9.88 1989		Ha	0.0102	12.59 1989		Ha	$0.0008$ 1.21		Ha	Suspect
Colaba	0.0008	1434	1985	Ha	0.0002	19.28 1986		Ha	0.0009	21.05 1986		Ha	$< 0.0001$ 0.97		Ha	Suspect
Santacruz	0.2332	309	1999	H <sub>0</sub>	0.5241	3.84	2001	H <sub>0</sub>	0.3883	6.89	2001	H <sub>0</sub>	0.1546 1.75		H <sub>0</sub>	Useful
Nashik	0.0573	122	1958	H <sub>0</sub>	0.11	7.53	1982	H <sub>0</sub>	0.3719		4.57 1958	H <sub>0</sub>	$0.0427$ 1.43		Ha	Useful
Osmanabad	0.443	59	1981	H <sub>0</sub>	0.1782	5.12	1978	H <sub>0</sub>	0.4096	4.06	1981	H <sub>0</sub>	0.1731 1.64		H <sub>0</sub>	Useful
Baramati	0.1527	154	1963	H <sub>0</sub>	0.0686	7.54	1990	H <sub>0</sub>	0.153	6.60	1963	H <sub>0</sub>	$0.0041$ 1.19		Ha	Useful
Alibag	0.0101	646	1994	Ha	0.0018	14.75 1997		Ha	0.0091	14.12 1994		Ha	$< 0.0001$ 0.85		Ha	Suspect
Ratnagiri	0.0019	1244	2000	Ha	$< 0.0001$ 32.63 2001			Ha	0.0006 21.30 2000			Ha	$< 0.0001$ 0.92		Ha	Suspect
Harnai	0.1782	347	1959	H <sub>0</sub>	0.13		7.55 1957	H <sub>0</sub>	0.0859		9.81 1959	H <sub>0</sub>	0.1567 1.76		H <sub>0</sub>	Useful
Sangli	0.1704	346	1939	H <sub>0</sub>	0.0155	11.05 1939		Ha	0.1477	8.90	1939	H <sub>0</sub>	$0.0008$ 1.27		Ha	Doubtful
Satara	0.0775	161	1999	H <sub>0</sub>	0.1514		5.86 2008	H <sub>0</sub>	0.2592		5.68 2006	H <sub>0</sub>	$0.0004$ 0.99		Ha	Useful
Mahabaleshwar 0.0324		530	1948	Ha	0.0159	11.10 1940		Ha	0.0525	11.14 1948		H <sub>0</sub>	0.0296 1.58		Ha	Suspect
Devgad	0.3327	174	1993	H <sub>0</sub>	0.0155	10.00 1993		Ha	0.3388		6.13 1993	H <sub>0</sub>	0.0129 1.39		Ha	Doubtful
Vengurla	0.0877	254	2000	H <sub>0</sub>	0.0155 10.10 2001			Ha	0.1183		7.99 2000	H <sub>0</sub>	0.0527 1.56		H <sub>0</sub>	Useful
Solapur	0.0632	918	2000	H <sub>0</sub>	0.0109	12.09 2001		Ha	0.0734	13.20 2000		H <sub>0</sub>	$< 0.0001$ 1.19		Ha	Doubtful
Jeur	0.1731	185	1981	H <sub>0</sub>	0.2896	5.13 1981		H <sub>0</sub>	0.1385	7.28 1981		H <sub>0</sub>	$0.0053$ 1.27		Ha	Useful

## **Homogeneity Test Statistics (***p* **value approximation) for Minimum Temperature after Correction**



inhomogeneities and can be used for further analysis after critical inspection.

*Class C : Suspect -* The series that rejects three or all null hypothesis under four tests at 5% level of significance are considered. The stations under this class have significant inhomogeneities and cannot be used for further analysis.

In this study stations with class 'doubtful' and 'suspect' are excluded from further climate trend analysis until they are thoroughly inspected for inhomogeneities and necessary corrections are made to adjust the data for homogeneity.

#### 2.5. *Homogeneity correction*

The results from homogeneity tests are analysed and based on categorisation, stations with class 'doubtful' and 'suspect' are selected for correction procedure. Stations with category 'useful' are used as reference station. Adjustment factor is calculated for both rainfall and temperature, which is needed to correct the original data series.

## 2.5.1*. Selection of reference station*

For selection of reference station, influence of spatial variation and eventual inhomogeneities in the reference

## **Homogeneity tests before and after correction summary**



## **Homogeneity test results**



## **TABLE 15**

## **Homogeneity Test Regression Statistics Results**





**Figs. 1(a-f).** Rainfall series before and after homogeneity correction

series will be reduced by using more than one reference series. The criteria for selecting reference series stations must satisfy conditions such as, it must be at minimum distance from test station, it should have same climate properties as test station, the availability of data must be of same or more period as test station and it the data of reference series should highly correlate with test series data. Based on our testing results, we have decided to use only single reference station because the aforementioned selection criteria was difficult to achieve for multiple reference station as stations with 'useful' categories were at far distance from each other having different data availability period and difference in climate patterns.

#### 2.5.2. *Adjustment factor*

After identifying the break dates in test series, an adjustment for homogeneity is made by calculating adjustment factor. This adjustment factor is then used to correct the data before the break date by multiplying original rainfall data or by adding in original temperature data. As mentioned earlier, rainfall and temperature series have different correction procedure; the details of which are given below.

#### 2.5.3. *Rainfall series correction*

Generally, ratio in test and reference series q in a specific year *i* is given as,

$$
q_i = \frac{f(P_i)}{g(Q_{ij}, j=1, K_i)}
$$

Here,  $i$  is the function of precipitation  $P_i$  at the test station,  $Q_{ij}$  is the precipitation at the  $j<sup>th</sup>$  reference station, and g is the function of precipitation at all the  $k_i k_i$ reference stations. Functions of *f* and *g* are defined as,

$$
f(P_i) = \frac{P_i}{\overline{P}} \text{ and } g(Q_{ij}, j = 1, K_i) = \frac{\sum_{j=1}^{k_i} V_j(Q_{ij} / \overline{Q}_j)}{\sum_{j=1}^{k_i} V_j}
$$

where,  $\overline{P}$  and  $\overline{Q}_i$  are mean precipitation throughout the observation period for the test station and the  $j<sup>th</sup>$ reference station, respectively, and  $k_i$  is the number of reference stations in the *i*th year of observation.  $V_i$  is a weight factor for reference station *j* which is defined by the correlation coefficients between the test series and  $j^{\text{th}}$ reference series (Peterson *et al*., 1998).

Inhomogeneous series is adjusted by multiplying the precipitation values for the period before the break date with an adjustment factor,

$$
AF = \frac{q_a}{q_b}
$$

Here,  $q_a$  and  $q_b$  are mean values of  $q_i$  after and before break year, respectively.



**Figs. 2(a-l).** Maximum temperature series before and after homogeneity correction

#### 2.5.4. *Temperature series correction*

First, select a time interval  $\Delta t$  before and after break year  $t<sub>break</sub>$ . Next, for each interval  $\Delta t$ , take average of each individual month separately and let us call it  $t<sub>break</sub>$  and  $t<sub>after</sub>$ Now, calculate the correction *dT* for each month separately by taking difference of these averages *t*before and *t*after.

After calculating *dT*, make smoothing 12 monthly *dT* by taking 3-month adjacent running average to achieve reasonable variation of *dT* throughout the year. By doing this for each of 12 months individually, we will get our adjustments factor AF for each month. Corrected series can be achieved by adding this adjustment factor in original temperature series for the period before the break year.



**Figs. 3(l-w).** Maximum temperature series before and after homogeneity correction

## **3. Results and discussions**

Total 30 rainfall and temperature stations were tested for data homogeneity. For rainfall series, homogeneity test results show 80% as useful, 7% as doubtful and 13% as suspect (Table 7), For maximum temperature series, homogeneity test result shows 17% useful, 7% doubtful and 76% suspect stations (Table 8) and for minimum temperature series, homogeneity test result shows 21%

useful, 10% doubtful and 69% suspect stations (Table 9). Use of various test found to be useful in detecting break year because of their characteristics as parametric, nonparametric, rank, *etc*. It is also observed that, break year detected by most tests were same for a respective station. The stations with doubtful and suspect category were taken up for correction by using useful category stations as reference series for rainfall. After applying correction procedures, the results show all 100% rainfall stations as



**Figs. 4(a-l).** Minimum temperature series before and after homogeneity correction

'useful' (Table 10), for maximum temperature series the results are 62% as 'useful', 17% as 'doubtful' and 21% as 'suspect (Table 11), while for minimum temperature series these results are 69% as 'useful', 21% as 'suspect' and 10% as 'doubtful' (Table 12). The 'useful' category stations may be used further in climate and trend analysis studies. Table 13 shows the summary of homogeneity tests before & after correction for all three parameters,

whereas Table 14 shows the consolidated homogeneity test results.

Table 15 shows the linear regression statistics (change in trend) for all three parameters before and after applying the homogeneity correction. Fig. 1 shows the linear trend line for rainfall and it shows that 33% stations show-decreasing trend after applying homogeneity



**Figs 5(m-v).** Minimum temperature series before and after homogeneity correction

correction. Figs. 2&3 shows the linear trend line for maximum temperature and it shows that 9% of the stations show the increasing trend whereas 13% of the total stations show the decreasing trend after applying homogeneity correction. Figs. 4&5 shows the linear trend line for minimum temperature, which shows that 5% of the stations show the increasing trend whereas 18% of total stations show the decreasing Trend. For majority of stations, it has been seen then there is an overall reduction in magnitude of a trained to applying the homogeneity correction for all the rainfall and temperature series.

One of the biggest consequences of biased trends are reflected in the climate change studies. Several climate studies show the increasing trend in temperature and decreasing trend in rainfall over last century, concluding as an effect of global warming. Such studies have failed to incorporate effects of non-climatic jumps in data and the inhomogeneities arrived due to various external factors are also not considered. Here, the trends in series before and after correction will also help to determine if there is an effect of climate change or not.

## **4. Conclusions**

In the present study, homogeneity tests are proven most important part of climate and trend analysis studies. Different characteristics of these tests not only helped in identifying significant non-climatic jump in a time series but also location of a break period. The correction

methods used are proved to be very effective in removing non-climatic jump from time series and making it homogeneous.

## *Acknowledgement*

Authors are thankful to India Meteorological Department for providing resources and research platform for carrying out this work.

*Disclaimer* : The contents and views expressed in this research paper/article are the views of the authors and do not necessarily reflect the views of the organizations they belong to.

#### **References**

- Agha, O. M. A. M. Bağçacı, S. Ç. and Şarlak, N., 2017, "Homogeneity analysis of precipitation series in North Iraq", *IOSR Journal of Applied Geology and Geophysics*, **5**, 3, 57-63.
- Ahmad, N. H. and Deni, S. M., 2013, "Homogeneity test on daily rainfall series for Malaysia", *MATEMATIKA* : *Malaysian Journal of Industrial and Applied Mathematics*, **29**, 141-150.
- Alexandersson, H. and Moberg, A., 1997, "Homogenization of Swedish temperature data. Part I : Homogeneity test for linear trends", *International Journal of Climatology : A Journal of the Royal Meteorological Society*, **17**, 1, 25-34.
- Buishand, T. A., 1982, "Some methods for testing the homogeneity of rainfall records", *Journal of hydrology*, **58**, 1-2, 11-27.
- Byakatonda, J. Parida, B. P. Kenabatho, P. K. and Moalafhi, D. B., 2018, "Analysis of rainfall and temperature time series to detect longterm climatic trends and variability over semi-arid Botswana", *Journal of Earth System Science*, **127**, 2, 25.
- González-Rouco, J. F. Jiménez, J. L. Quesada, V. and Valero, F., 2001, "Quality control and homogeneity of precipitation data in the southwest of Europe", *Journal of climate*, **14**, 5, 964-978.
- Guhathakurta, P. Rajeevan, M. Sikka, D. R. and Tyagi, A., 2015, "Observed changes in southwest monsoon rainfall over India during 1901-2011", *International Journal of Climatology*, **35**, 8, 1881-1898.
- Hanssen-Bauer, I. and Førland, E. J., 1994, "Homogenizing long Norwegian precipitation series", *Journal of Climate*, **7**, 6, 1001-1013.
- Hawkins, D. M., 1977, "Testing a sequence of observations for a shift in location", *Journal of the American Statistical Association*, **72**, 357, 180-186.
- Jaruskova, D., 1996, "Change-point detection in meteorological measurement", *Monthly Weather Review*, **124**, 7, 1535.
- Kang, H. M. and Yusof, F., 2012, "Homogeneity tests on daily rainfall series", *Int. J. Contemp. Math. Sciences*, **7**, 1, 9-22.
- Khaliq, M. N., and Ouarda, T. B., 2007, "On the critical values of the standard normal homogeneity test (SNHT)", *International Journal of Climatology : A Journal of the Royal Meteorological Society*, **27**, 5, 681-687.
- Morozova, A. L. and Valente, M. A., 2012, "Homogenization of Portuguese long-term temperature data series : Lisbon, Coimbra and Porto", *Earth System Science Data*, **4**, 187-213.
- Owen, D. B., 1962, "Handbook of statistical tables", *Addison Wesley: Reading, UK*.
- Peterson, T. C., Easterling, D. R., Karl, T. R., Groisman, P., Nicholls, N., Plummer, N. and Parker, D., 1998, "Homogeneity adjustments of in situ atmospheric climate data: a review", *International Journal of Climatology: A Journal of the Royal Meteorological Society*, **18**, 13, 1493-1517.
- Pettitt, A. N., 1979, "A non‐parametric approach to the change‐point problem", *Journal of the Royal Statistical Society: Series C (Applied Statistics)*, **28**, 2, 126-135.
- Schönwiese, C. D., and Rapp, J., 1997, "Climate trend atlas of Europe based on observations 1891-1990", *Springer Science & Business Media*.
- Suhaila, J., Deni, S. M., and Jemain, A. A., 2008, "Detecting inhomogeneity of rainfall series in Peninsular Malaysia", *Asia-Pacific Journal of Atmospheric Sciences*, **44**, 4, 369-380.
- Taxak, A. K. Murumkar, A. R. and Arya, D. S., 2014, "Long term spatial and temporal rainfall trends and homogeneity analysis in Wainganga basin, Central India", *Weather and Climate Extremes*, **4**, 50-61.
- Tuomenvirta, H., 2001, "Homogeneity adjustments of temperature and precipitation series-Finnish and Nordic data", *International Journal of Climatology: A Journal of the Royal Meteorological Society*, **21**, 4, 495-506.
- Von Neumann, J., 1941, "Distribution of the ratio of the mean square successive difference to the variance", *The Annals of Mathematical Statistics*, **12**, 4, 367-395.
- Wijngaard, J. B. Klein Tank, A. M. G. and Können, G. P., 2003, "Homogeneity of 20<sup>th</sup> century European daily temperature and precipitation series", *International Journal of Climatology: A Journal of the Royal Meteorological Society*, **23**, 6, 679-692.